

An Automated Chamber Network for Evaluation of Carbon Budget of Asian Terrestrial Ecosystems



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Center for Global Environmental Research (CGER)

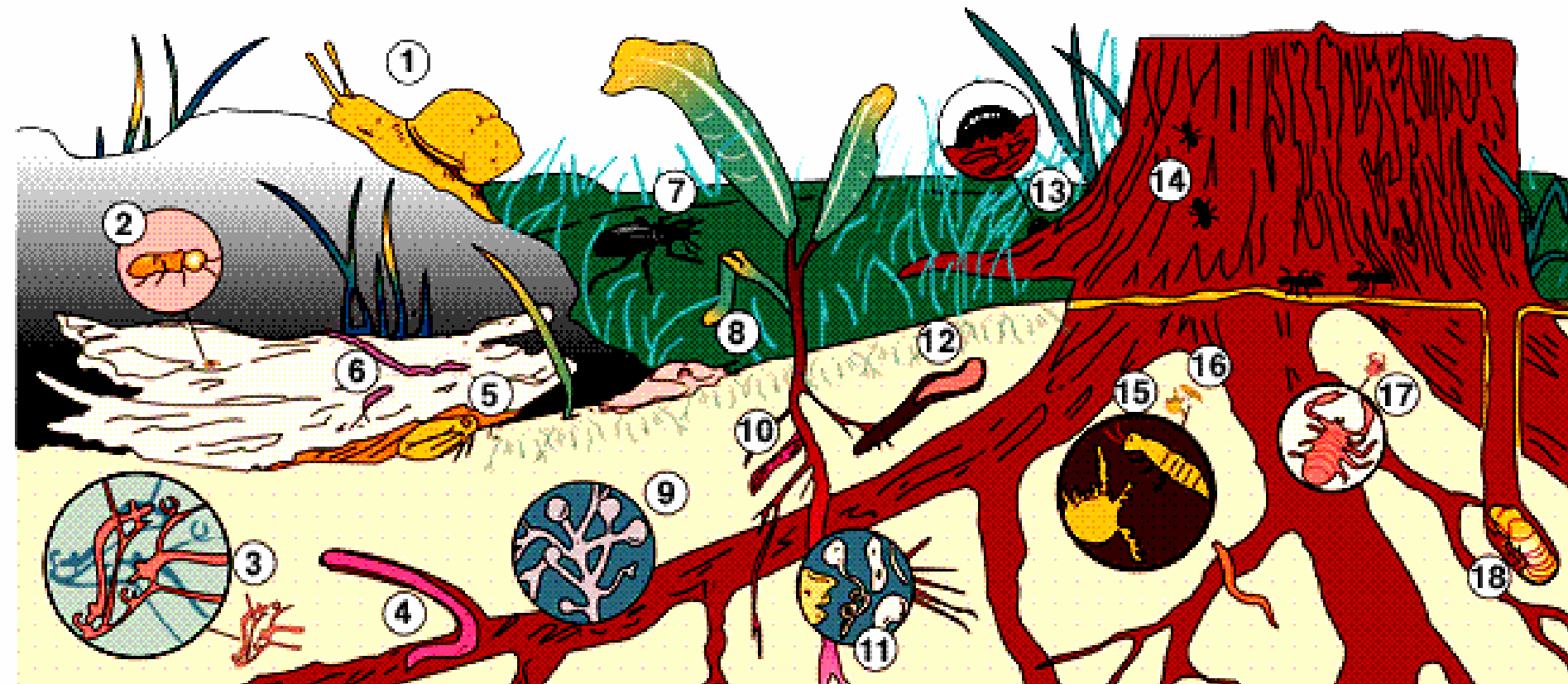
National Institute for Environmental Studies (NIES)

Outline

1. **What** (is soil respiration (R_s))
2. **Why** (study on R_s)
3. **How** (chamber network)
4. **Find** (results)
5. **Future** (plan)

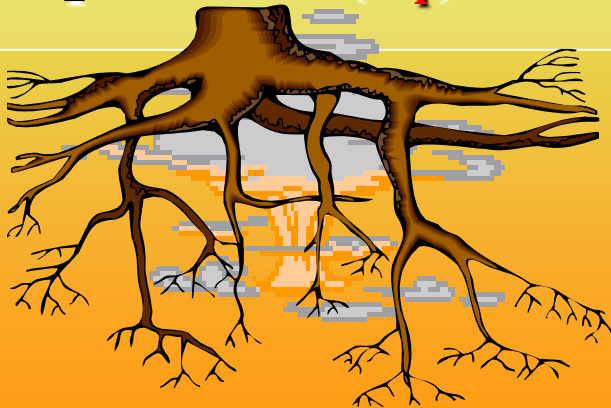
Soil Ecosystem

includes roots and numerous consumer organisms

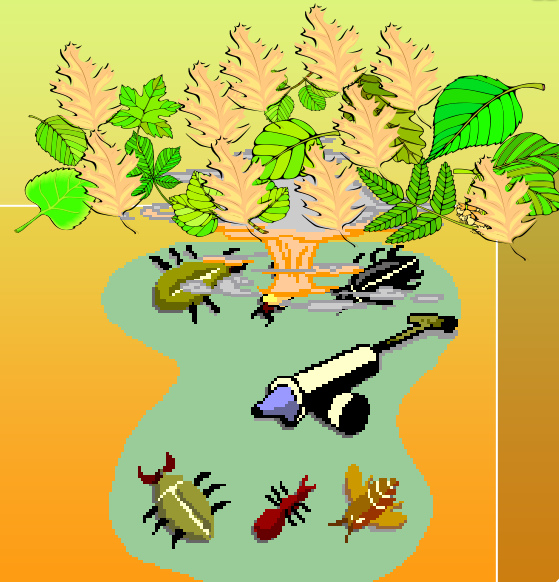


Soil Respiration (R_s)

Autotrophic (Root)
respiration (R_r)



Heterotrophic
respiration (R_h)



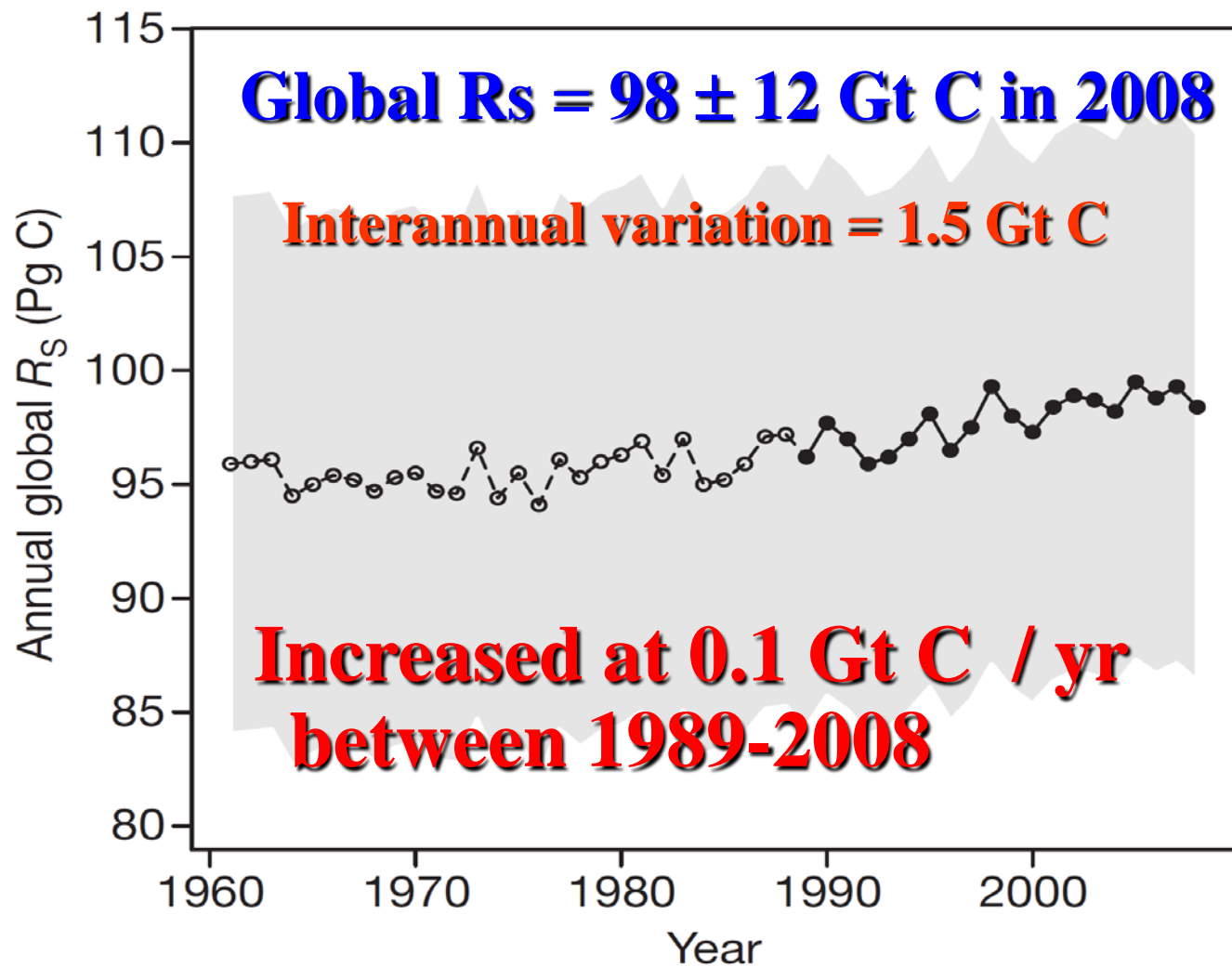


Figure 2 | Estimated annual global R_s . The dashed line indicates results outside the time period covered by main data set, S1 (1989–2008), but within the period covered by the entire R_s database, S0 (1961–2008), and should be considered speculative. The grey region shows the standard deviation of the Monte Carlo simulations ($N = 1,000$). **Nature, 464: 579-582 (25 March 2010)**

Global Heterotrophic Respiration (Rh)

Potter & Klooster 1998 (*GBC*); IPCC2007

by CASA model

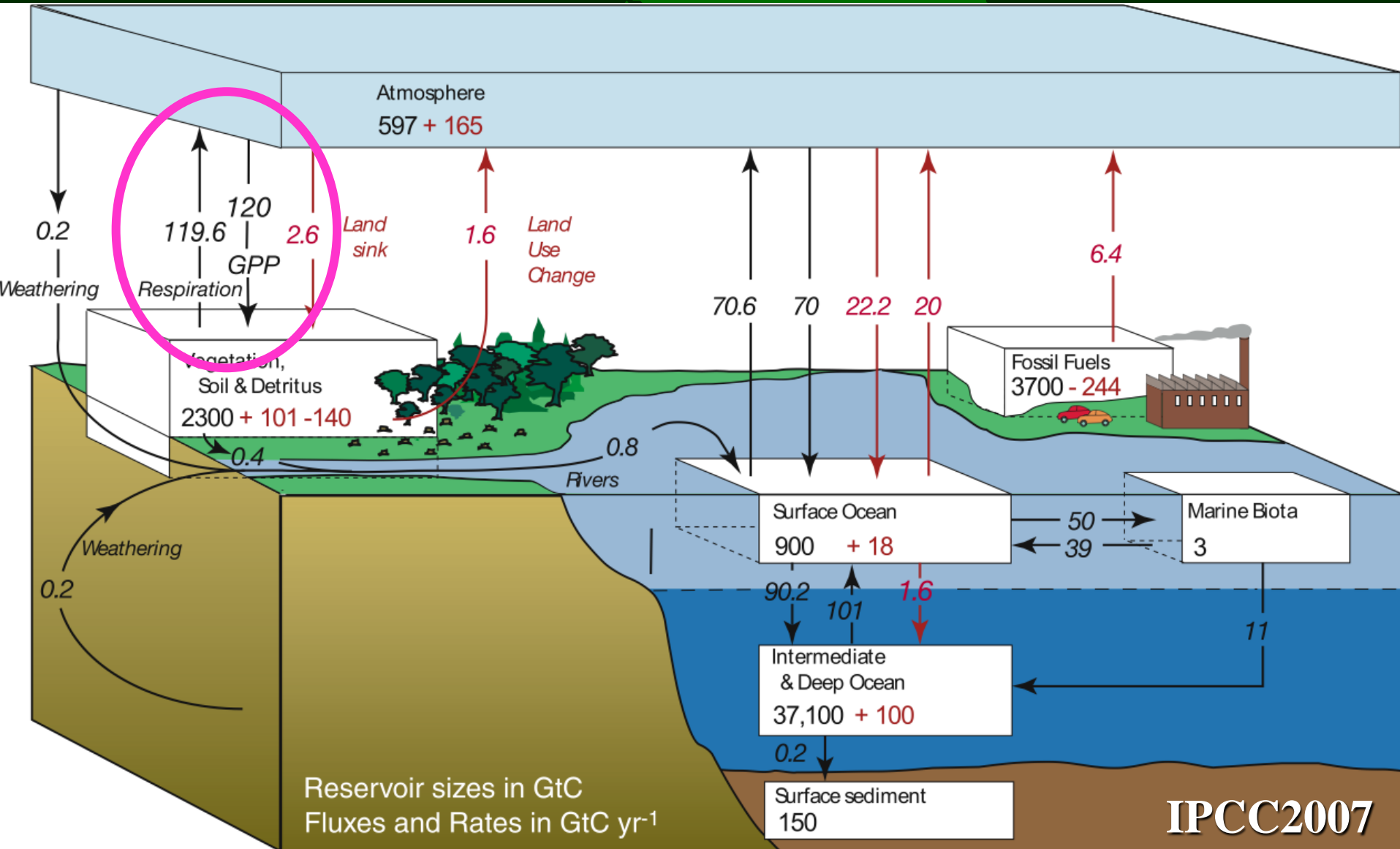
Global Rh = 71% of Rs (**69.6 Gt C y⁻¹**)

10 times of fossil fuel emission (7.2 GtC y⁻¹)

30 times of land C sink (2.6 GtC y⁻¹)

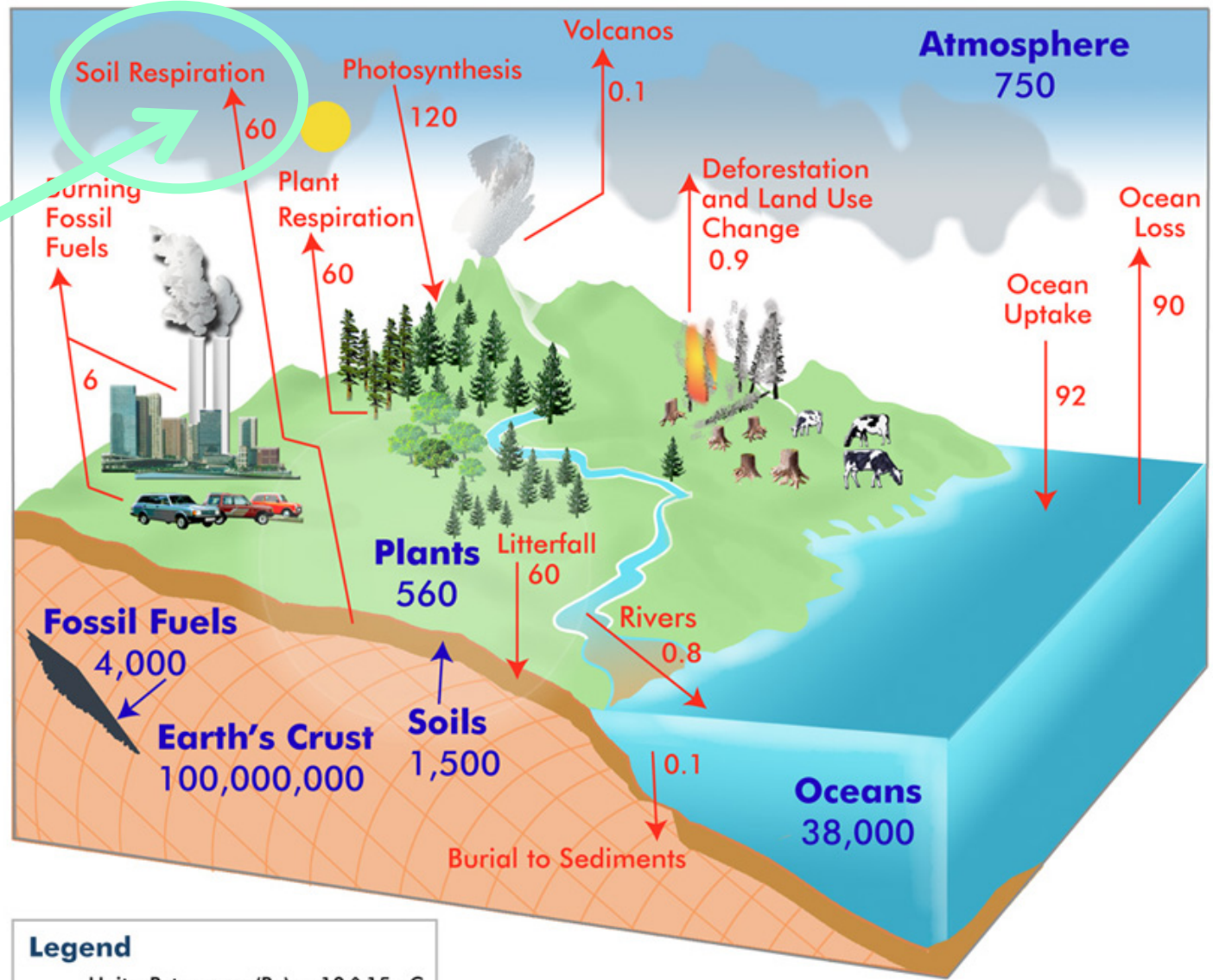
☞ **Plays a key role in global carbon cycle**

Global Carbon Balance



Global Carbon Cycle

Rh = 70



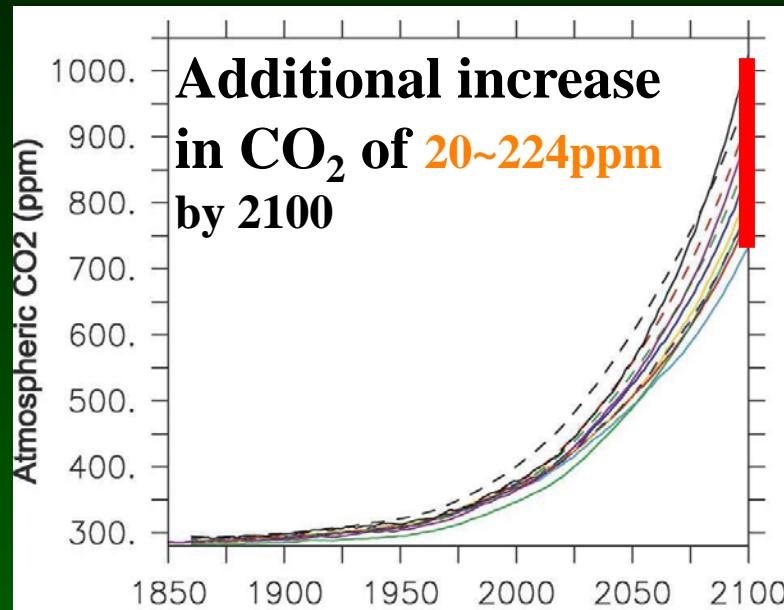
Legend

Units: Petagrams (Pg) = 10¹⁵ gC

- Pools: Pg
- Fluxes: Pg/year

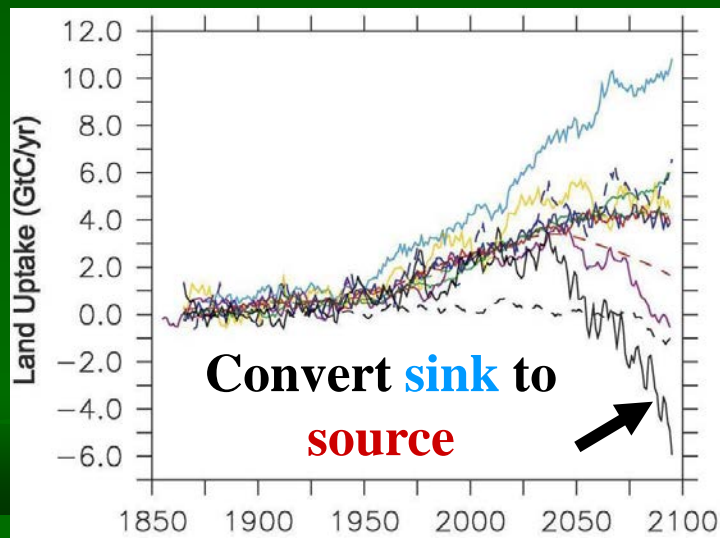
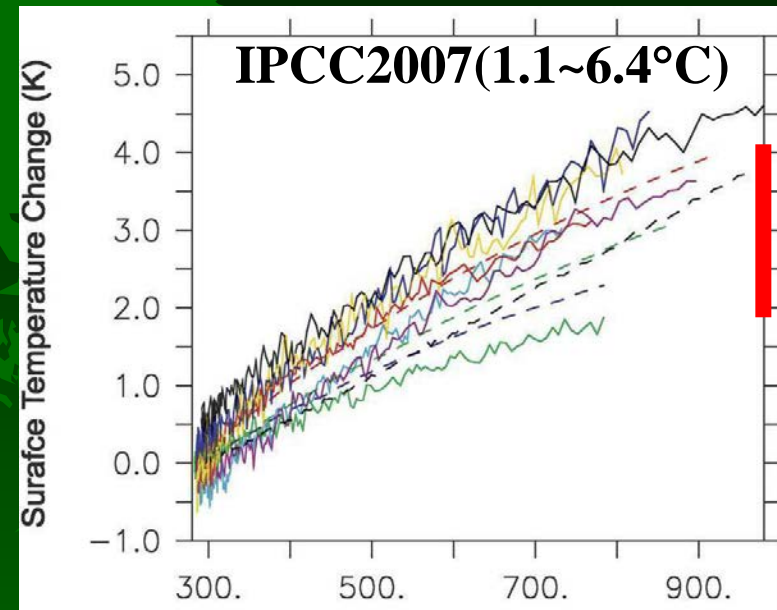
Feedback of Heterotrophic Respiration to Global Warming

(Friedlingstein et al. 2006; IPCC 2007)



Vicious cycle

0.1~1.5°C

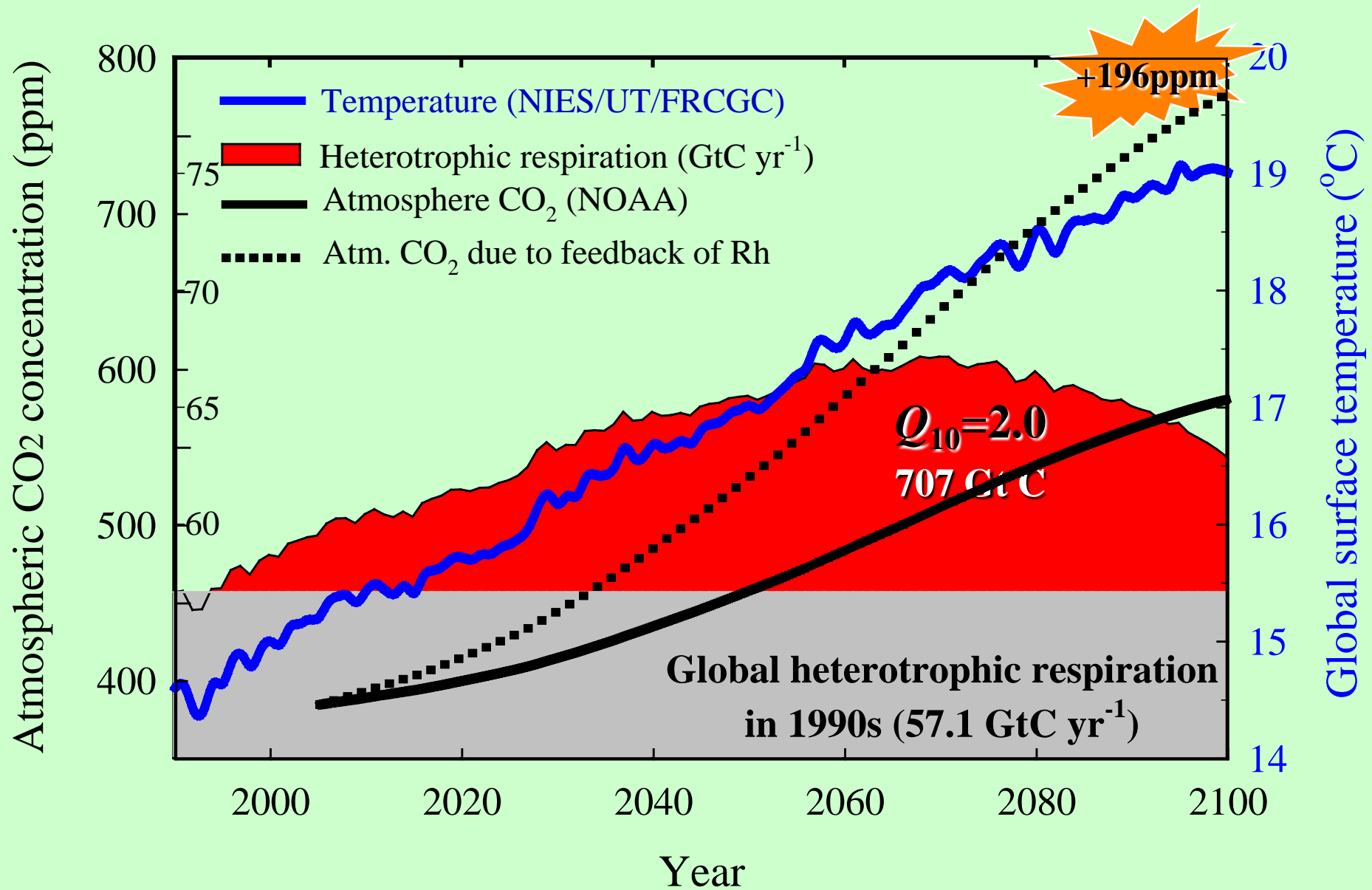


Positive feedback

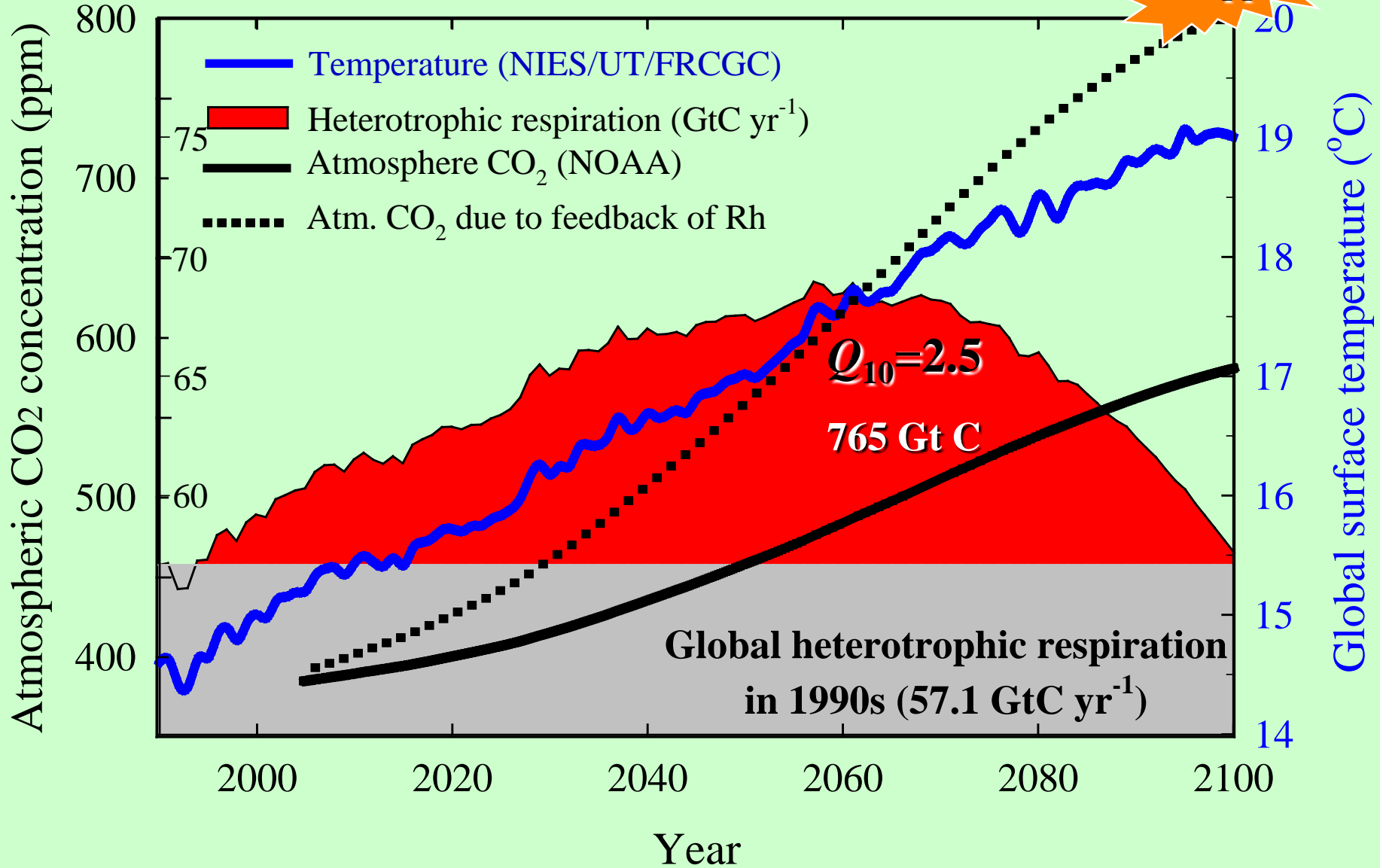
Heterotrophic respiration

$$Q_{10}=1.1\sim 2.2$$

Feedback of Heterotrophic Respiration to Global Warming



Feedback of Heterotrophic Respiration to Global Warming



Open Questions

- **With global warming, will Asian terrestrial ecosystems continuous be carbon sink?**
- **or potentially convert to carbon source?**

Eddy Covariance for Soil Respiration



Bare soil site

Eddy Covariance

A photograph of an eddy covariance flux tower in a forest. The tower is a vertical metal pole with various instruments attached. At the top, there is a sonic anemometer and a gas analyzer. Below that, there are two horizontal arms, each with a sensor. A white control box is mounted on the pole. The tower is surrounded by dense green foliage, including ferns and trees. A tree trunk is visible on the right side of the frame.

- ◆ Daytime: underestimation
- ◆ Nighttime: overestimation

Chamber Methods

Static Chamber

**Alkali
Soda lime**

Syringe

Dynamic Chamber

**Steady
state**

**Non-steady
state**

Open-top

Closed-top

Commodity

Static Chamber Technique



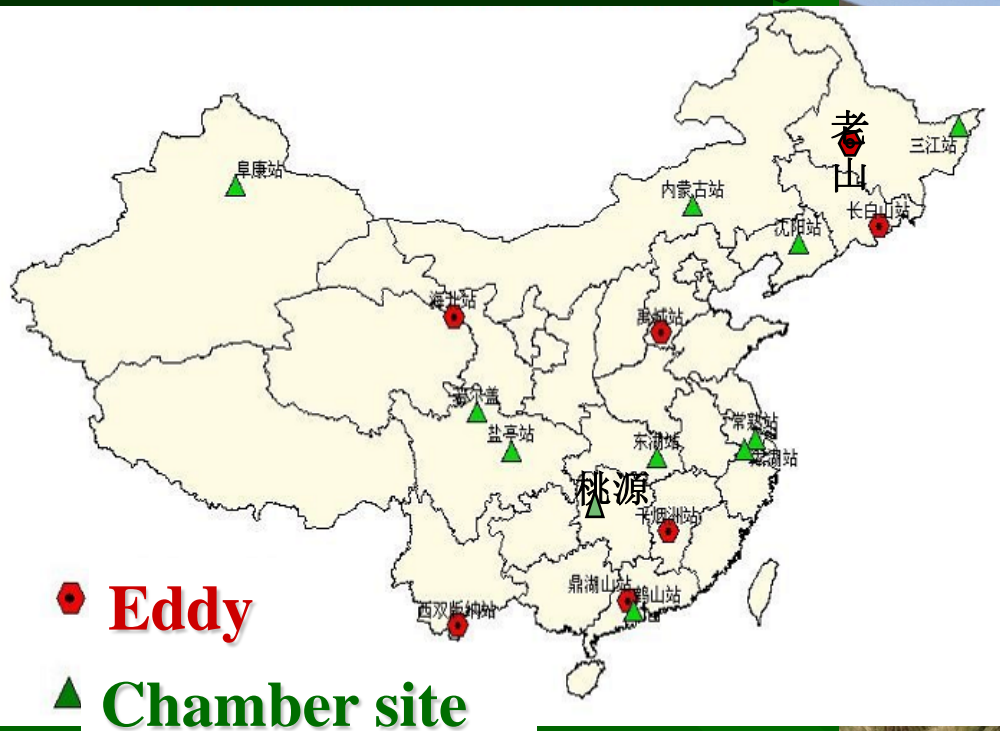
Alkali Absorption Method

Science Camp for High School Students



July 27-29, 2004 in Tomakomai

ChinaFlux



Manual static chamber-GC method

(b)

Gas sampling



Gas chromatography



Sampling frequency: once a week

Time: 9:00-12:00

御見積書

No. IN0-094

年 月 日

独立行政法人
国立環境研究所 御中

下記の通り御見積申し上げます。

会社

製造会社名： 株式会社
納期： 受注後 日の予定
御支払条件： 現金払
御見積有効期間： 60 日

4-2323
)
9-0811
0-7702
8-0561

御見積金額合計 ¥17,662,890._

型式	品名	数量	価格	
			単価	合計
LI-8100	全自動土壌呼吸測定システム	1 台	35,000._	2,850,000._
LI-8150-16	マルチプレクサー (LI-8100用)	1 台	12,000._	3,400,000._
8100-104/16	ロングタームチャンバー16個セット	1 セット	11,200,000._	11,200,000._
	小計			17,450,000._
	値引			- 628,200._
	合計			16,821,800._
	消費税			841,090.
	総計			¥17,662,890._

上記見積中には据付、試運転、整備又は調整等に関する費用は原則として含まれて居りません。御契約に際し下記による訂正御承認賜り度御願ひ申し上げます。

担当者



Soil CO₂ Gradient Technique

NDIR Gas Analyzer



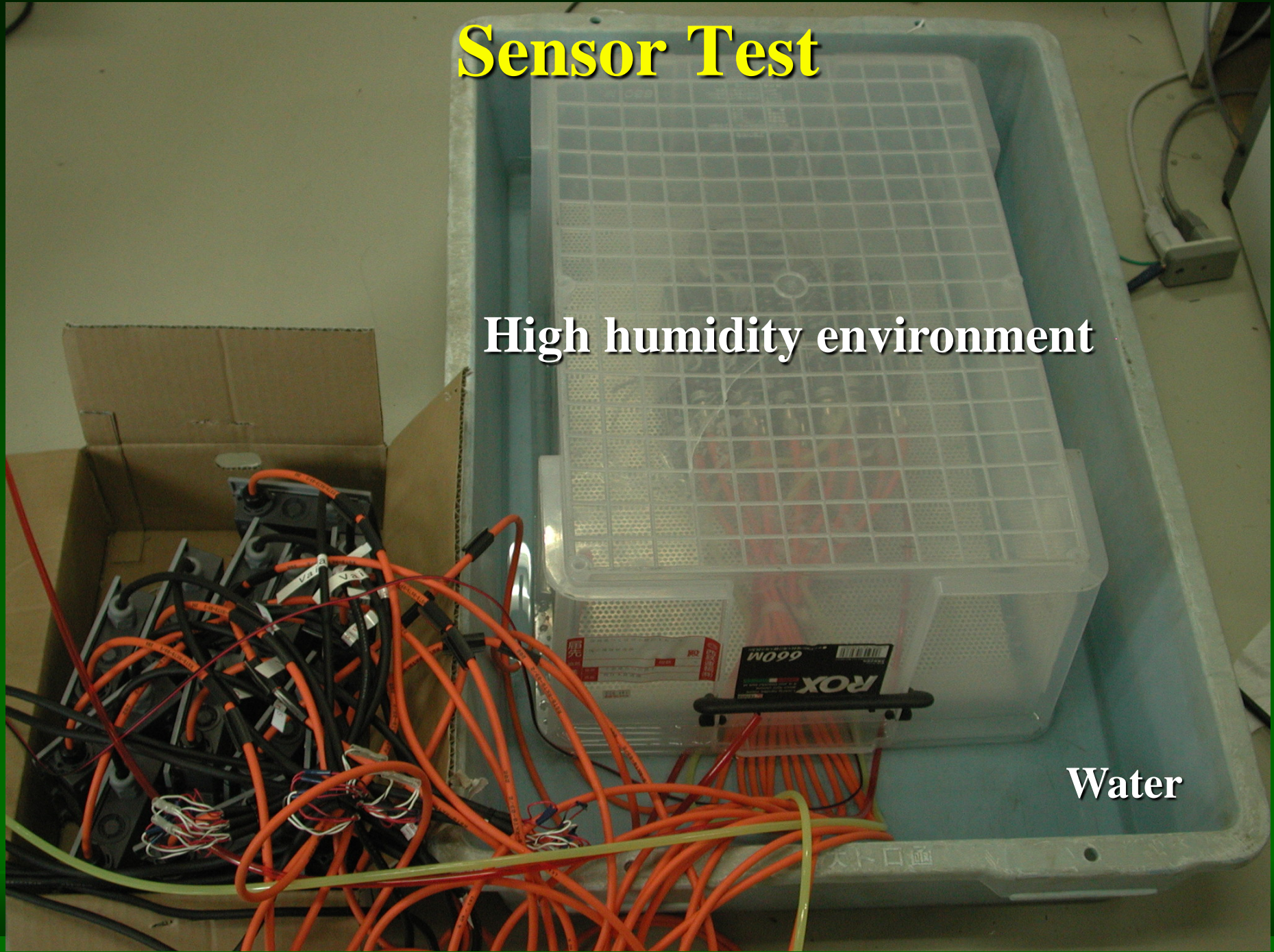
Transmitter

Probe

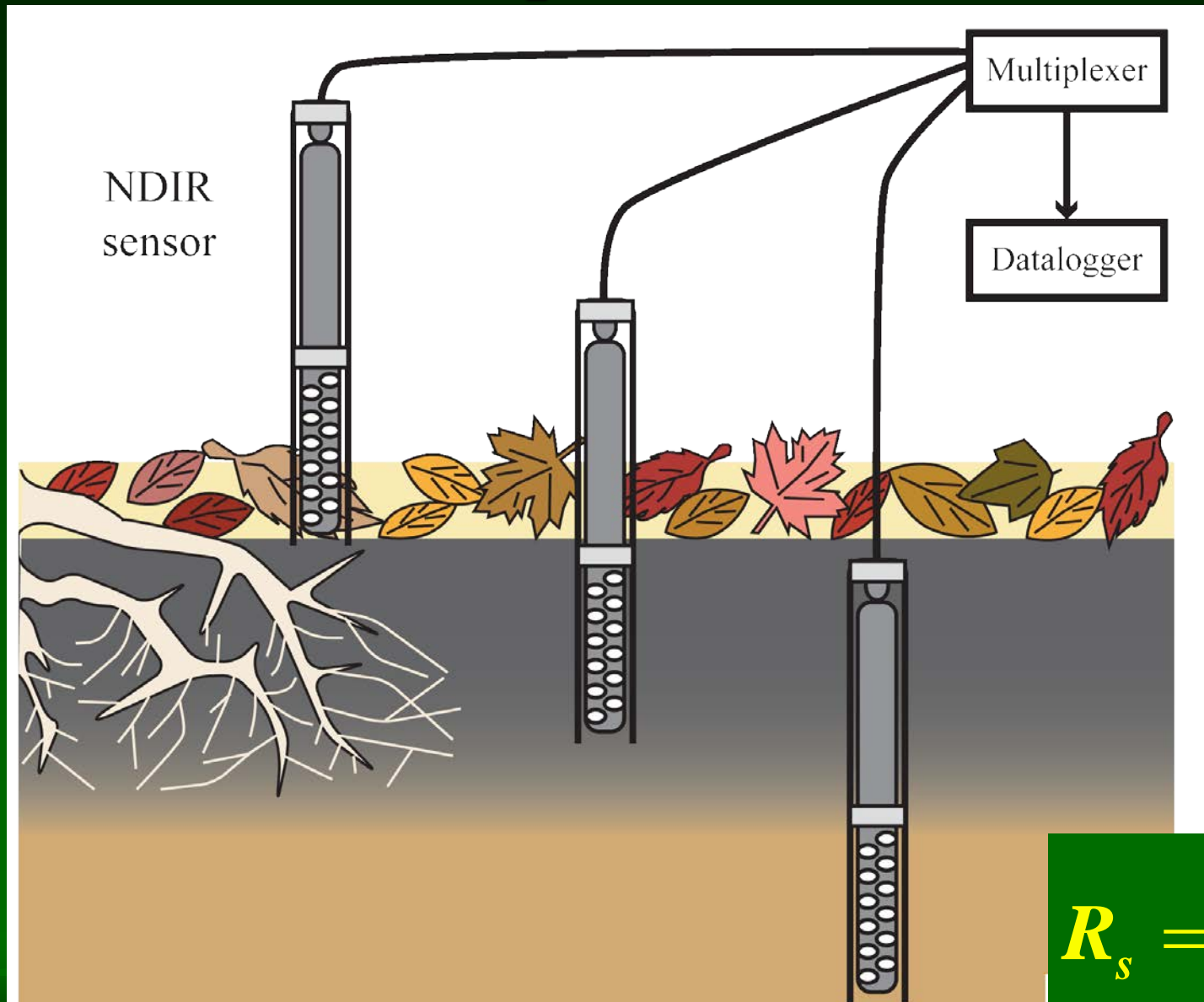
Sensor Test

High humidity environment

Water



Soil CO₂ Gradient Technique

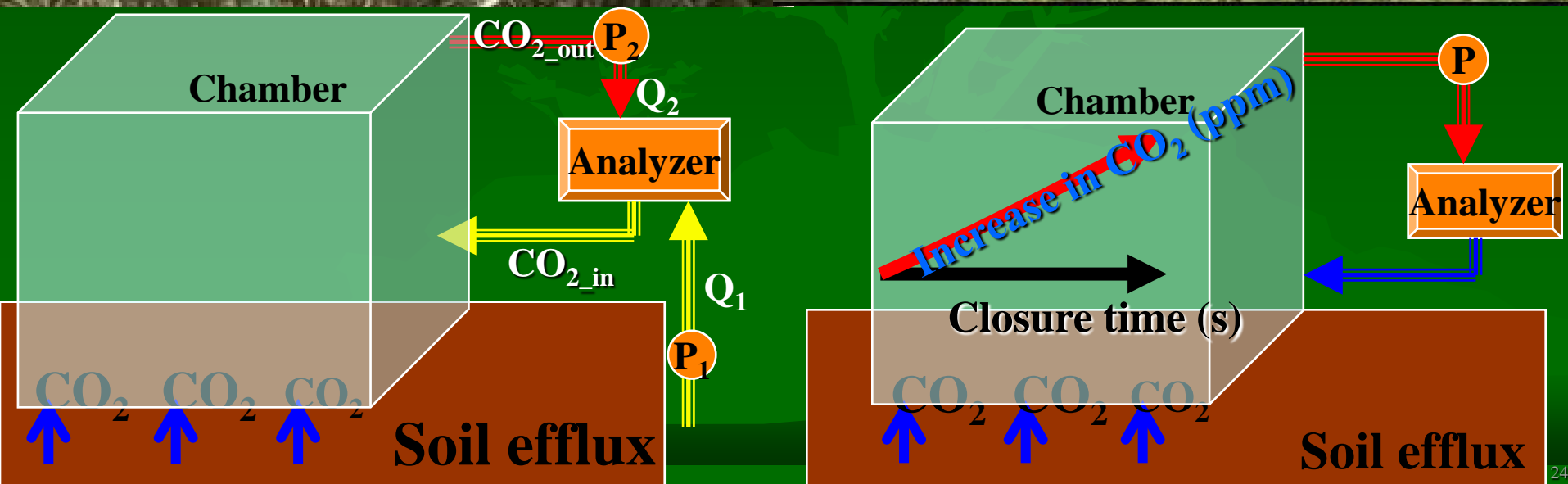


$$R_s = -D \frac{\partial C}{\partial X}$$

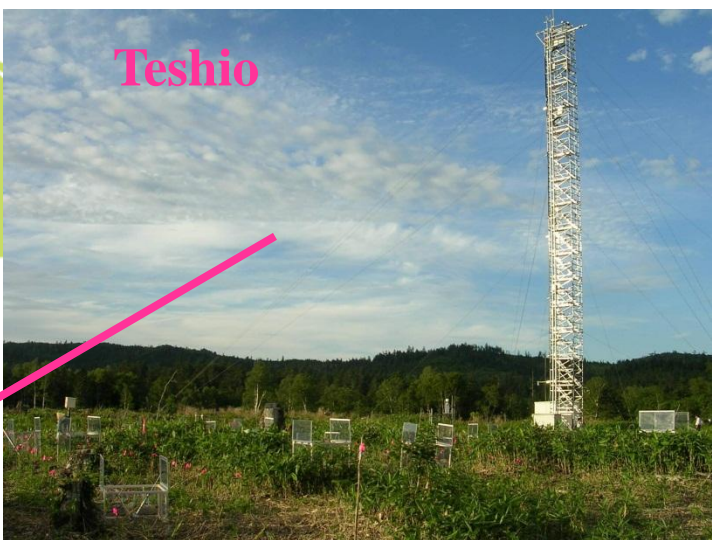
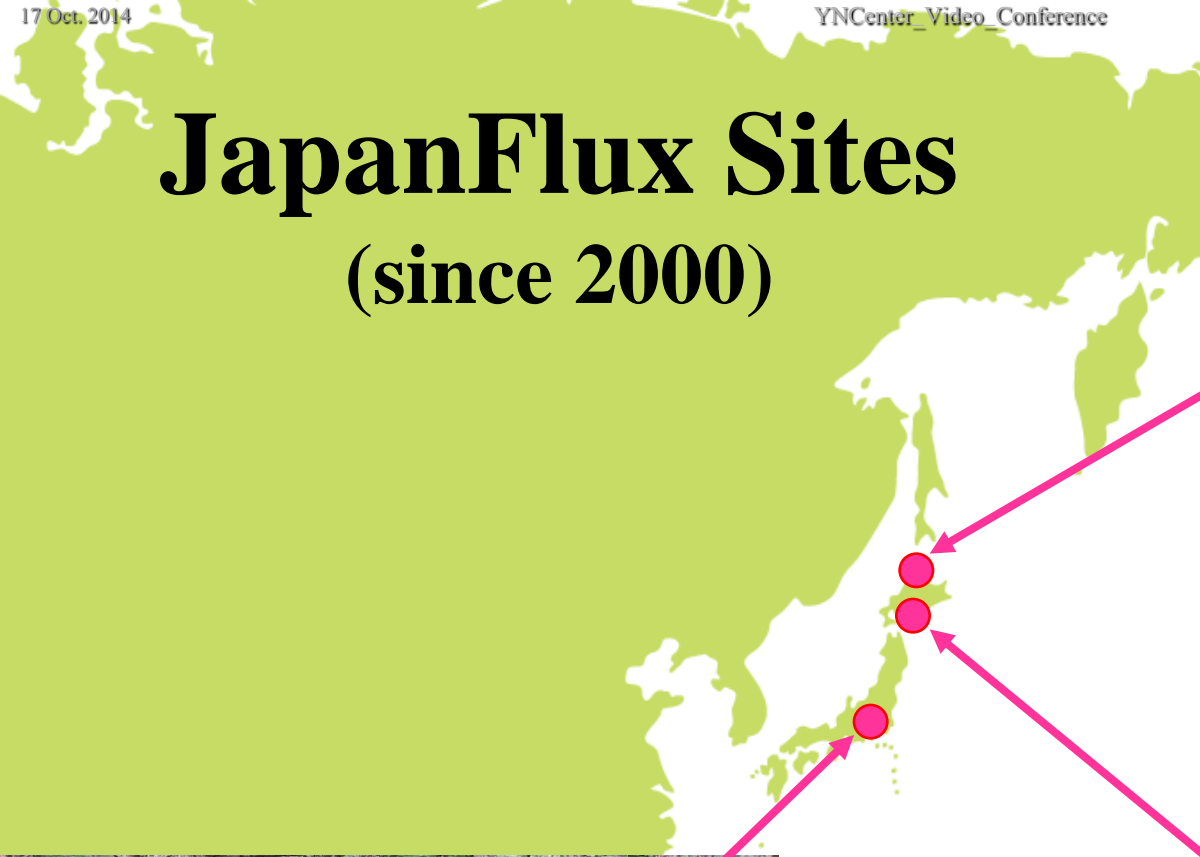
Automated Chambers for AsiaFlux

1997 model

2002 model



JapanFlux Sites (since 2000)



Teshio



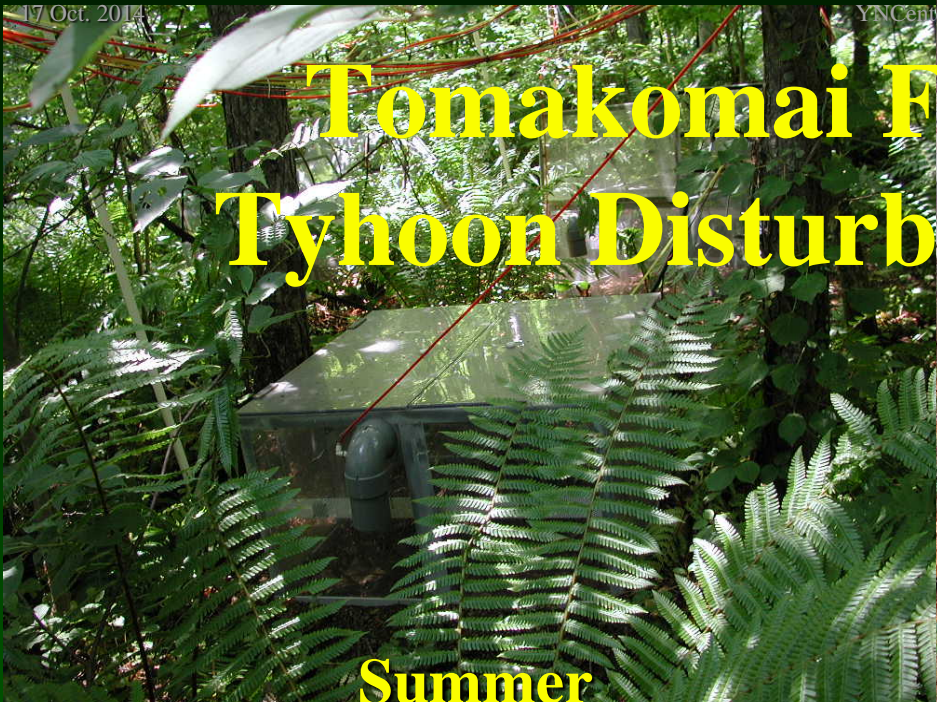
Tomakomai



Mt. Fuji



Tomakomai Flux Site Before Typhoon Disturbance (2000-2004)



Summer



Autumn

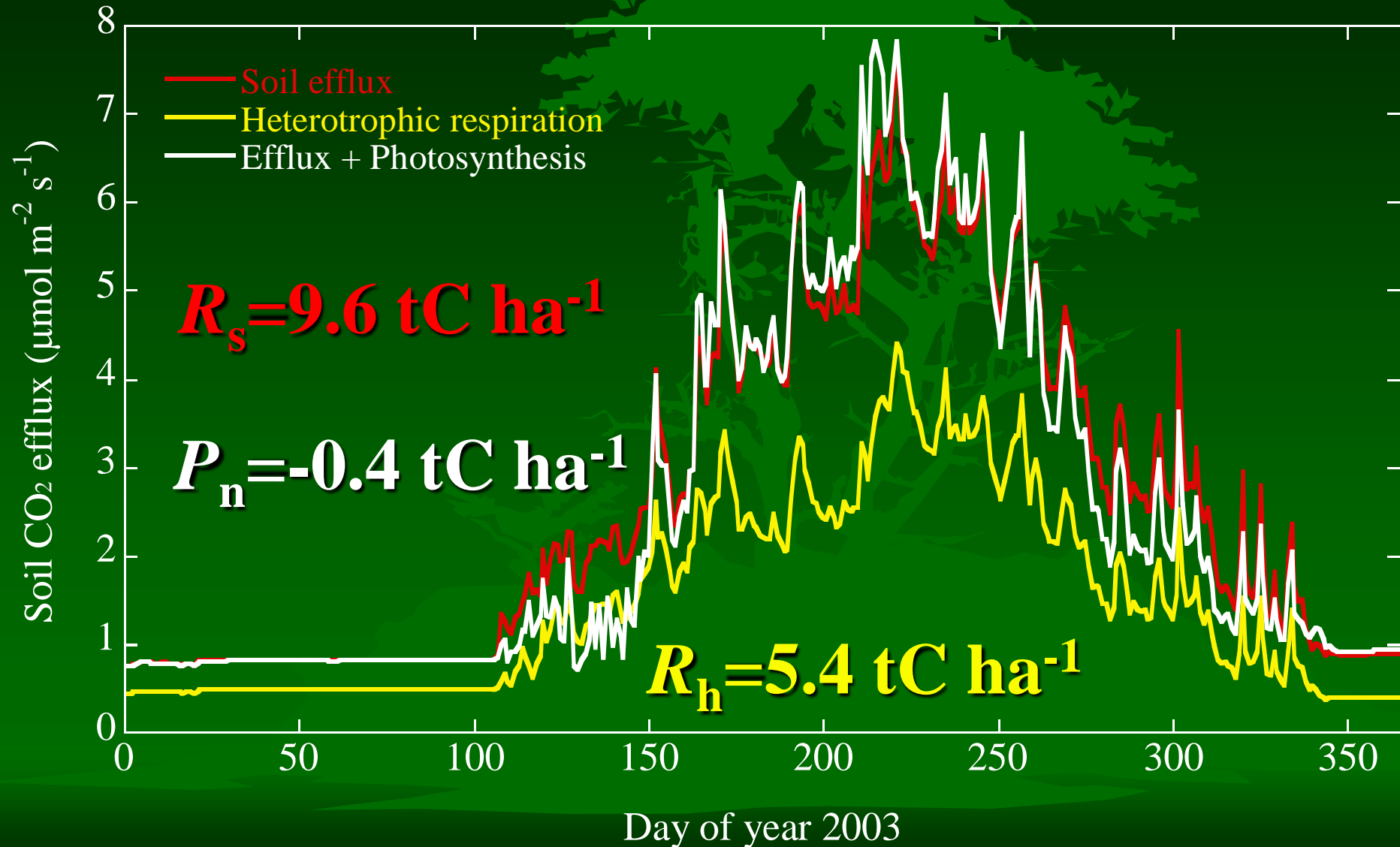


Winter



Spring

Soil Efflux, Heterotrophic Respiration, Understory Photosynthesis at Tomakomai



High temporal & Spatial Variations



Automated

Size = 90*90 cm



Open-top

D = 30 cm



LI6400-09

D = 9.5 cm

N. Liang et al. / Agricultural and Forest Meteorology 123 (2004) 97–117

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Table 1

Number of sampling points required for the automated chamber, open-top chamber, and LI-6400 chamber approaches to achieve different degrees of precision (within $\pm 10\%$ to within $\pm 20\%$ of the full sample mean) with 95% confidence interval

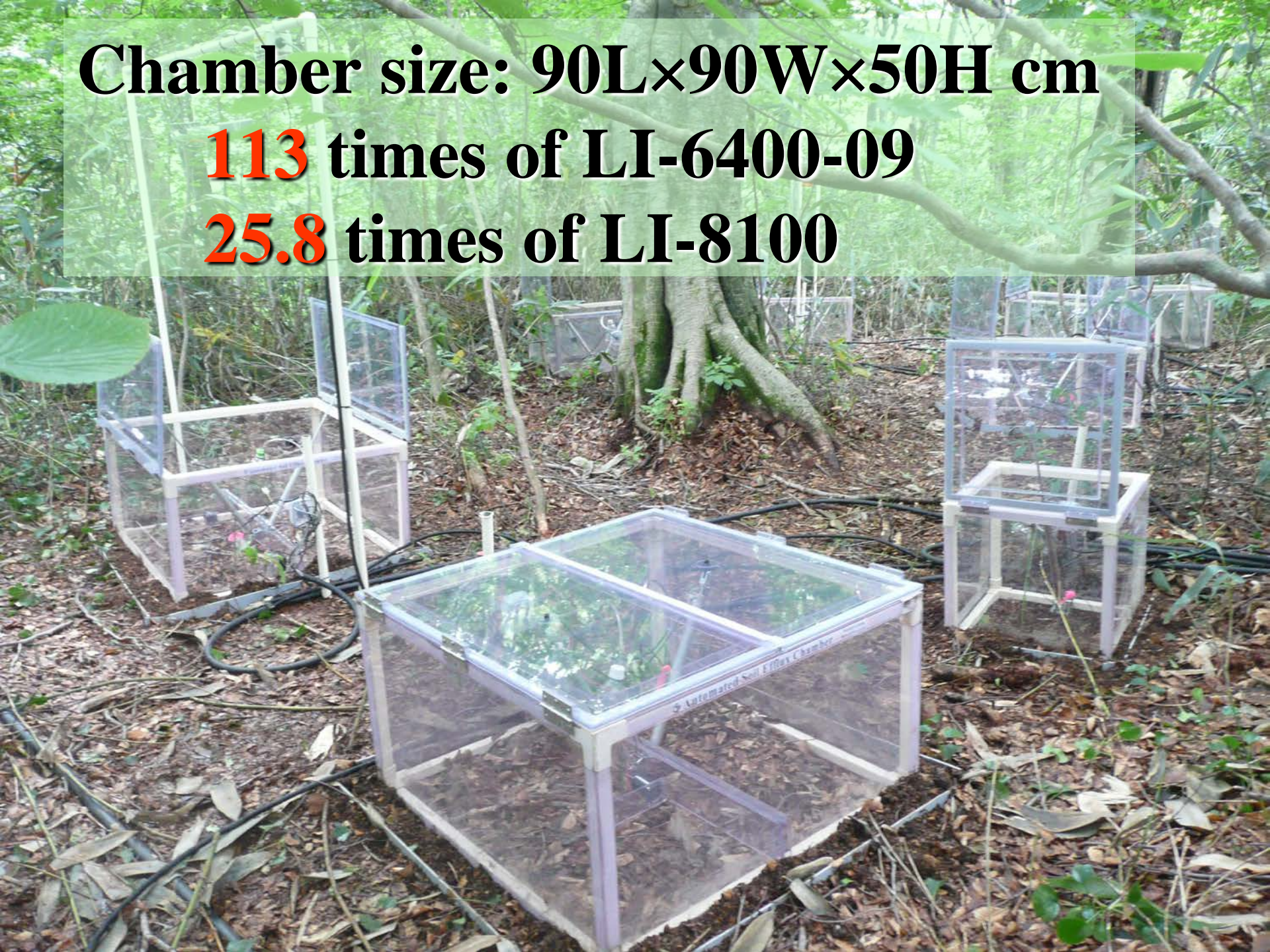
Chamber type	No. of sampling points actually measured	CO ₂ efflux (mean + S.D.) (mg C m ⁻² h ⁻¹)	No. of sampling points required for measurements	
			Within $\pm 10\%$	Within $\pm 20\%$
Automated	16	126 \pm 20	18	5
Open-top	9	118 \pm 35	65	16
LI-6400	20	202 \pm 90	355	89

Calculation is based on results averaged for the entire measurement period.

Chamber size: 90L×90W×50H cm

113 times of LI-6400-09

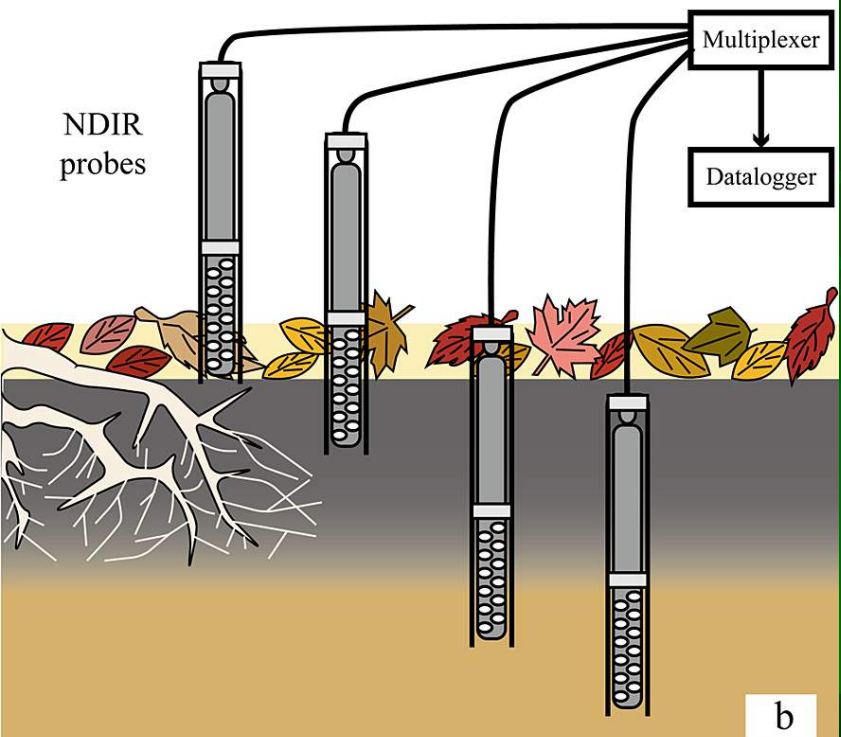
25.8 times of LI-8100





Control Unit





Soil CO₂ efflux of a larch forest in northern Japan

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Revised: 15 October 2010 – Accepted: 29 October 2010 – Published: 5 November 2010

Abstract. We had continuously measured soil CO₂ efflux (R_s) in a larch forest in northern Japan at hourly intervals for the snow-free period in 2003 with an automated chamber system and partitioned R_s into heterotrophic respiration (R_h) and autotrophic respiration (R_t) by using the trench method. In addition, we applied the soil CO₂ concentration gradients method to continuously measure soil CO₂ profiles under snowpack in the snowy period and to partition R_s into topsoil (O_a and A horizons) CO₂ efflux (F_t) with a depth of 0.13 m and sub-soil (C horizon) CO₂ efflux (F_c). We found that soil CO₂ effluxes were strongly affected by the seasonal variation of soil temperature but weakly correlated with soil moisture, probably because the volumetric soil moisture (30–40% at 95% confidence interval) was within a plateau region for root and microbial activities. The soil CO₂ effluxes changed seasonally in parallel with soil temperature in topsoil with the peak in late summer. On the other hand, the contribution of R_t to R_s was the largest at about 50% in early summer, when canopy photosynthesis and plant growth were more active. The temperature sensitivity (Q_{10}) of R_t peaked in June. Under snowpack, R_s was stable until mid-March and then gradually increased with snow melting. R_s summed up to 79 gC m⁻² during the snowy season for 4 months. The annual R_s was determined at 934 gC m⁻² y⁻¹ in 2003, which accounted for 63% of ecosystem respiration. The annual contributions of R_h and R_t to R_s were 57% and 43%, respectively. Based on the gradient approach, R_s was partitioned vertically into litter (O_l and O_e horizons) with a depth of 0.01–0.02 m, topsoil and sub-soil respirations with proportions of 6, 72 and 22%, respectively, on an annual basis. The vertical distribution of CO₂ efflux was consistent with those of soil carbon and root biomass.

1 Introduction

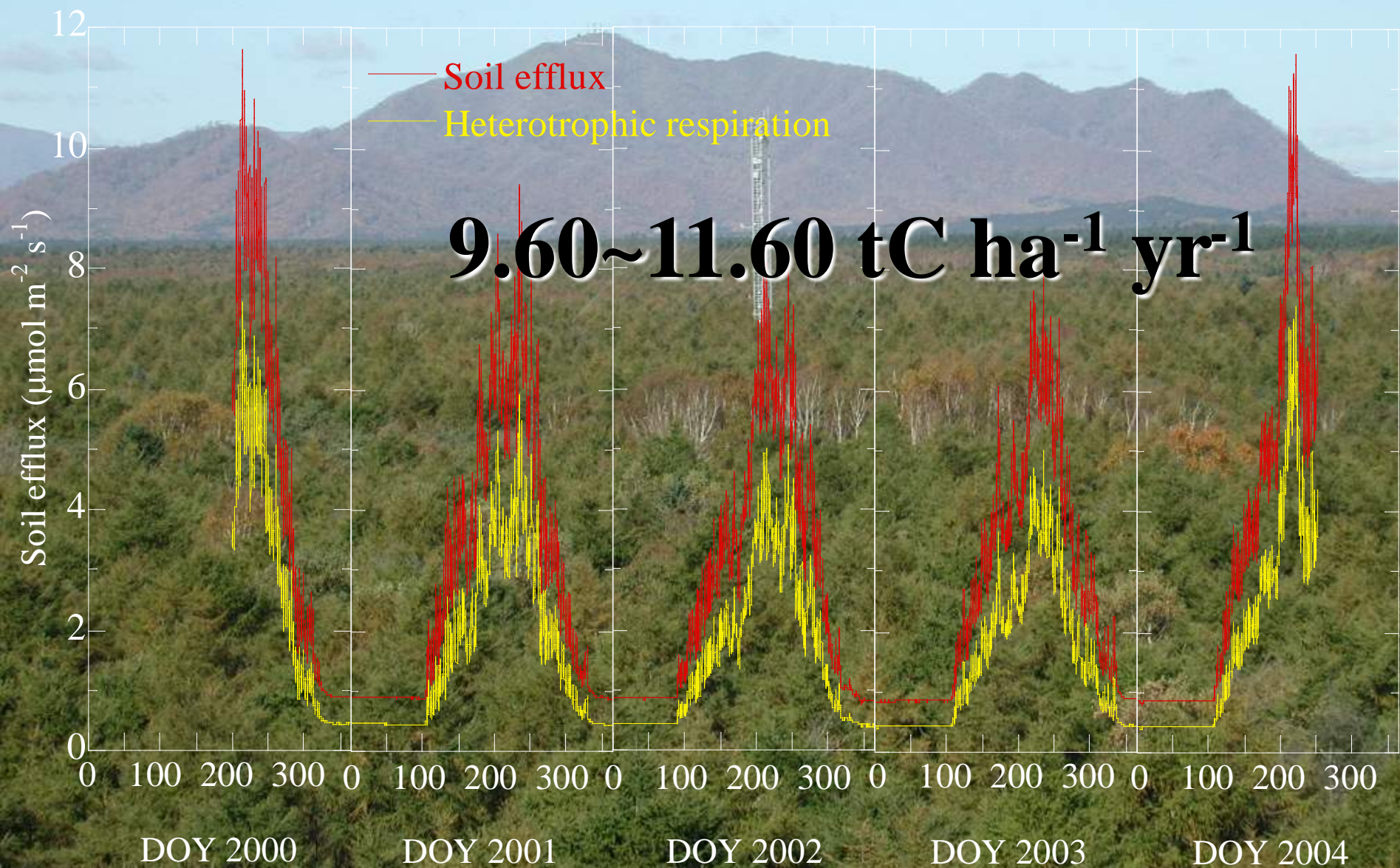
The world's soils contain about 1550 Pg of organic carbon, which is more than twice the amount in the atmosphere (IPCC, 2007). Forests worldwide contain about 45% of the global carbon stock, a large part of which is in the forest soils. Recently, Bond-Lamberty and Thomson (2010) estimated that the global soil CO₂ efflux, widely referred to as soil respiration (R_s), was about 98 Pg C y⁻¹ in 2008 based on a five-decade record of chamber measurements, which is more than 13 times the rate of fossil fuel combustion (IPCC, 2007), indicating that 20–40% of the atmospheric CO₂ circulates through soils every year. Overall, R_s is the largest component of ecosystem respiration (RE) and the second largest flux in the global carbon cycle after gross primary production (GPP). R_s is therefore a key process that is fundamental to our understanding of the terrestrial carbon cycle (Davidson and Janssens, 2006). A relatively small change in the carbon flow into or out of soils can strongly influence the global carbon cycle. For example, it was reported that the global R_s increased by 0.1 Pg C y⁻¹ between 1989 and 2008 (Bond-Lamberty and Thomson, 2010), and that the positive feedback from this enhancement of R_s by global warming would further raise atmospheric CO₂ concentration by 20–224 ppm by 2100 and resulting higher CO₂ levels would lead to an additional temperature increase ranging from 0.1 to 1.5 °C (Friedlingstein et al., 2006; IPCC, 2007).

In forest ecosystems, micrometeorological studies (i.e., eddy covariance) have shown that, on average, about 80% of GPP is respired back to the atmosphere (Law et al., 2002), and R_s has been estimated to account for 60–90% of RE, with marked temporal as well as spatial variations (Law et al., 1999; Janssens et al., 2001a, b; Liang et al., 2004). Therefore, R_s has recently received much attention from researchers and its accurate measurement is critical for developing a reliable model of carbon exchange in forest ecosystems (Jassal et al., 2007; Zhou et al., 2009).

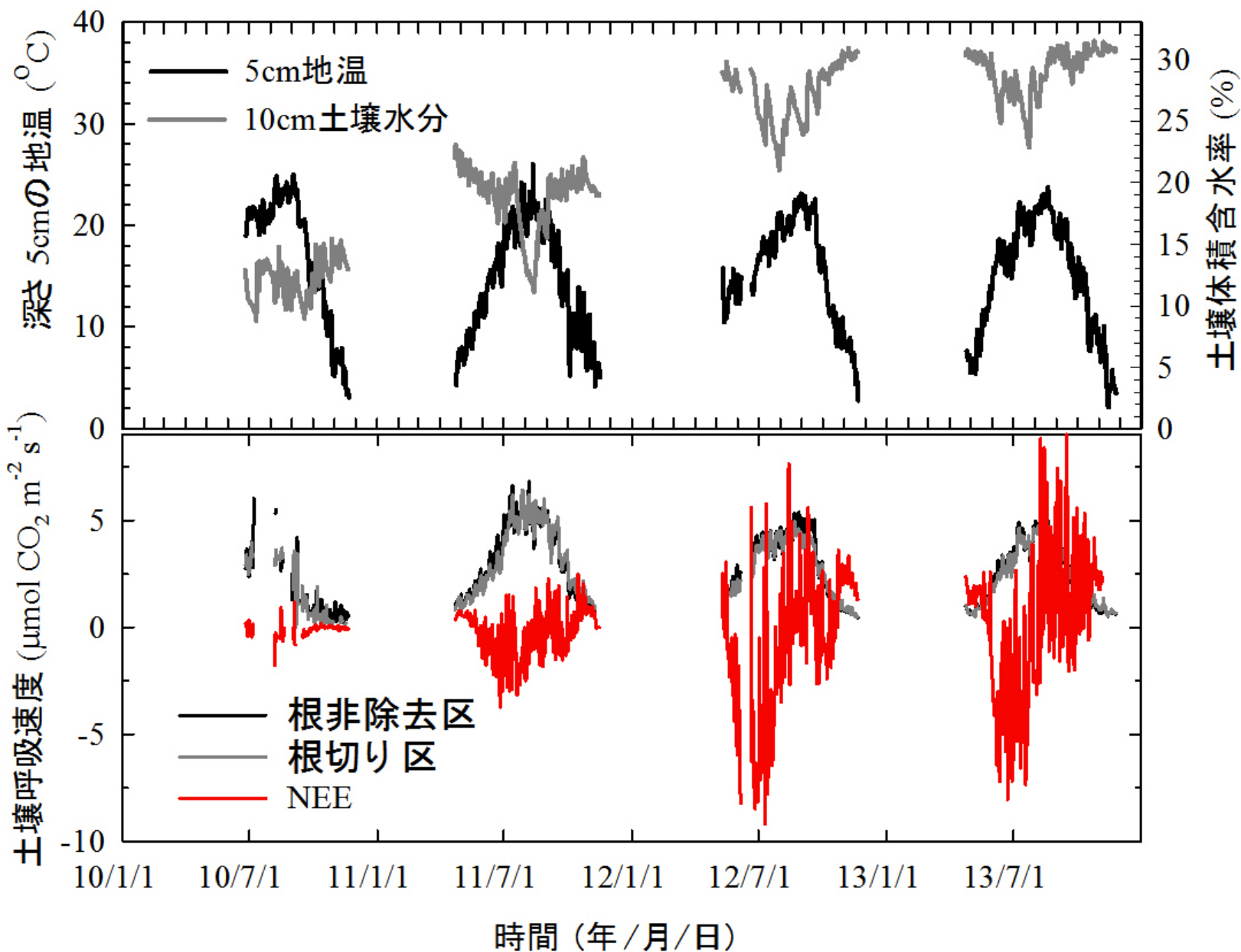


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Inter-annual Variation of Soil Efflux at Tomakomai



苦小牧の台風跡地





Larch



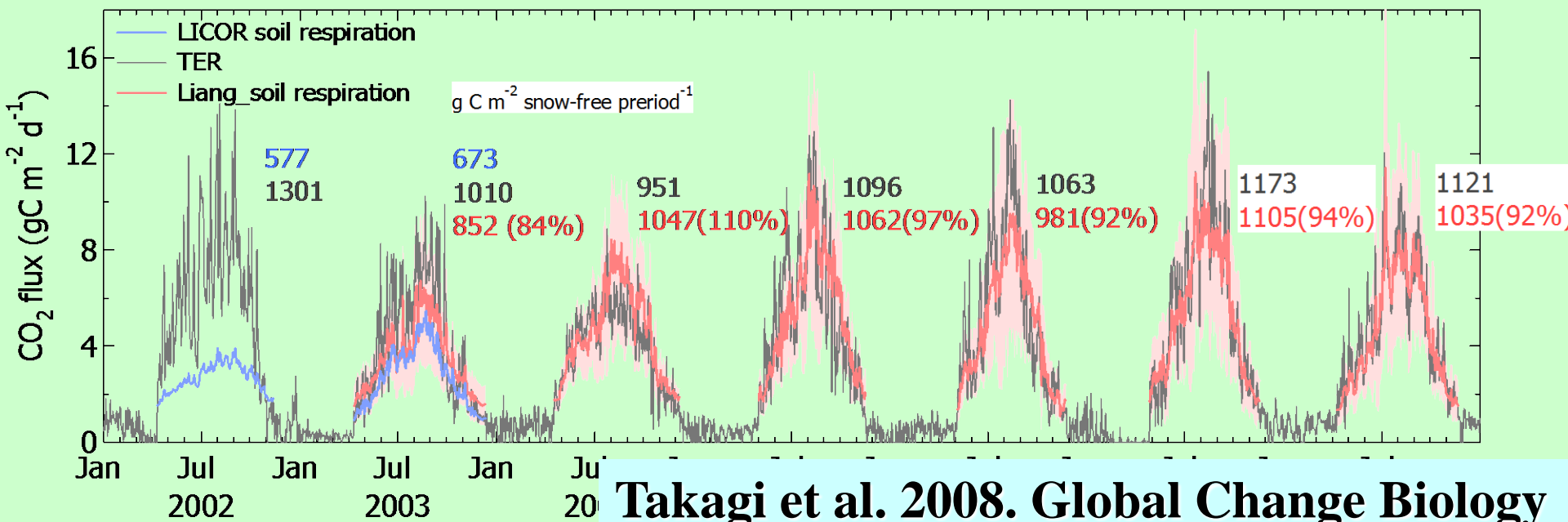
Teshio CC-Lag Site



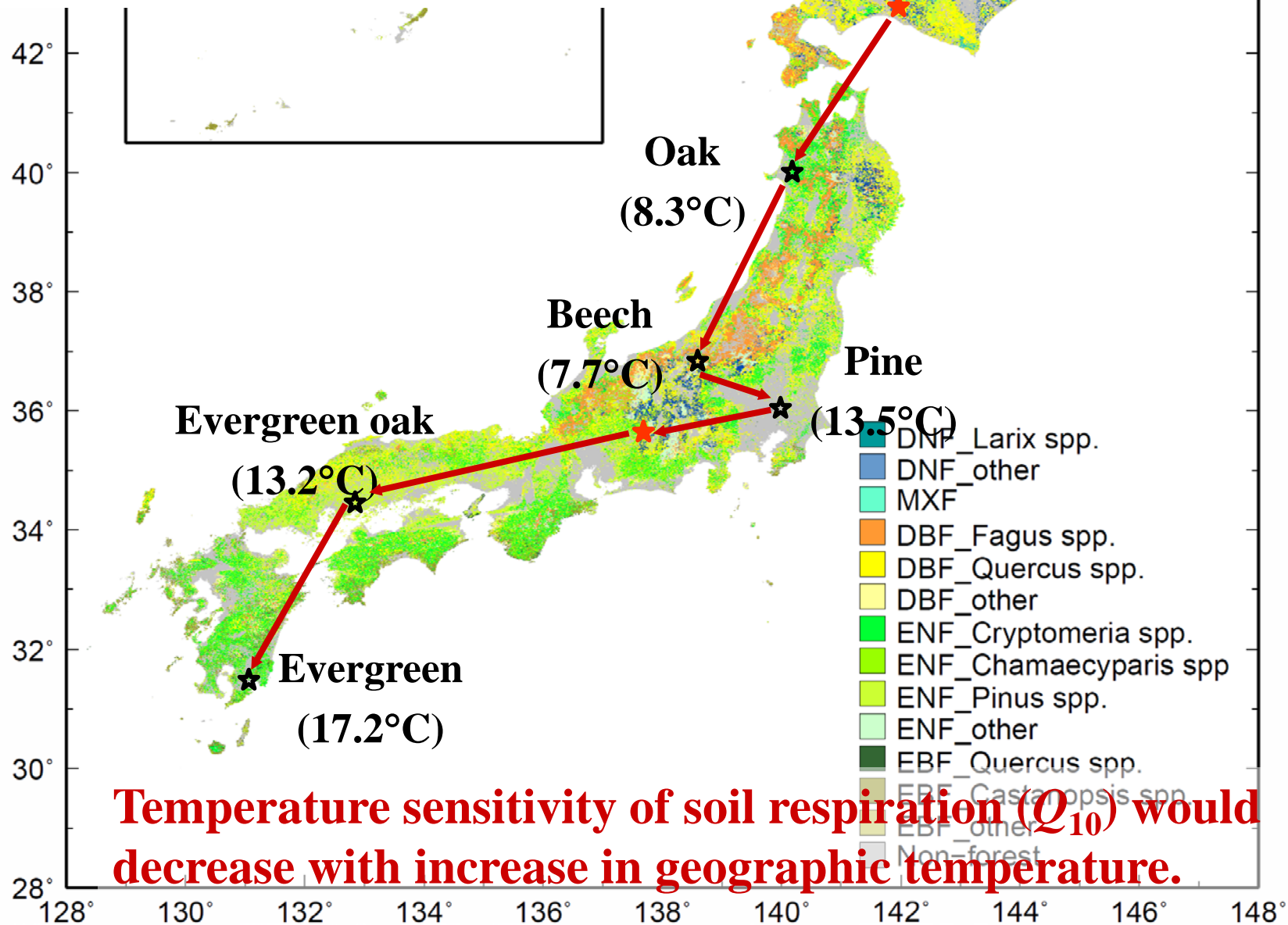
Sasa

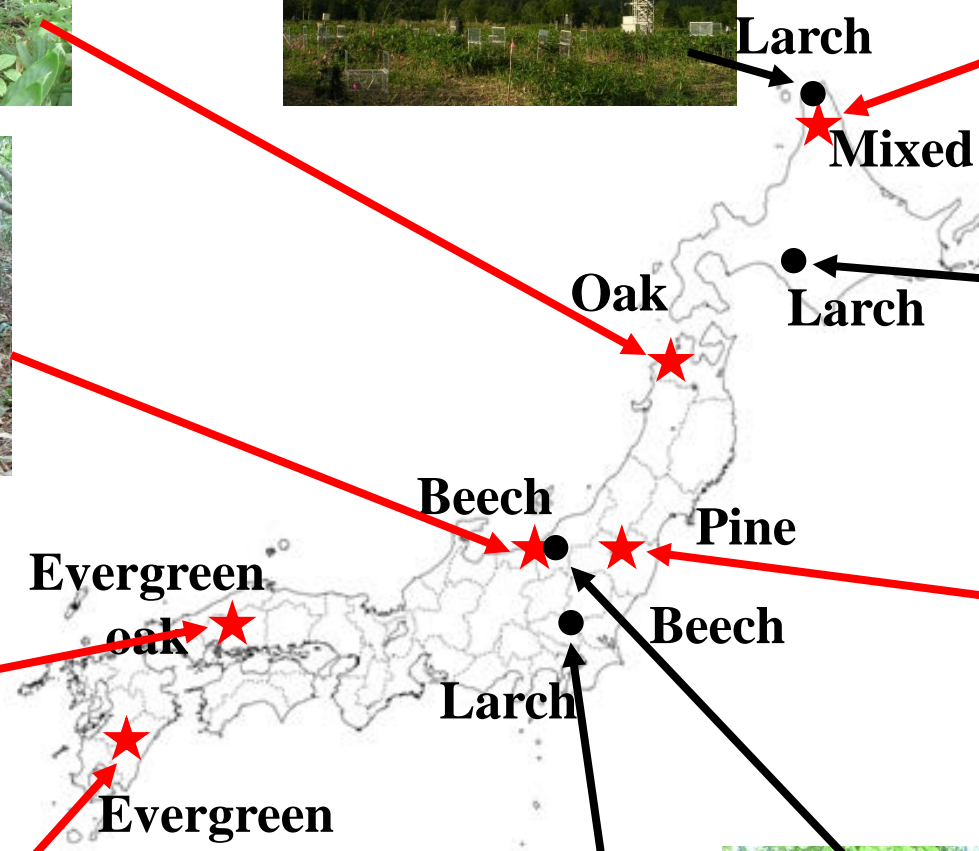


Soil



Hypothesis: Q_{10} vs. Temperature





Experiment Designing



15~24 chambers (90×90 × 50 cm)

5 heterotrophic plots



10 trench plots

5 warming plots

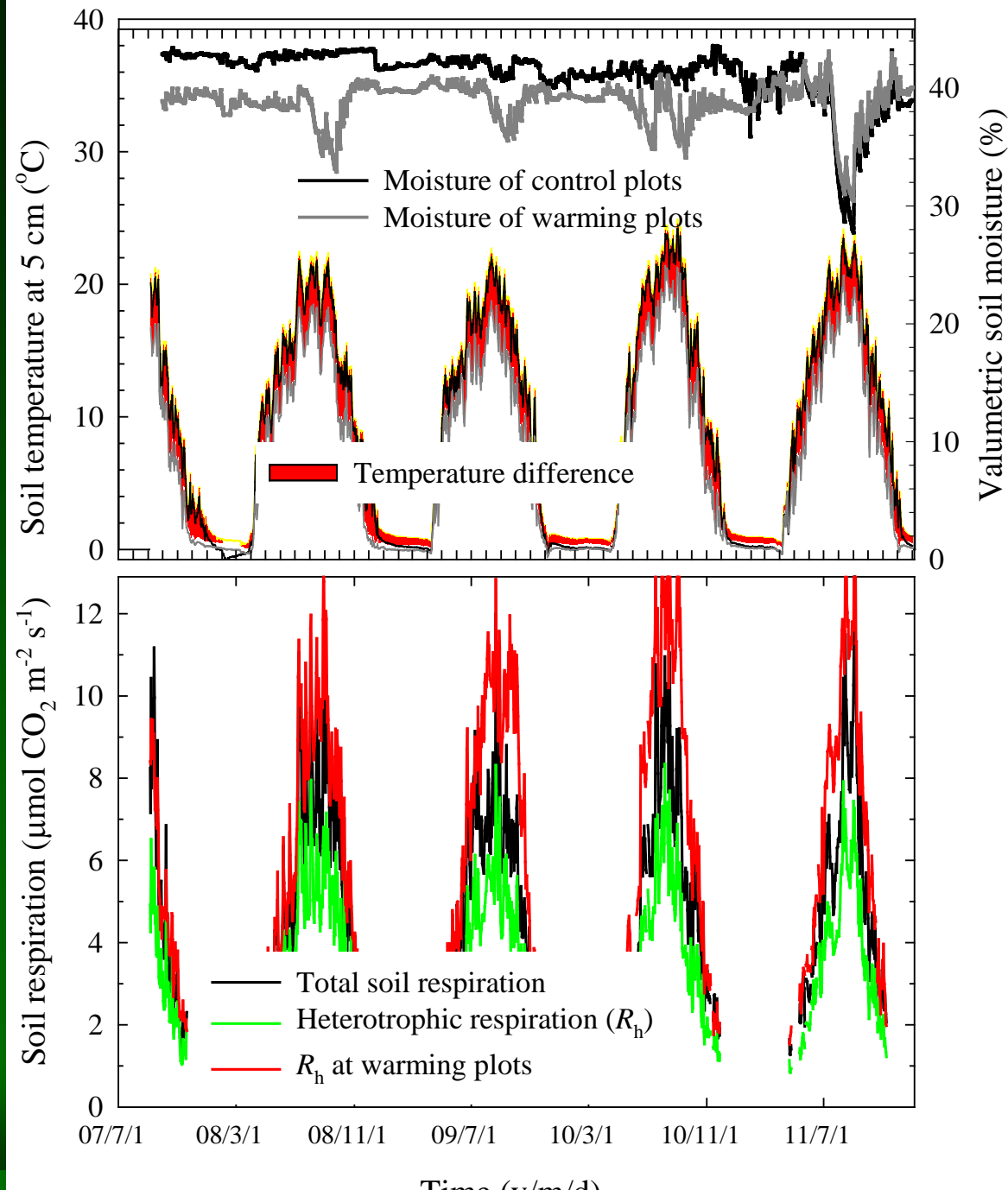
Cool-temperate mixed forest in northern Hokkaido

Warming:

↑ 3.0°C T_{soil} at 5-cm

Effect:

Rh ↑ 20% per °C



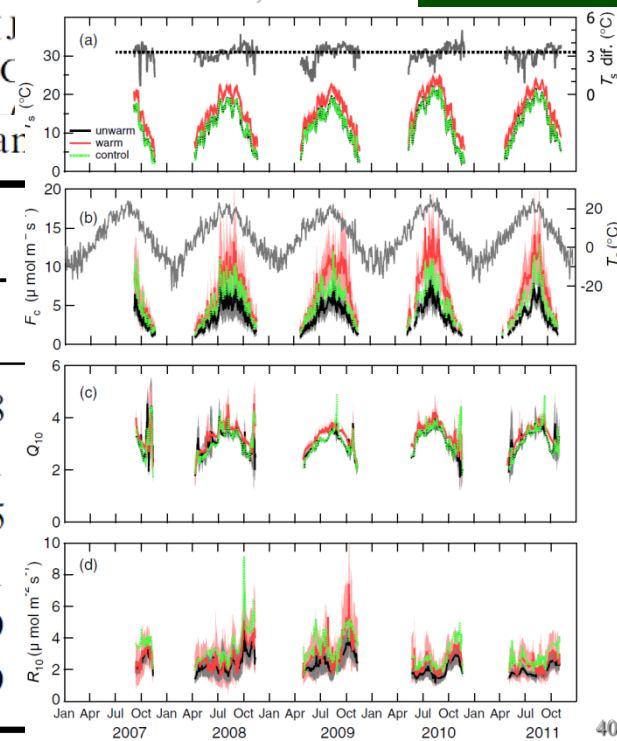
Sustained large stimulation of soil heterotrophic respiration rate and its temperature sensitivity by soil warming in a cool-temperate forested peatland

By MARICAR AGUILOS¹, KENTARO TAKAGI^{2*}, NAISHEN LIANG³,
YOKO WATANABE², MUNEMASA TERAMOTO³, SEIJ
YOSHIYUKI TAKAHASHI³, HITOSHI MUKAI³ and KAICHI

¹Graduate School of Environmental Science, Hokkaido University, Sapporo 060-0814

Table 4. Seasonal average (each measurement period for each year) and star

Year	Q_{10}		
	Unwarmed	Warmed	Control
2007	3.15 ± 0.35	3.18 ± 0.32	2.99 ± 0.18
2008	3.06 ± 0.24	3.27 ± 0.10	3.04 ± 0.11
2009	3.00 ± 0.04	3.15 ± 0.09	3.00 ± 0.05
2010	3.17 ± 0.06	3.47 ± 0.11	3.25 ± 0.11
2011	3.16 ± 0.19	3.33 ± 0.03	3.16 ± 0.09
Four seasons	3.10 ± 0.08	3.31 ± 0.06	3.11 ± 0.09



Japanese Forest Map

Mixed: 20%/°C

($Q_{10}=2.89$)

Oak : 8.5%/°C

($Q_{10}=3.4$)

Beech : 6.8%/°C

($Q_{10}=3.11$)

Pine : 4.0%/°C

($Q_{10}=3.29$)

Evergreen oak : 5.2%/°C

($Q_{10}=2.97$)

Evergreen : 11%/°C

($Q_{10}=3.12$)

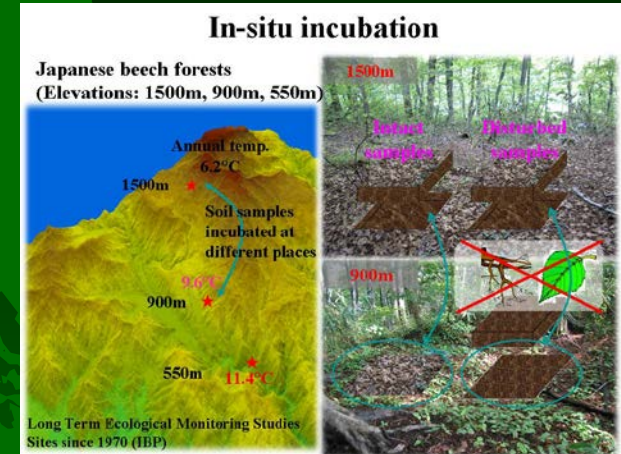
- DNF_Larix spp.
- DNF_other
- MXF
- DBF_Fagus spp.
- DBF_Quercus spp.
- DBF_other
- ENF_Cryptomeria spp.
- ENF_Chamaecyparis spp
- ENF_Pinus spp.
- ENF_other
- EBF_Quercus spp.
- EBF_Castanopsis spp.
- EBF_other
- Non-forest

46°
44°
42°
40°
38°
36°
34°
32°
30°
28°

128° 130° 132° 134° 136° 138° 140° 142° 144° 146° 148°

On-going Project

Warming Experiments



50-year-old plantation



Tomakomai: Disturbed by typhoon



Teshio: clear-cut & reforested



Evaluation of effects by short-term climate change, natural disturbance and human disturbance

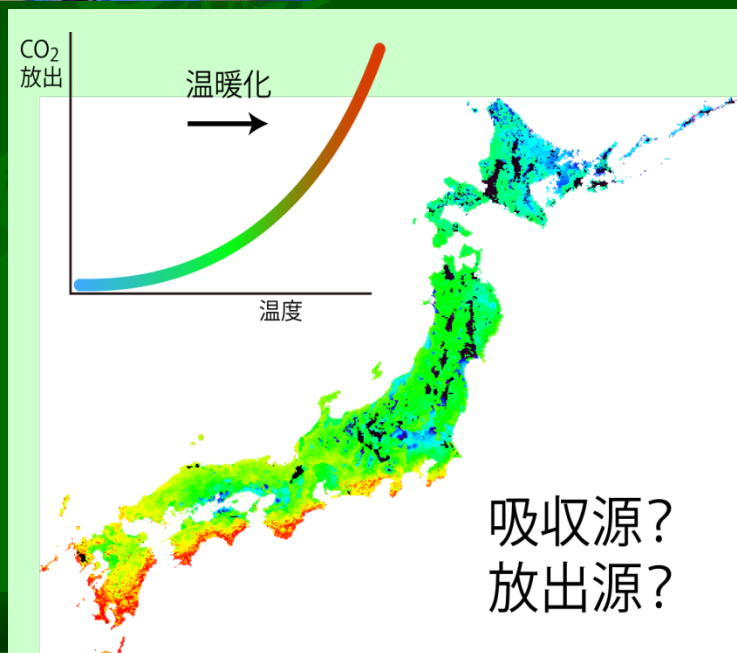
On-going Technology

NIES-AMS



Understanding mechanisms

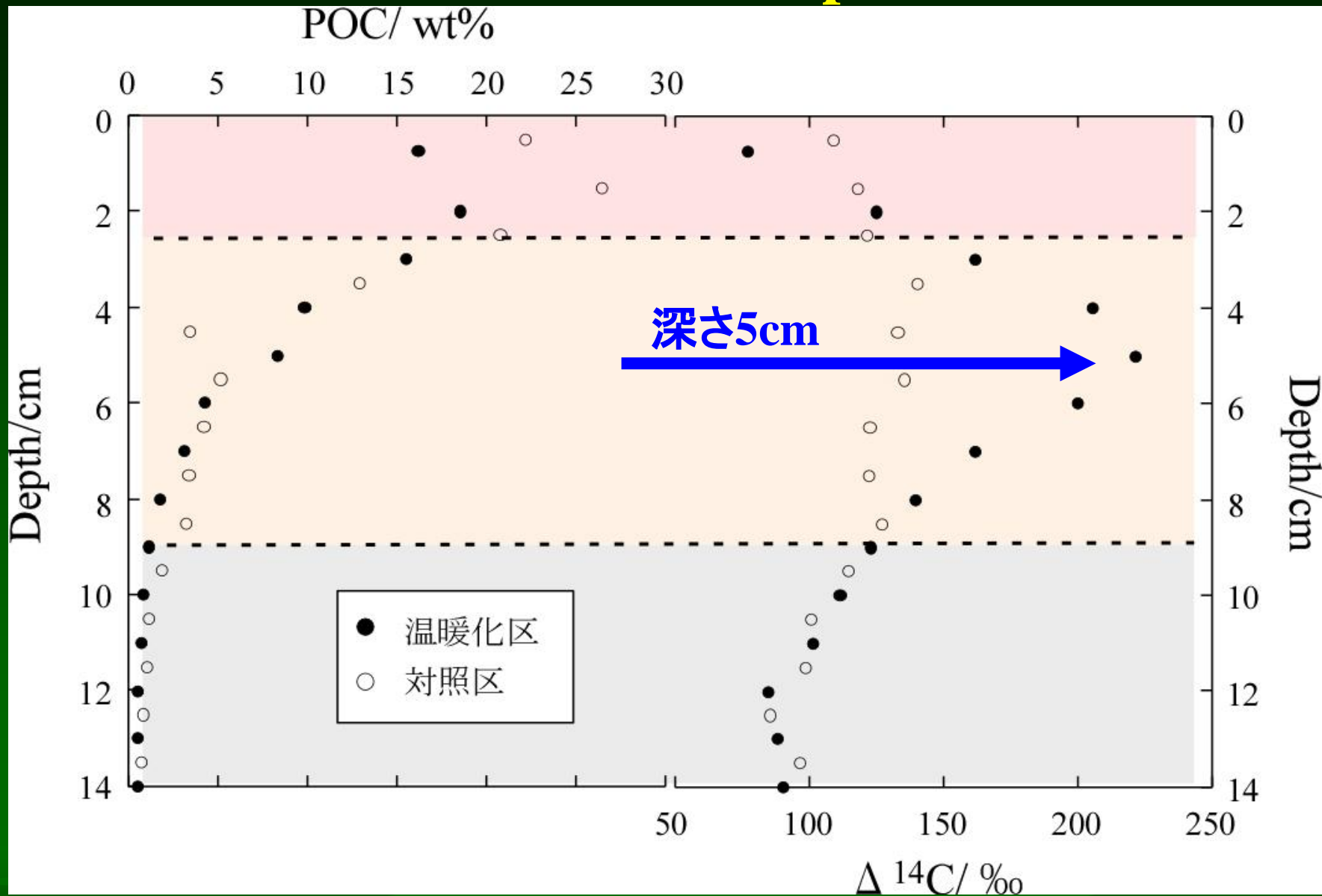
**Answering questions:
Carbon sink?
source?**



Sampling for Soil ^{14}C

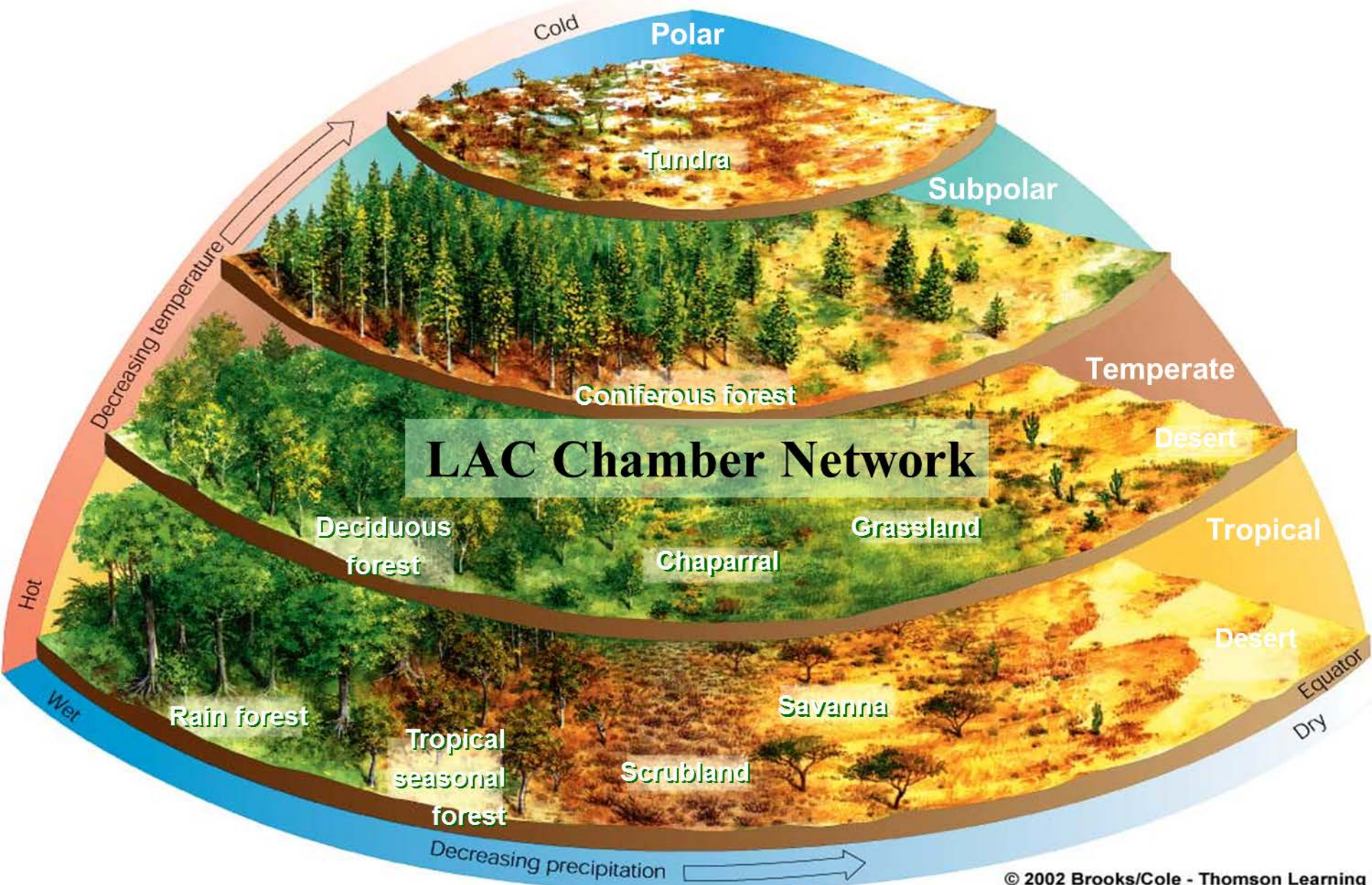


Profiles of POC & $\Delta^{14}\text{C}$ of Japanese Forest Soil



Older SOC seems more sensitive to temperature than that of new SOC

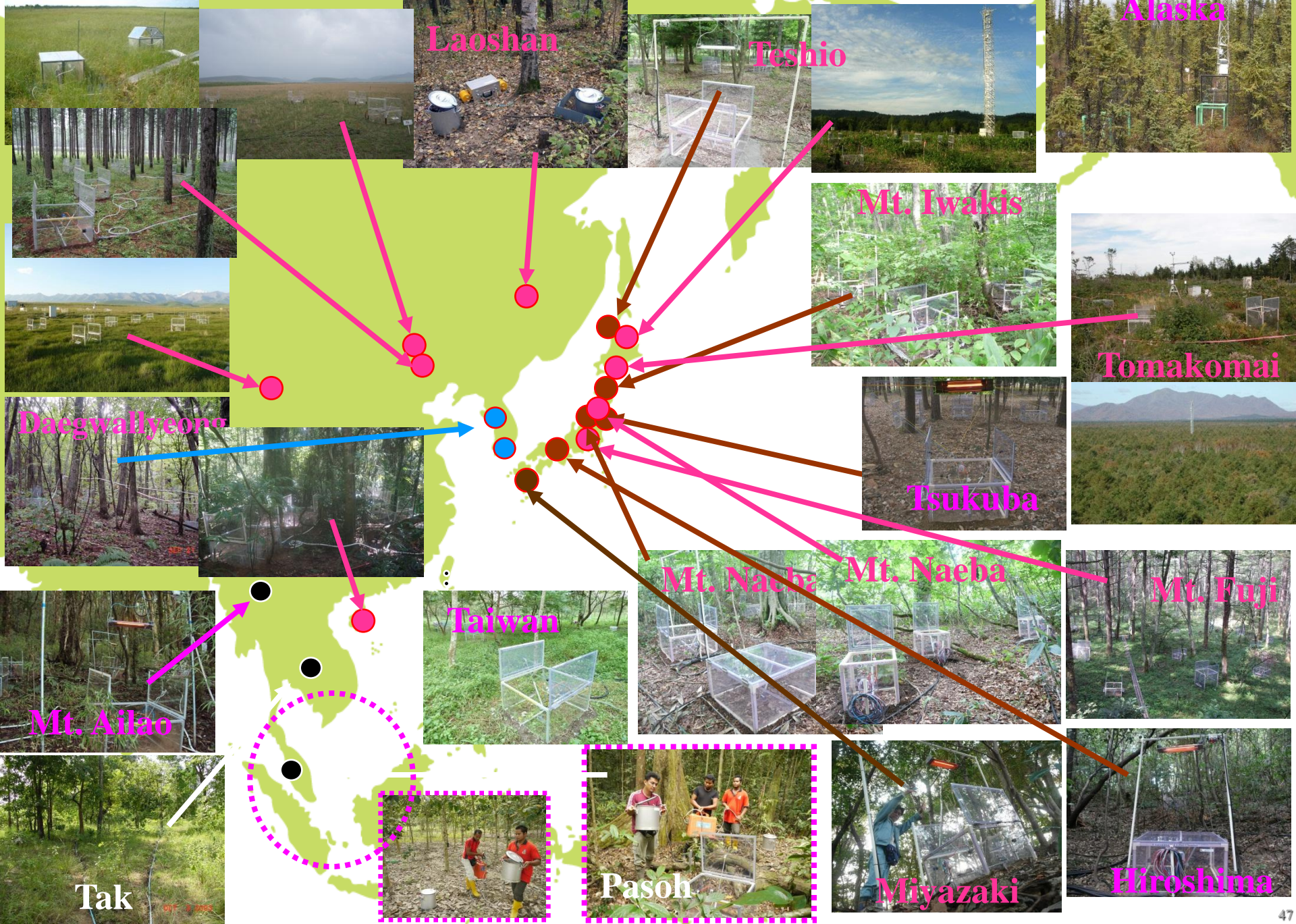
Ecosystems (Biomes) in Asia





Automated Chamber Networks

YNCenter_Video_Conference



Alaska

Laoshan

Teshio

Mt. Iwakis

Tomakomai

Daegwallyeong

Tsukuba

Mt. Naeba

Mt. Naeba

Mt. Fuji

Taiwan

Mt. Ailao

Pasoh

Miyazaki

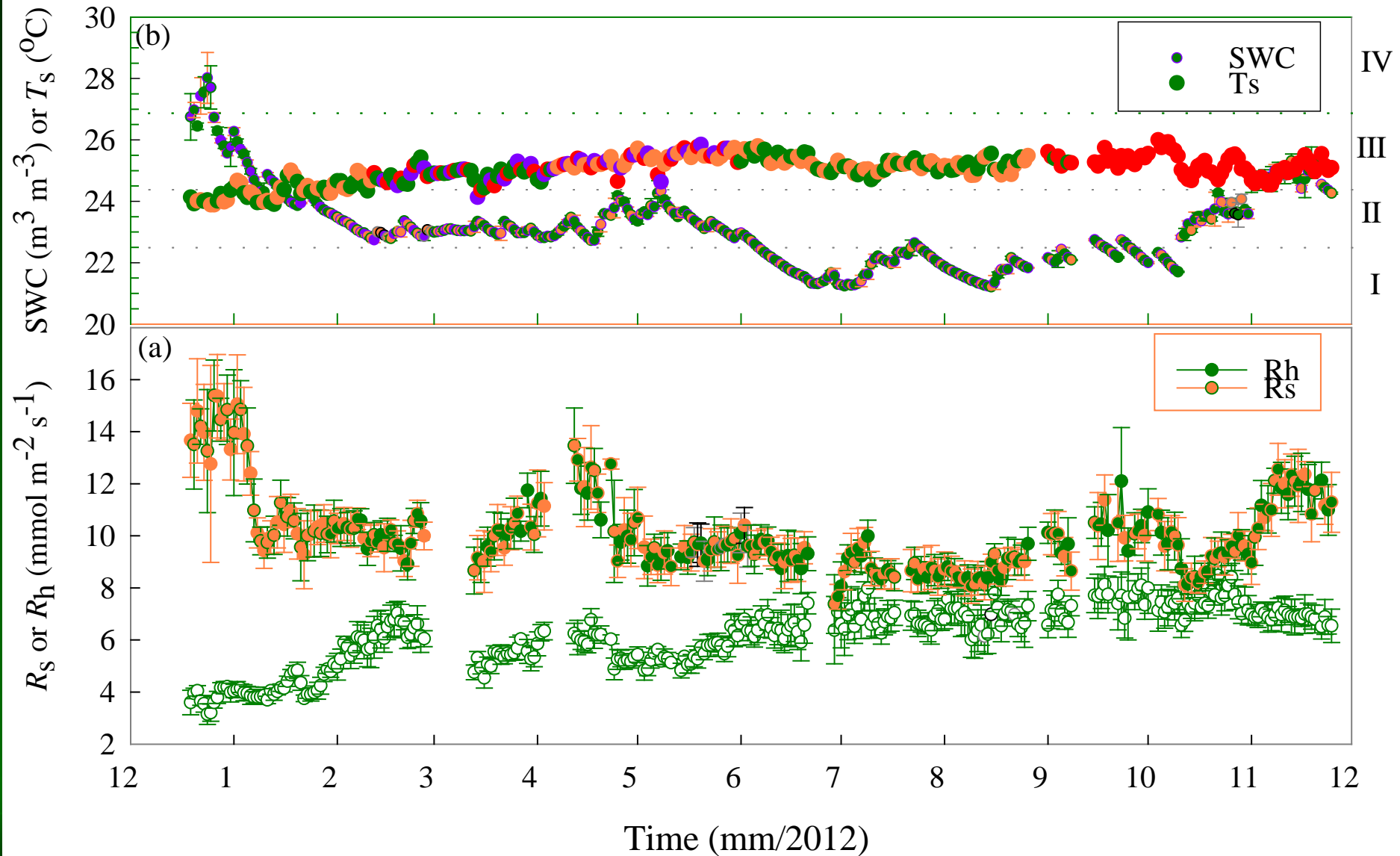
Hiroshima

Tak

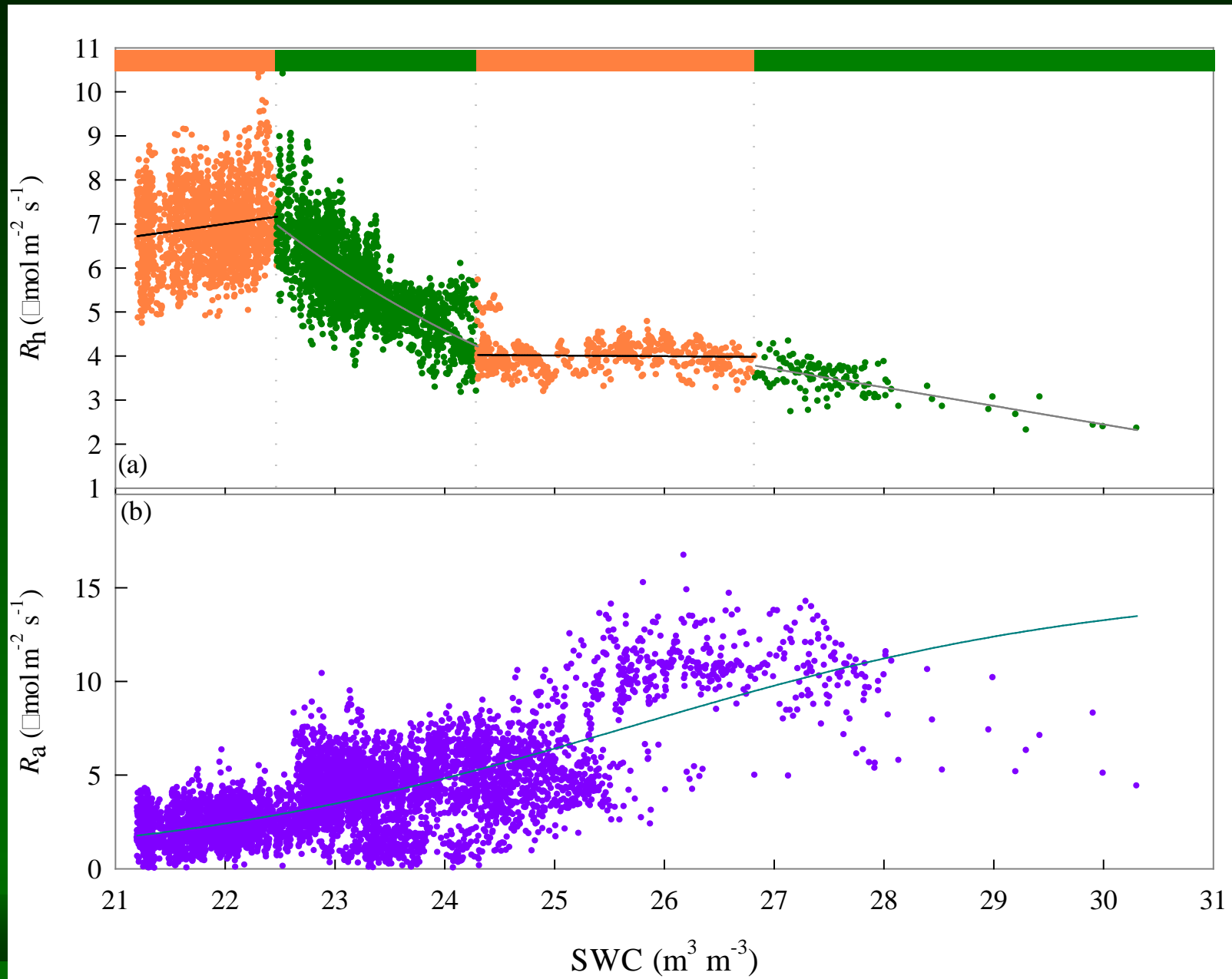
Continuous Measurement Rs at Malaysian Tropical Rainforests



Seasonal Changes in Soil Efflux at Pasoh Tropical Forest



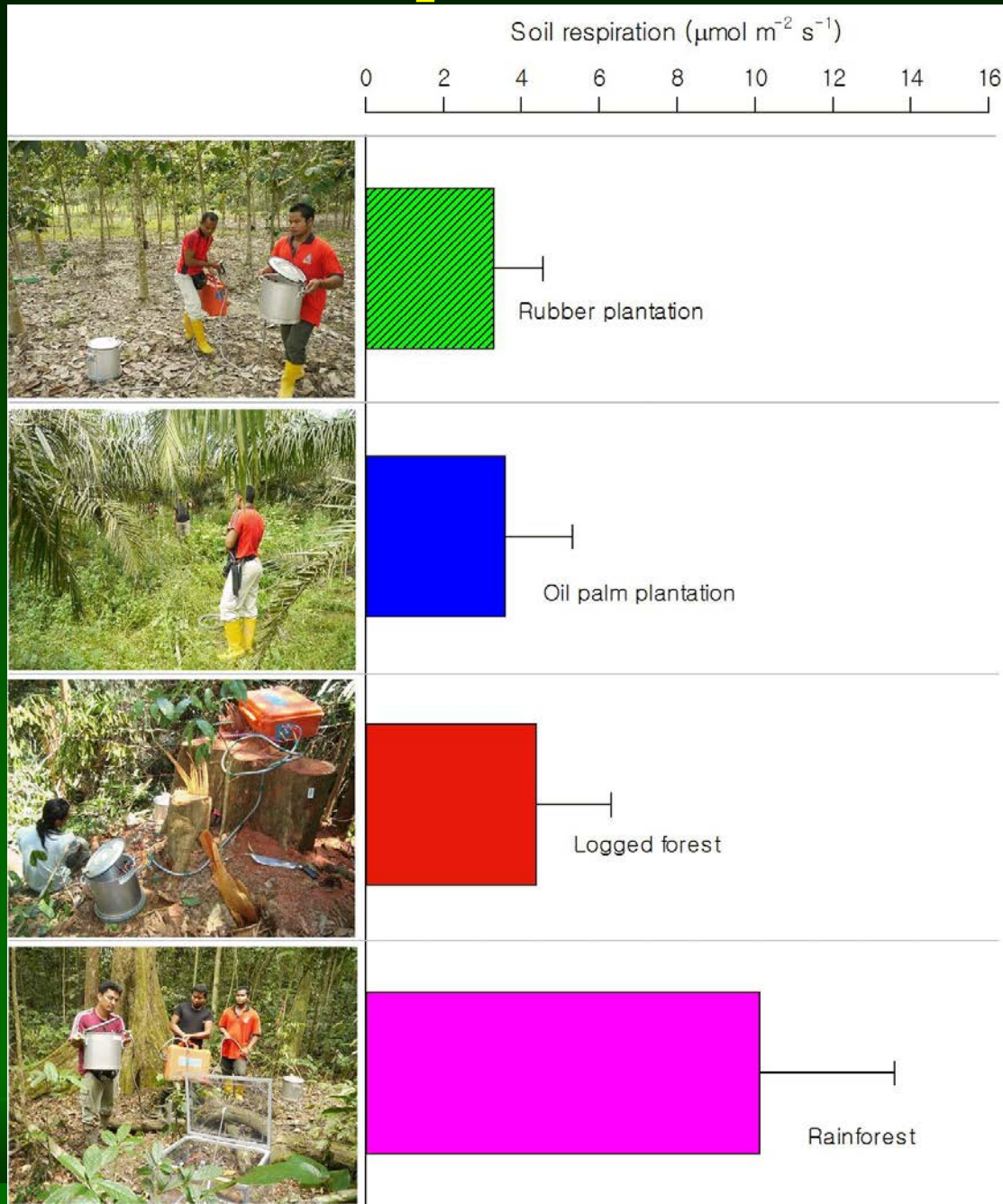
Soil Moisture Response of Auto- & Heterotrophic Respiration



Soil Respiration vs. Land Use Change



Soil Respiration at Different Land Use Types



Land use changes
caused soil
degradation (土壤
退質(劣化))



Logging site

Decomposition of harvest residue

Monitor & control by mobile phone



Soil CO₂ Efflux vs. REDD Credit

With 30% of logging intensive (low-impact logging):

$$T_2 = 23.5^\circ\text{C} + 1.5^\circ\text{C} \text{ (soil temperature increase)}$$

$$C = 38 \text{ tC/ha (harvest residue: AG=21, BG=17)}$$

$$d = 13\% \text{ (decomposition ratio of CWD)}$$

$$\$ = 15 \text{ US\$/tC}$$

Total carbon lost at the first year after the logging:

$$\sum \$ = (3.32 + 5.70) \times 15 = 135 \text{ US\$ / ha}$$

$$\text{REDD credit} = 247 - 135 = 112 \text{ US\$ / ha}$$

Low Power Consumption



Chamber Network in China

疏勒河高原冻土
(秦大河)



疏勒河上游高寒草甸生态系统观测场

海北 (贺金生)



乱海子湿地

アジア

多伦 (李胜功)



塞罕坝 (朴世龙)



天塩



国立環境研究所

富士山

広島

宮崎

麗江

哀牢山



麗江 (张一平?)

溪頭

屏東



哀牢山 (张一平)



尖峰岭 (方精云)

尖峰嶺



22°

Subtropical Broadleaf Forest at Mt. Ailao



JOURNAL OF GEOPHYSICAL RESEARCH: ATMOSPHERES, VOL. 118, 2981–2990, doi:10.1002/jgrd.50300, 2013

Soil respiration in an old-growth subtropical forest: Patterns, components, and controls

Zheng-Hong Tan,^{1,2} Yi-Ping Zhang,^{1,3,4} Naishen Liang,² Qing-Hai Song,^{1,3,5}
Yu-Hong Liu,^{1,4} Guang-Yong You,^{1,3,5} Lin-Hui Li,^{1,3,5} Lei Yu,^{1,3,5} Chuan-Shen Wu,^{1,4}
Zhi-Yun Lu,^{1,4} Han-Dong Wen,^{1,4} Jun-Fu Zhao,^{1,3,5} Fu Gao,⁶ Lian-Yan Yang,^{1,3}
Liang Song,¹ Yong-Jiang Zhang,^{1,7} Teramoto Munemasa,² and Li-Qing Sha¹



Continuous Automated Chamber (CAC) with Picarro G1301



Seasonal Variation of Soil Temperature and CH₄ Flux

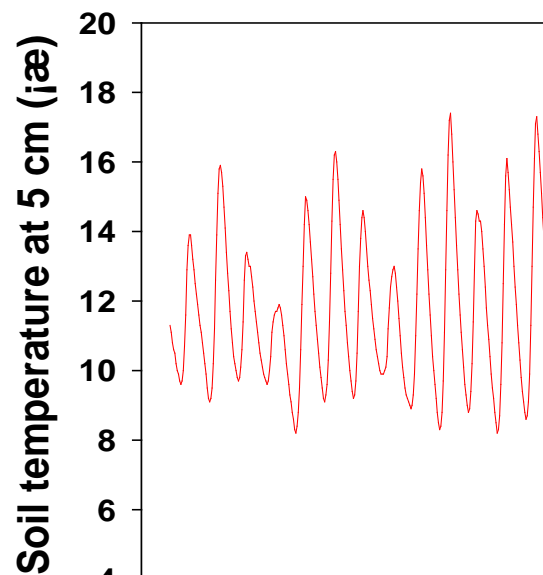
Environmental Pollution 181 (2013) 81–90

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Environmental Pollution

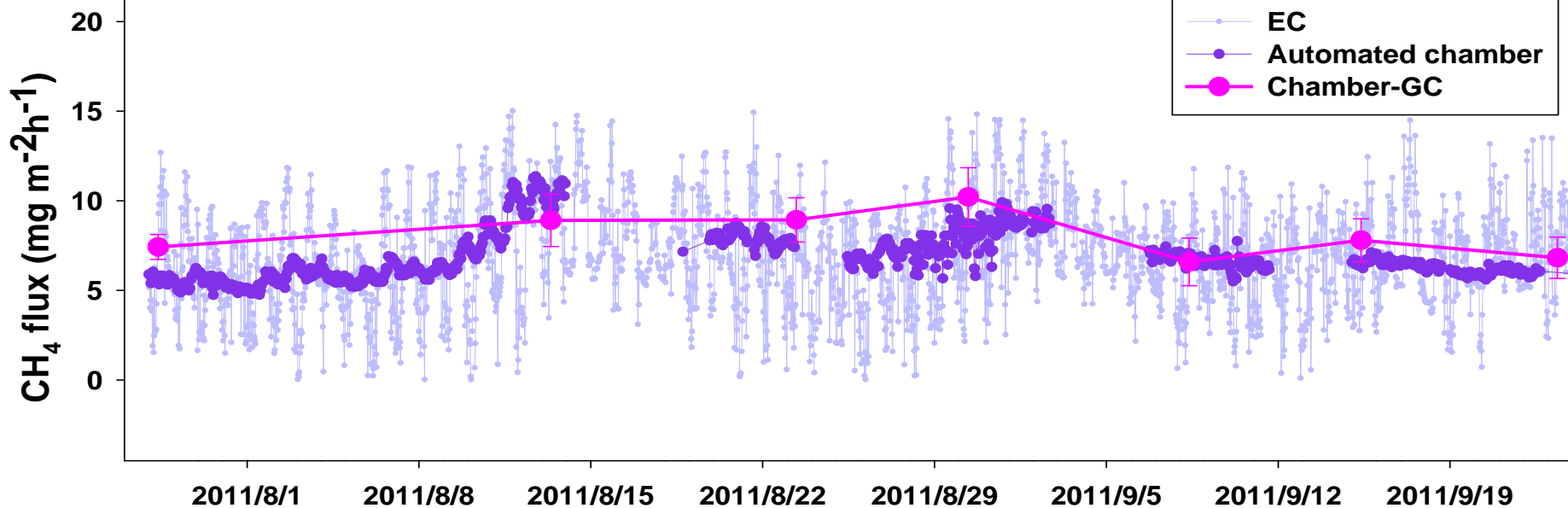
journal homepage: www.elsevier.com/locate/envpol

ELSEVIER



A comparison of methane emission measurements using eddy covariance and manual and automated chamber-based techniques in Tibetan Plateau alpine wetland

Lingfei Yu^a, Hao Wang^b, Guangshuai Wang^b, Weimin Song^a, Yao Huang^c, Sheng-Gong Li^d, Naishen Liang^e, Yanhong Tang^e, Jin-Sheng He^{a,b,*}



EC

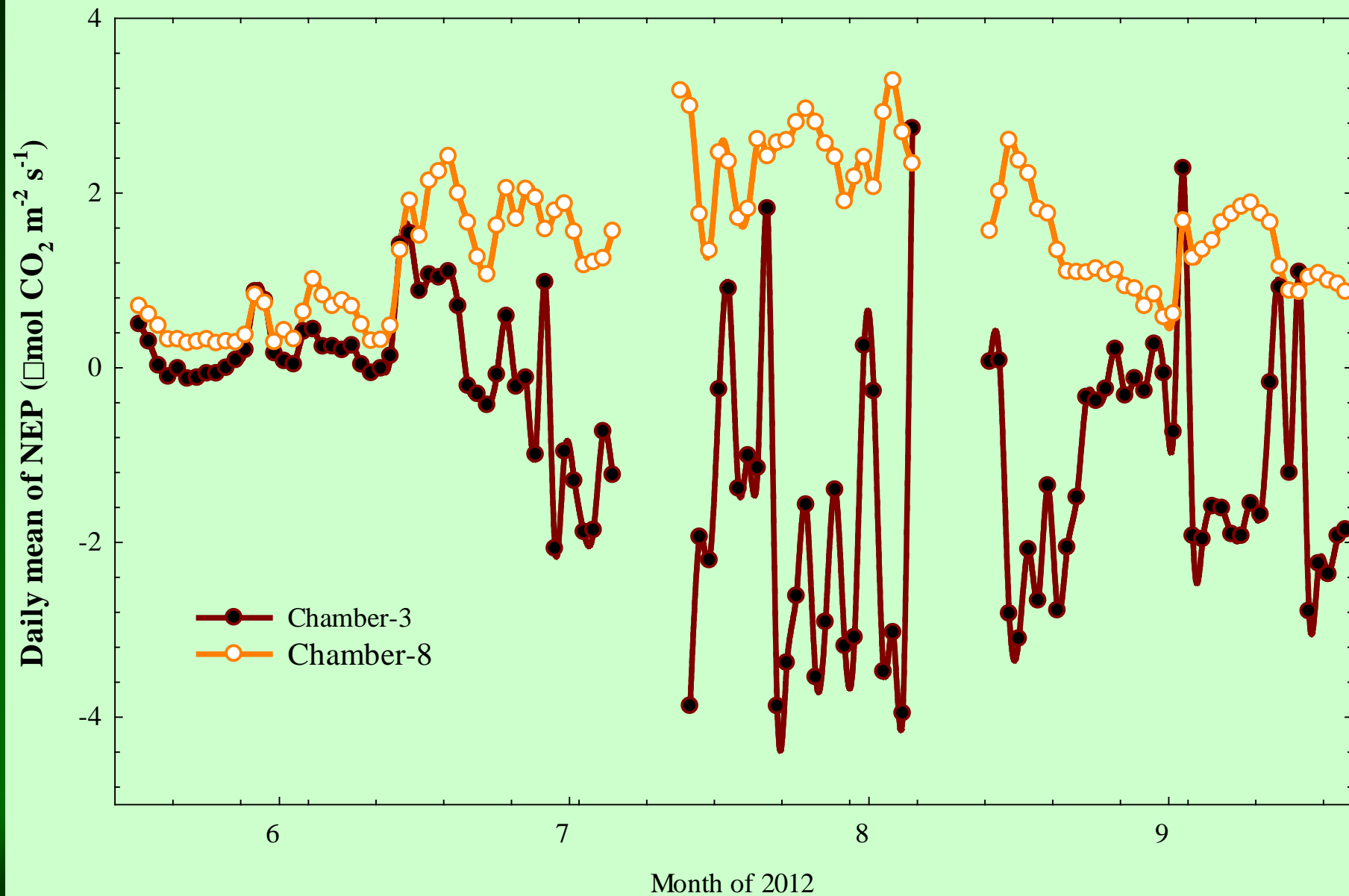
Automated chamber

Chamber-GC

Inner-Mongolian Grassland



Seasonal Change in NEP & Rs at Duolun Grassland Ecosystem



塞罕坝落叶松林生态系统



尖峰岭热带季雨林生态系统



麗江高山針葉林生態系統土壤呼吸增溫實驗

27°14' N, 100°23' E; 3200m asl.

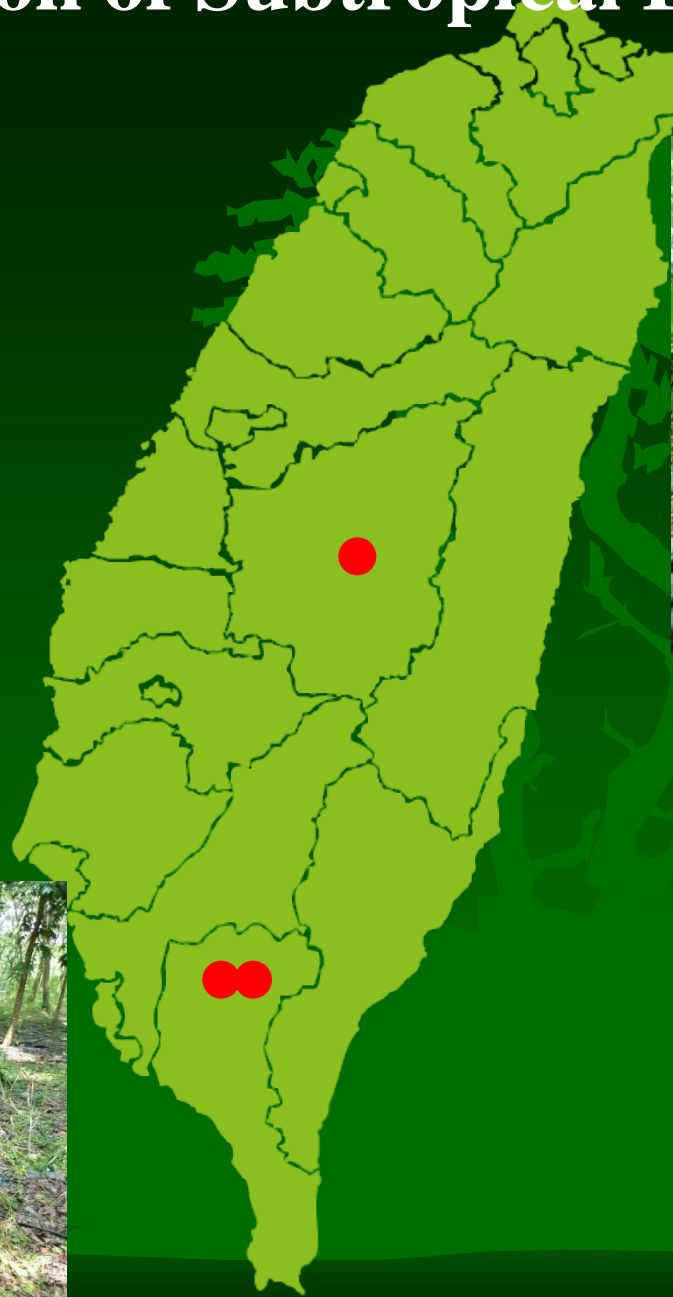


疏勒河上游高原凍土草甸生態系統

Seasonal Permafrost meadow, 4100m, Tibet Plateau



Soil Respiration of Subtropical Forests



Xitou Flux Site



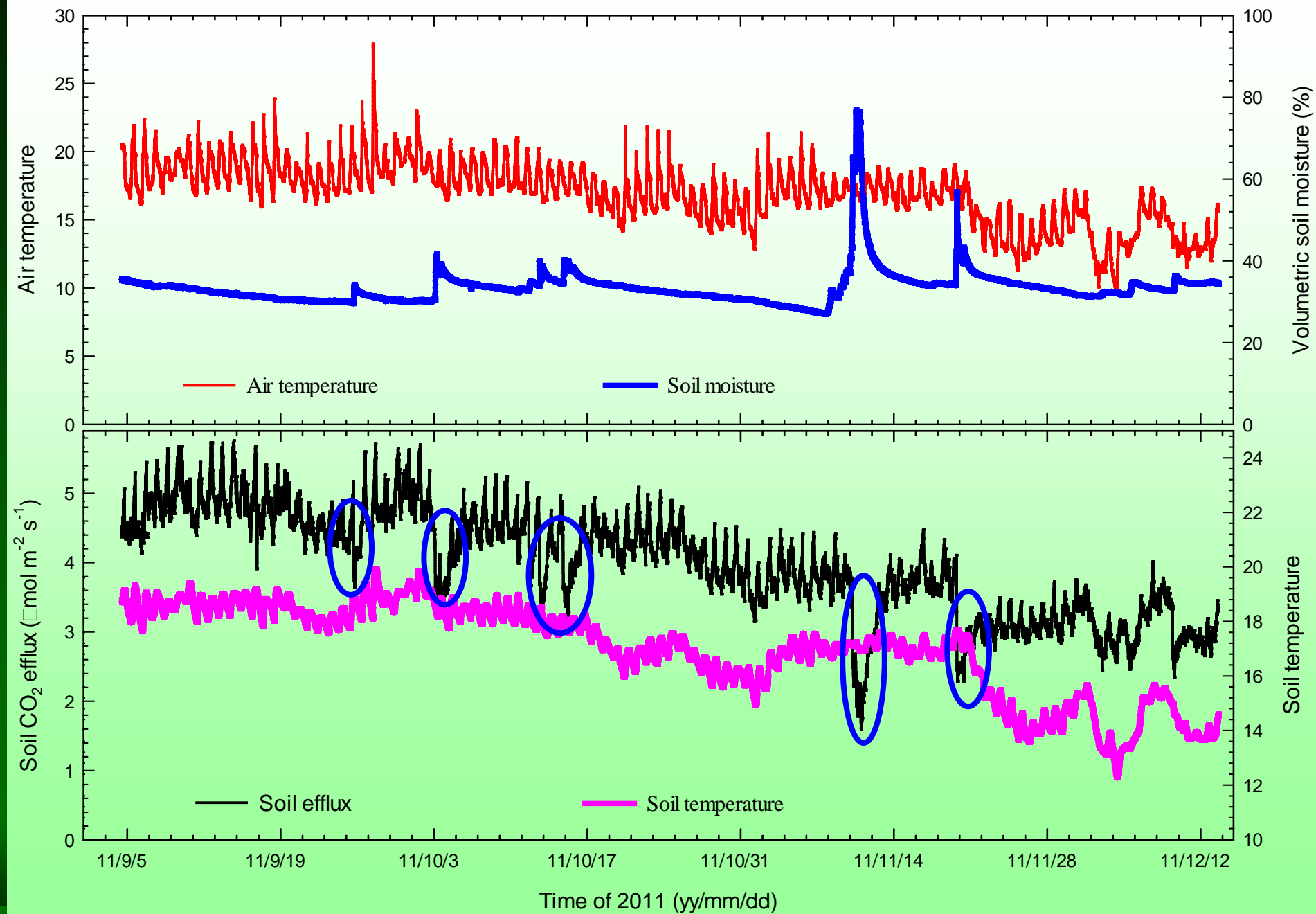
Pingdong Flux Site



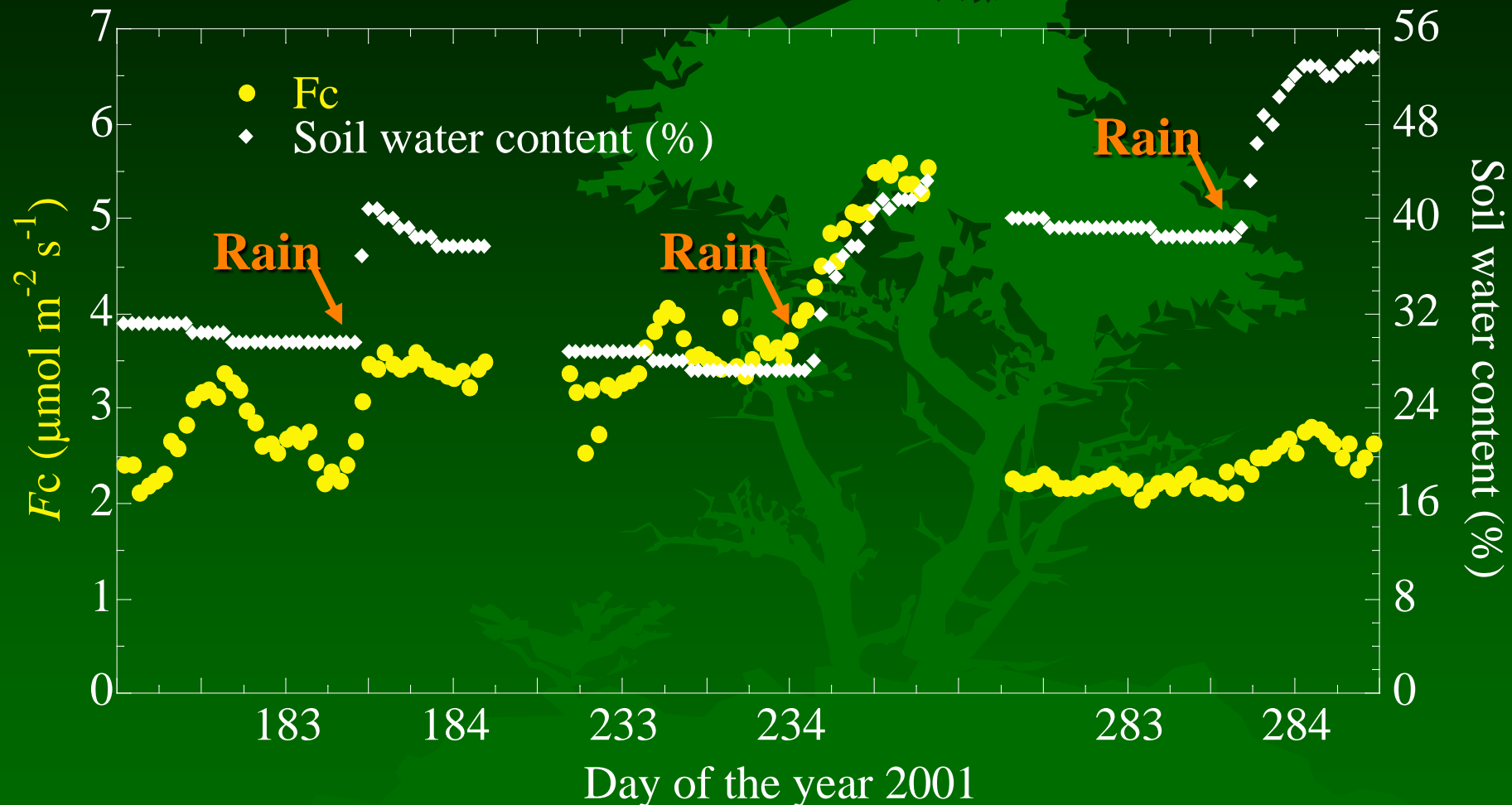
Pingdong Flux Site



Seasonal Changes in Soil CO₂ Efflux at Xitou Flux Site



Is soil respiration regulated by surface soil water content?



Following rainfall, soil respiration increased immediately to a high value. However, the rate had returned to the general lower level by the next day.

Wood Respiration



24 Automated Branch Chambers



Weight: < 200g

Multichannel Automated Chamber Systems



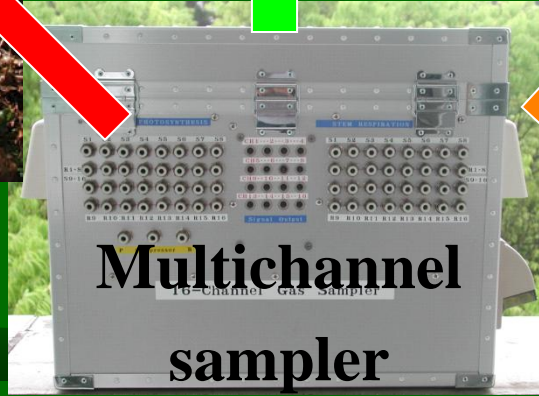
Photosynthesis



Soil efflux

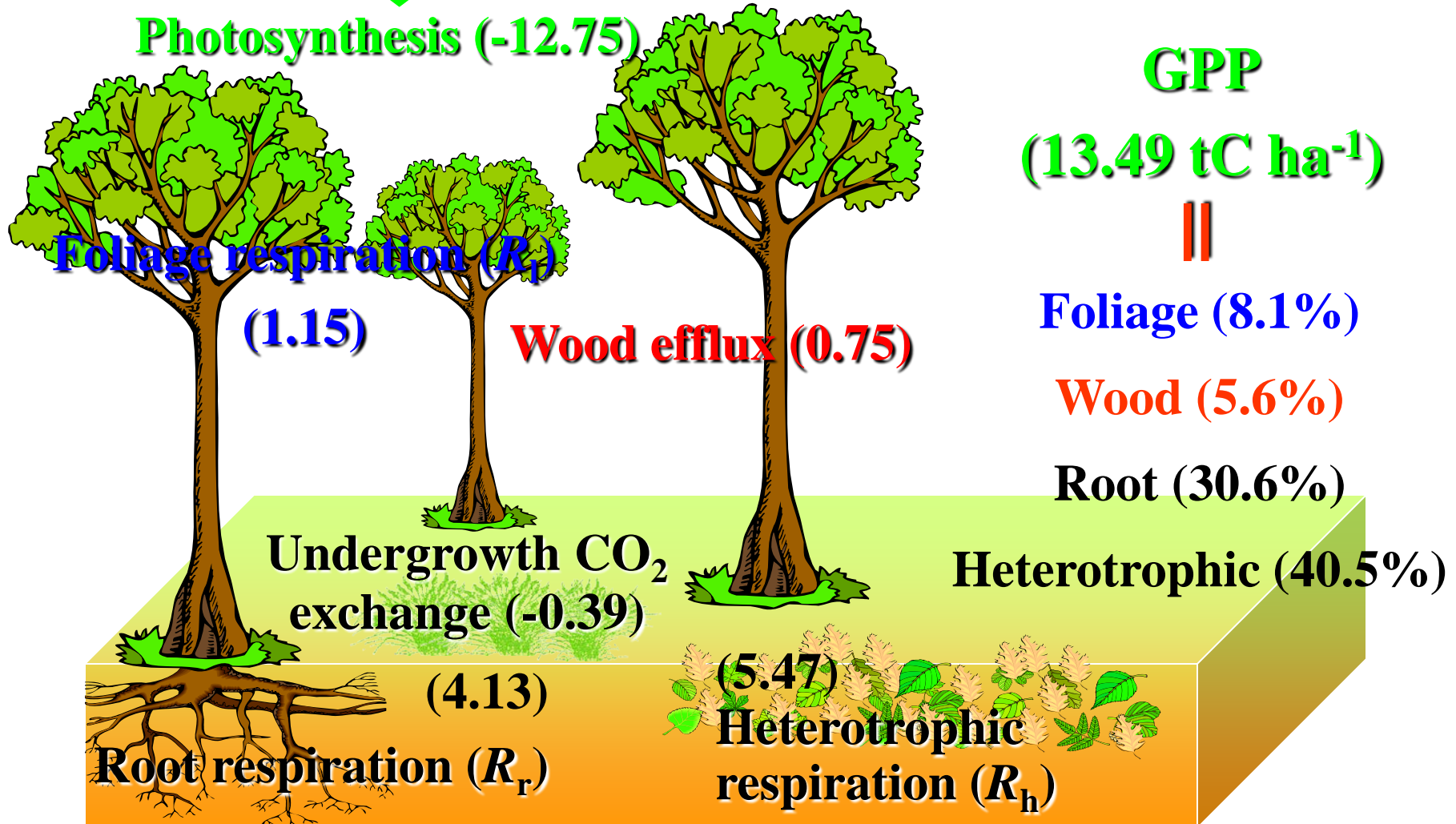


Stem efflux



Larch Forest Carbon Balance

Eddy Covariance (NEP) = 200~550 tC ha⁻¹ y⁻¹



Chamber Network Publications

- 
- Liang et al. 2003. Tree Physiol.**
Liang et al. 2004. Agr. For. Meteorol.
Liang et al. 2005. Phytos
Liang et al. 2005. J. Agr. Meteorol.
Takahashi & Liang. 2007. Geochem. J.
Takahashi, Liang et al. 2008. JGR.
Takagi et al. 2008, GCB
Sano, Hirano, Liang et al. 2010. FEM
Liang, Hirano et al. 2010. BG
Aguilos et al. 2011, BD
Yao et al. 2012, JFR
Tan et al. 2013, JGR
Yu et al. 2013, EP
Aguilos et al. 2013, Tellus B
東ら. 2013,北海道の農業気象
Wang et al. 2013, Jan J For Environ
Aguilos et al. 2014, AFM
阿部ら. 2014, 関東森林研究



Thank you for attentions!

