Responses of heat stress to the climate warming from the perspective of temperature and humidity

Jiacan Yuan^{1,3}

Michael I. Stein², Robert E. Kopp³, Dawei Li^{3,4}

- 1. Department of Atmospheric and Oceanic Sciences, Fudan University, CN
- 2. Department of Statistics, University of Chicago, US
- 3. Department of Earth and Planetary Sciences, Rutgers University, US
- 4. Department of Geosciences, University of Massachusetts Amherst, US



Motivation





Motivation

Impact of extreme events during 2001-2010 compared with 1991-2000



700 000

Increase in surface air temperature (T)



Multi-model ensemble of global-mean surface air temperature deviating from 1986-2005 (The Climate Science Special Report Fig.1.4)

39-model average of surface air temperature in 2081-2100 deviating from 1986-2005 (IPCC AR5 Fig. 12.11)



Increase in extreme temperature



Humidity is another important contributor to heat stress

People feel hotter when relative humidity is higher even the temperature is the same



At higher elevations in the mountains



In the desert



High humidity





In a rainforest or jungle



At the beach or near the ocean

It's harder for sweat to evaporate because the air is already so moist. The body doesn't cool down as easily. That's why it feels hotter.



US National Weather Service

Over land, specific humidity increases while relative humidity (RH) decreases



IPCC AR5 (Fig. 11.14)



Outline

Part I Emulate the evolving distribution of relative humidity conditional on daily maximum temperature in a warming climate

Part II Responses of Wet-Bulb Globe Temperature to climate warming





Outline

Part I Emulate the evolving distribution of relative humidity conditional on daily maximum temperature in a warming climate

Part II Responses of Wet-Bulb Globe Temperature to climate warming





Common assumptions for evolving marginal distribution in a warming climate





Hansen et al. (2012)



Evolving marginal distributions in a real case



July in Chicago (CESM LENS RCP8.5)



Data: 35 members from CESM Large Ensemble 6-hourly data RCP8.5

Evolving Joint Distributions between Tmax and Relative Humidity (RH)

New York City July



12

Evolving Joint Distributions between Tmax and Relative Humidity (RH)



- > NYC: New York City, humid temperate climate
- CHI: Chicago, hot-summer humid continental climate
- PHX: Phoenix, hot-desert climate
- > NOLA: New Orleans, humid subtropical climate





Enthalpy(Megajoule/kg)



How to predict the evolving distribution of RH conditional on Tmax?



Quantile Regression

Multi-linear regression: $\widehat{Y} = X \widehat{\beta}$

Quantile regression: $\widehat{Y}_{\tau} = X \hat{\beta}_{\tau}$

Advantages of quantile regression :

- Without assuming any particular statistical form for the distribution
- Allows any changes in shape of distribution





Emulating the distribution of RH on Tmax through Quantile Regression

$$\widehat{RH}_{\tau}(Tmax) = \theta + \gamma (Tmax - T_0)_+ + \sum_{j=1}^m \eta_j K_j(Tmax)$$

$$\underbrace{\text{Kink function}}_{1} \quad \text{Cubic Spline function}$$

$$\underbrace{(Tmax - T_0)_+}_{1} = \begin{cases} Tmax - T_0, & Tmax \ge T_0 \\ 0, & Tmax < T_0 \end{cases}$$



Quantile regression solves the following minimization problem:

$$\tau \sum_{i=1}^{n} \left(RH_i - \widehat{RH}_{\tau}(Tmax_i) \right)_{+} + (1-\tau) \sum_{i=1}^{n} \left(\widehat{RH}_{\tau}(Tmax_i) - RH_i \right)_{+}$$





Dots: raw data from CESM LENS





Model evaluation

Empirical inverse quantiles + cross validation of the RH data in a small Tmax interval (1 degree C)





99 quantiles from 0.01 to 0.99





the solid lines mark the 0.025 (lower line) and 0.975 (upper line) of the binomial distribution the counts should follow if the model is accurate





At fixed temperature (e.g. 95th quantile in 1990-2005)

Increases in RH will amplify the heat stress in a future day when climate gets warmer.





Heat Index: a metric measures intensity of heat stress considering both temperature and relative humidity (Rothfusz, 1990)



At fixed quantile of temperature (e.g. 95th quantile in each period) In a warming climate, despite a modest decrease in RH, heat stress will tend to increase faster than temperature alone would indicate







Outline

Part I Emulate the evolving distribution of relative humidity conditional on daily maximum temperature in a warming climate

Part II Responses of Wet-Bulb Globe Temperature to climate warming





Wet bulb globe temperature

1. It is measurable



2. Has been widely used in health guidelines

Table 1. ISO 7243: WBGT reference values

Metabolic rate (Wm ⁻²)	WBGT reference value	
	Acclimatized (°C)	Not acclimatized (°C)
Resting $M < 65$	33	32
65 < M < 130	30	29
130 < M < 200 ~ walki	ing 28	26
200 < M < 260	25 (26)*	22 (23)*
M > 260	23 (25)*	18 (20)*

The values given have been established allowing for a maximum rectal temperature of 38°C for the persons concerned.

*: Figures in brackets refer to sensible air movement.

Parsons (2006)

25

Wet bulb globe temperature

Wet-bulb globe temperature (WBGT):

 $WBGT = 0.7 \times T_w + 0.2 T_g + 0.1 \times T_a$

Simplified WBGT (appropriate for shaded conditions):

 $WBGT^* \approx 0.7 \times T_w + 0.3 \times T_a$



Wet bulb temperature

- > Empirical relationship based on psychrometric chat (Stull 2011) ~ limited by ranges of observations
- > Wet bulb potential temperature (DAVIES-JONES 2008) ~ designed for severe weather forecast
- \succ Isobaric wet bulb temperature (Li et al. 2020) ~ the cooling of body through perspiration



$$C_{pa}T_w + L_v q_s(T_w) = C_{pa}T + L_v q$$

Enthalpy of saturated moist air

Enthalpy of initial moist air

Data and methods





What is the historical heat stress like, measured by WBGT?





25 26 27 28 29 30 31 32 33

)

What is the implication for heat stress, measured by WBGT, under the different level of global warming?



Increasing GMST



31 °C ~ the peak of 1995 Chicago Heatwave

- 33 °C ~ humans maintain a normal core body temperature at resting
- 35 °C ~ maximum value calculated from historical observations



Increasing GMST



31 °C ~ the peak of 1995 Chicago Heatwave

- 33 °C ~ humans maintain a normal core body temperature at resting
- 35 °C ~ maximum value calculated from historical observations

Changes in Frequency of Exceedance with $\Delta GMST$ (Selected Sites)





World Population Exposure to Extreme WBGT



Key Takeaways

- > Humidity is an important contributor to the impact of heat stress:
 - > At fixed temperature, increases in RH will amplify the heat stress in a future day
 - In a warming climate, despite a modest decrease in RH, heat stress will tend to increase faster than temperature alone would indicate
- The frequency of heat-humidity extremes, measured using WBGT, will increase dramatically in response to global warming
- ≻ The benefit of limiting ∆GMST to 1.5°C rather than 2°C is evident in reducing world

population exposure to life threatening heat stress.





FDU-IRDR-ICOE-RIG-WECEIPHE



Risk Interconnectivity and Governance on WEather/Climate Extremes Impact and Public HEalth

极端天气/气候事件与人体健康风险互联与治理国际卓越中心



Recruitment

Position: Ph.D candidate, Postdoc

Research field: impact of weather/climate extremes on human health



Contact: Jiacan Yuan (袁嘉灿) jcyuan@fudan.edu.cn



Thank You!

- J. Yuan*, M. I. Stein, and R. E. Kopp (2020). The evolving distribution of relative humidity conditional upon daily maximum temperature. Journal of Geophysical Research-Atmospheres. https://doi.org/10.1029/2019JD032100
- D. Li*, J. Yuan*, and R. E. Kopp (2020). Escalating Global Exposure to Compound Heat-Humidity Extremes with Warming. Environmental Research Letters. https://doi.org/10.1088/1748-9326/ab7d04



Convert sea level pressure to surface pressure

$$P = \rho RT \qquad (1)$$

$$dP = -\rho g dz \qquad (2)$$

$$dT = \Gamma dz \qquad (3)$$

$$(1) + (2): \qquad dP = -\frac{g}{RT} P dz \qquad \longrightarrow \qquad dlnP = -\frac{g}{RT} dz$$

$$(3) + (4): \qquad dlnP = -\frac{g}{R\Gamma} \times \frac{dT}{T} = -\frac{g}{R\Gamma} dlnT$$

$$ln\left(\frac{Ps}{Psl}\right) = -\frac{g}{R\Gamma} ln\left(\frac{Ts}{Tsl}\right)$$

$$Ps = Psl \times \left(1 - \left(\frac{\Gamma z}{Ts}\right)\right)^{\frac{g}{R\Gamma}}$$

(4)