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Yale-NUIST Center on Atmospheric Environment

Spatial variations of transfer coefficients of heat, water vapor and momentum across Lake Taihu

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Video conference on April 27, 2012

Outline

1. Introduction
2. Theory
3. Sites and instruments
4. Results
5. Discussion

1. Introduction

- Lake is an important land surface type for atmospheric research. The inclusion of lakes in numerical weather prediction (NWP) and climate models improves model performance. A critical issue here concerns the accuracy of the **transfer coefficients** because any bias in them will propagate directly to the flux variables.
 - To identify spatial variations of transfer coefficients of heat, water vapor and momentum across Lake Taihu;
 - To compare the transfer coefficients of Lake Taihu with physical ocean parameterization.

(Rouse et al. 2003, 2005; Long et al. 2007; Downing et al. 2006; Herderson-Sellers 1986; Bonan 1995; Betts and Ball 1997; Liu et al. 2005).

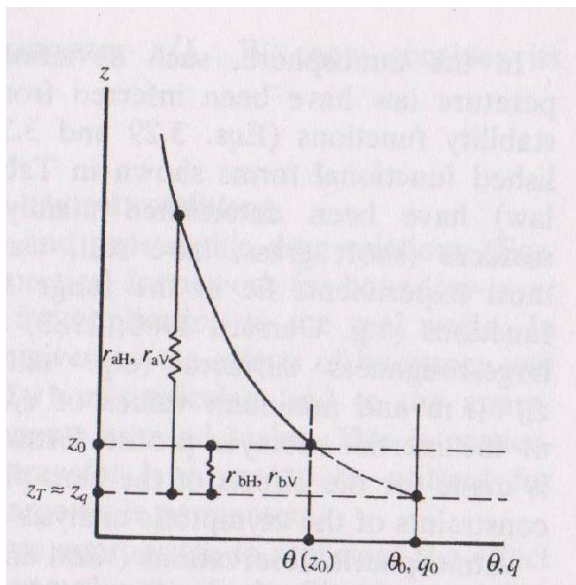
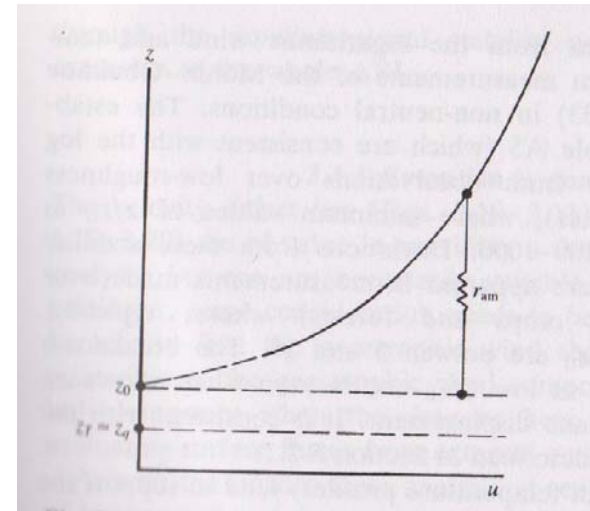
2. Theory

2.1 ROUGHNESS LENGTH VS TRANSFER COEFFICIENTS

Roughness length

- **Aerodynamic roughness length** (z_0) is a parameter of some vertical wind profile equations that model the horizontal mean wind speed near the ground; in the log wind profile, it is equivalent to the height at which the **wind speed** theoretically becomes **zero**.

$$ku / u_{*0} = \ln(z / z_0)$$



- **Scalar roughness length:** the surface-layer temperature and humidity profiles are defined in the equations with scalar roughness lengths z_T and z_q replacing z_0 in the wind profile relation.

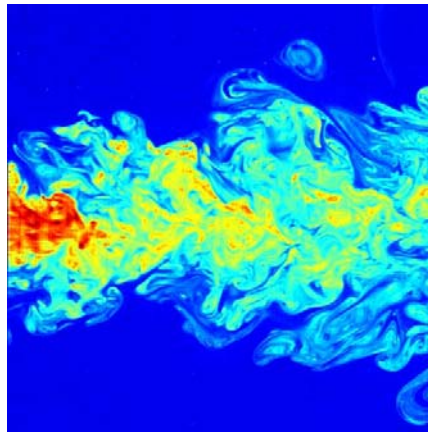
$$k(q - q_0) / q_{*0} = \ln(z / z_q) - \Psi_w(\zeta)$$

$$k(\theta_v - \theta_0) / \theta_{v*0} \approx \ln(z / z_T) - \Phi_H(\zeta)$$

Aerodynamic vs scalar roughness length

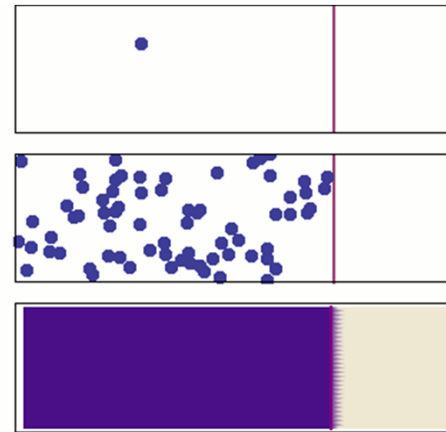
Conceptually, the distinction between z_T , z_q and z_0 is suggested by the consideration of **transport mechanisms** for heat, water vapour and momentum in the presence of aerodynamically rough flow close to the surface.

Turbulence



The transfer of **momentum** is effected by **pressure fluctuations** in the turbulent wakes behind the roughness elements.

Molecular diffusion



Heat and **water vapour** must ultimately be transferred by **molecular diffusion** across the interfacial sublayer.

Transfer coefficients - the surface roughness

$$C_{DN} = k^2 / [\ln(z / z_0)]^2$$

Momentum roughness

$$C_{EN} = k^2 / [\ln(z / z_0) \ln(z / z_q)]$$

Roughness length for water vapor

$$C_{HN} = k^2 / [\ln(z / z_0) \ln(z / z_T)]$$

Roughness length for temperature

The neutral transfer coefficients at the height of 10 m above the water surface

$$C_{D10N} = k^2 / [\ln(10 / z_0)]^2$$

$$C_{E10N} = k^2 / [\ln(10 / z_0) \ln(10 / z_q)]$$

$$C_{H10N} = k^2 / [\ln(10 / z_0) \ln(10 / z_T)]$$

Mass transfer equations

Momentum flux

$$\tau = \rho_a C_D u^2$$

Drag coefficient

Latent heat flux

$$LE = \rho_a L_v C_E u (q_s - q_a)$$

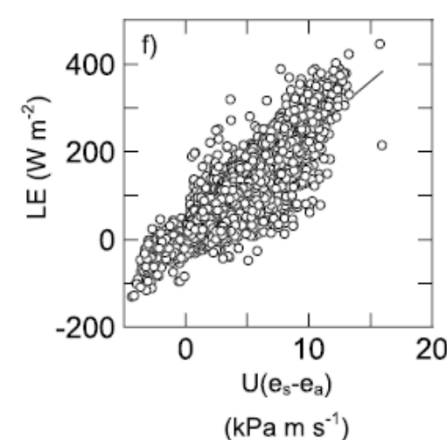
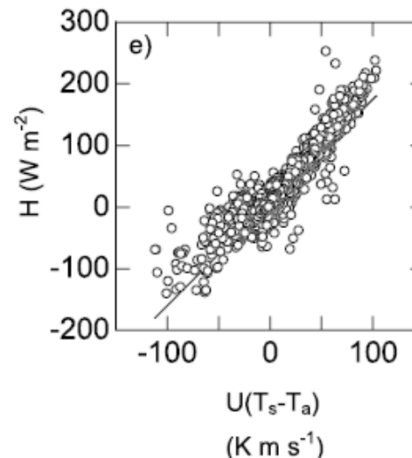
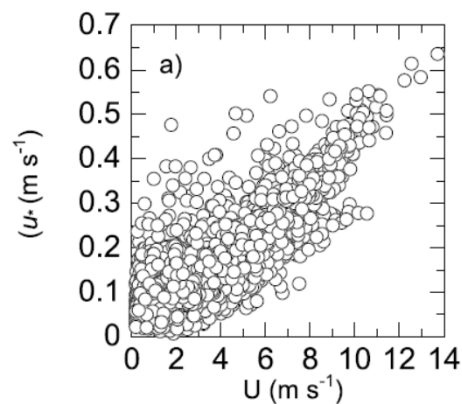
Transfer coefficient for moisture

Sensible heat flux

$$H = \rho_a c_p C_H u (\theta_s - \theta_a)$$

Transfer coefficient for heat

The coefficients C_{DN} , C_{EN} and C_{HN} were optimized by minimizing the bias between the observation and calculation of τ , LE and H .



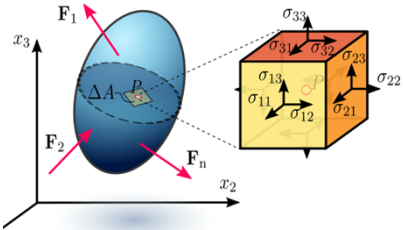
(Liu et al. 2009)

2. Theory

2.2 GARRATT/BRUTSAERT MODEL

Smooth vs Rough surface over water

Shear stress



Turbulent boundary layer

Viscous sublayer δ_1

Smooth surface

Aerodynamically rough surface

$$h_r \gg \delta_1$$

$$u_* h_r / \nu > 75$$

Aerodynamically smooth surface

$$h_r \ll \delta_1$$

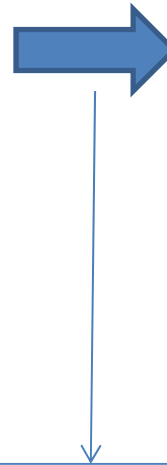
$$u_* h_r / \nu < 5$$

h_r – the height of the roughness elements

- Aerodynamically smooth:
 $Re_* \approx 0.11$



- Fully rough:
 $Re_* > 2$



- **Wave breaking:** a transfer of energy from the already saturated wave components (ripples) to turbulence.
- $Re_* \approx 2, u_{*0} \approx 0.23 \text{ m s}^{-1} (u_{10} \approx 5.5 \text{ m s}^{-1})$

The aerodynamic roughness length over sea

- In very light winds --> aerodynamically smooth surface

$$z_0 \approx 0.11\nu / u_{*0}$$

- In rough flow conditions

$$z_0 = \alpha_c u_{*0}^2 / g$$

$$z_0 = (\alpha_c u_{*0}^2 / g) (\mu_w u_{*0} / \sigma_t)^{m-2}$$

Drag coefficients over sea

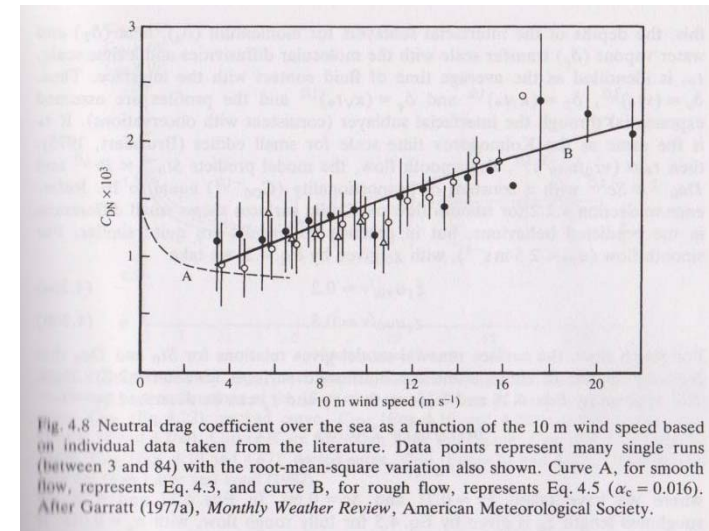
Aerodynamically smooth flow

$$C_{DN} = k^2 / [\ln(u_{*0} z / 0.11\nu)]^2$$

Rough flow

$$C_{DN} = k^2 / [\ln(zg / \alpha_c u_{*0}^2)]^2$$

Transitional regime: $2.5 < u_{10} < 5.5$ /s



Bulk parameterization for heat transfer

$$C_{HN} = C_{DN}^{1/2} / (B_H^{-1} + C_{DN}^{-1/2})$$

$$B_H^{-1} = St_0^{-1} - C_{D0}^{-1/2} = k^{-1} \ln(z_0 / z_T)$$

For smooth flow

$$z_T u_{*0} / \nu \approx 0.2$$

For fully rough flow

$$\ln(z_0 / z_T) = 2.48 Re_*^{1/4} - 2$$

$$Re_* = u_{*0} z_0 / \nu$$

Bulk parameterization for water vapor transfer

$$C_{EN} = C_{DN}^{1/2} / (B_V^{-1} + C_{DN}^{-1/2})$$

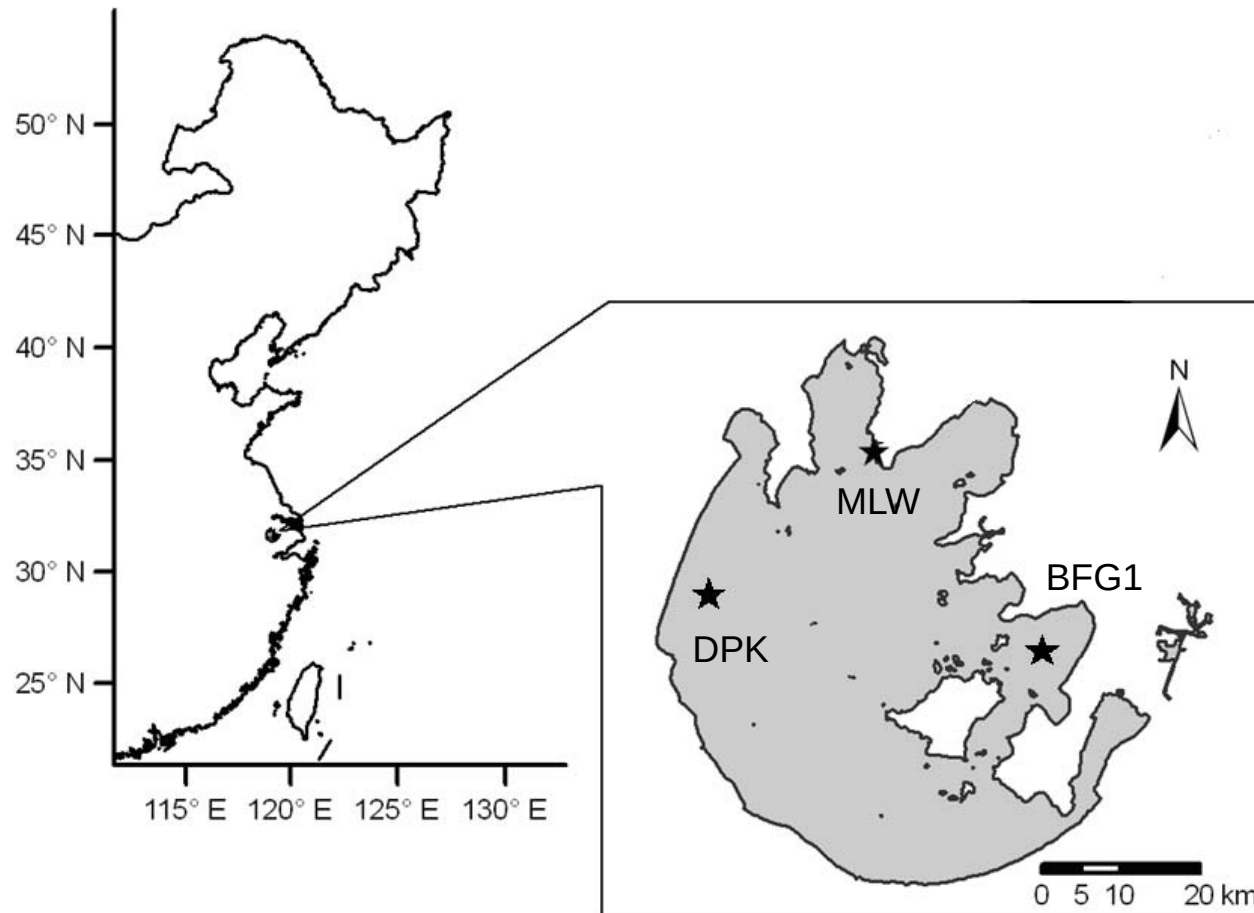
$$B_V^{-1} = Da_0^{-1} - C_{D0}^{-1/2} = k^{-1} \ln(z_0 / z_q)$$

smooth flow $z_q u_{*0} / \nu \approx 0.3$

rough flow $\ln(z_0 / z_q) = 2.28 Re_*^{1/4} - 2$

$$Re_* = u_{*0} z_0 / \nu$$

3. Sites and instruments



MLW site
梅梁湾站

Wind direction: 200-315°

AC

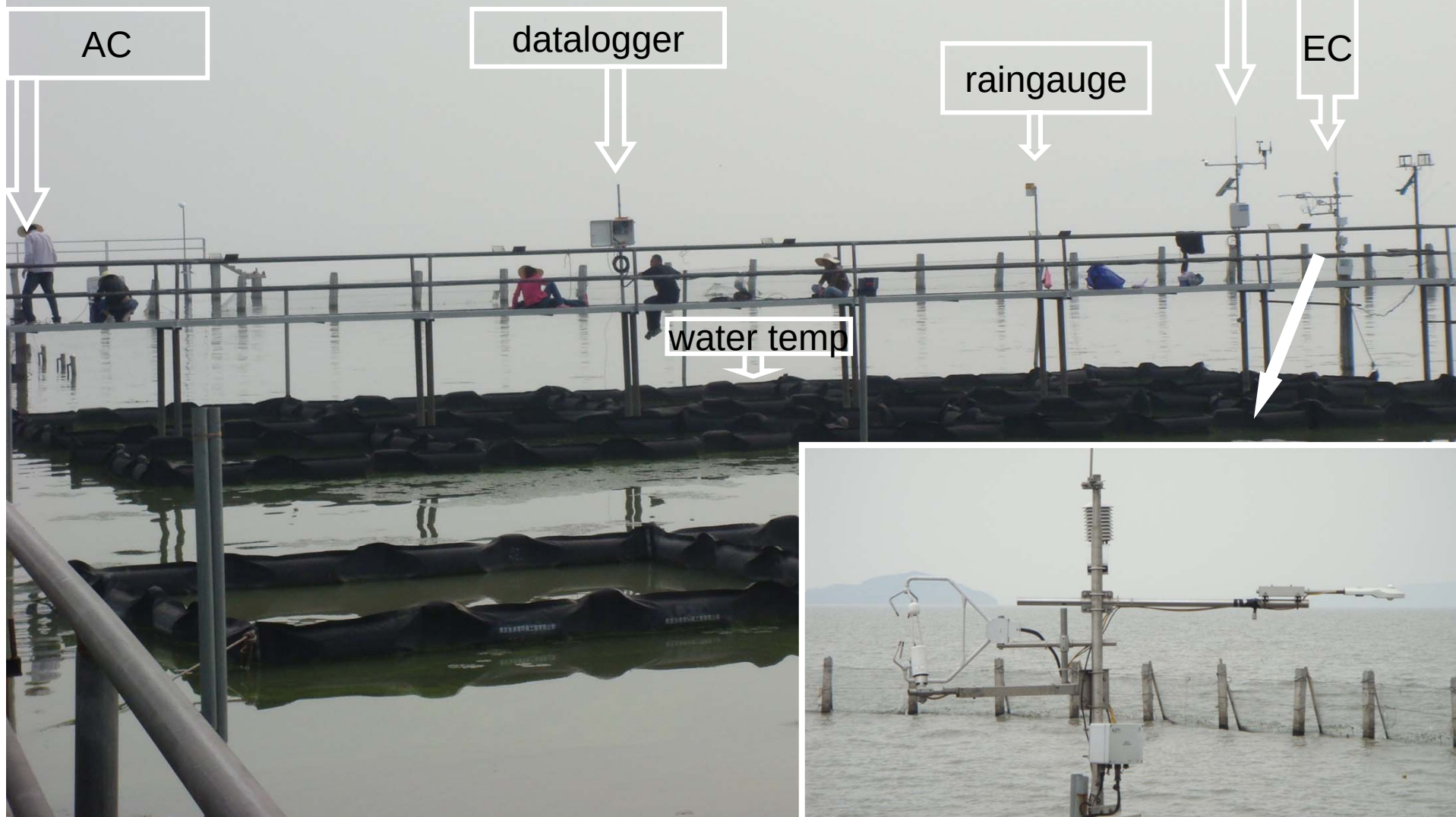
datalogger

microclimate

raingauge

EC

water temp



DPK site
大浦口站

Wind direction: 100-280°



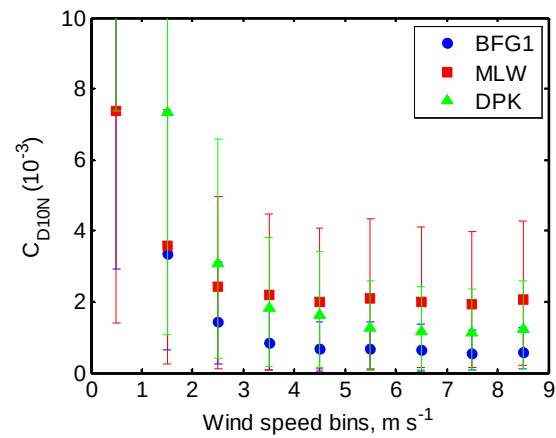
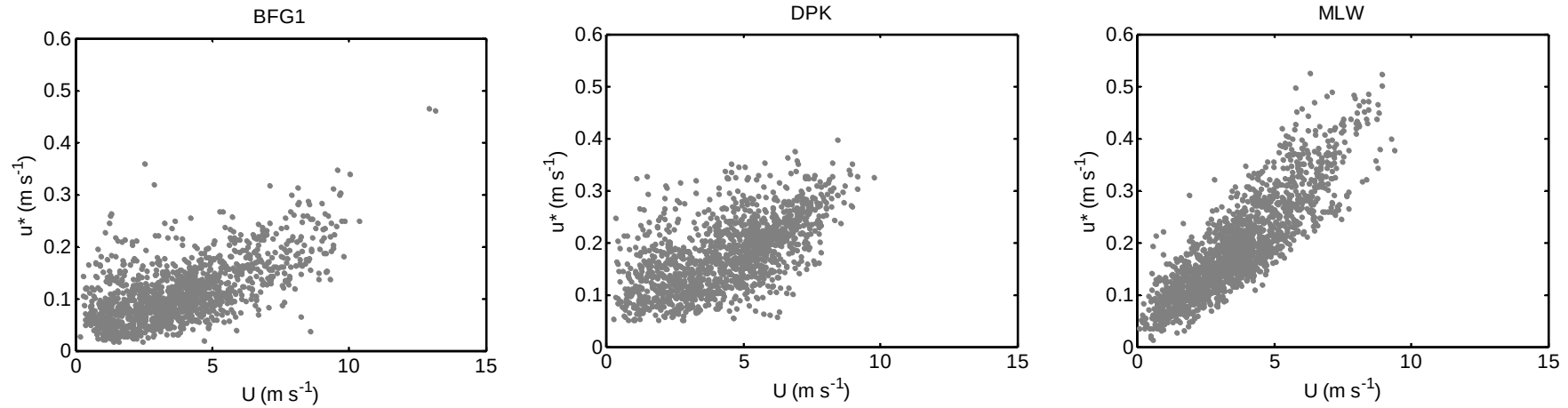
Wind direction: 135-315°



BFG1 site
1号避风港

4. RESULTS

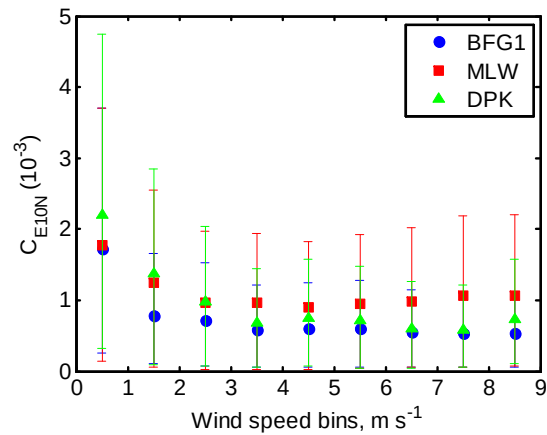
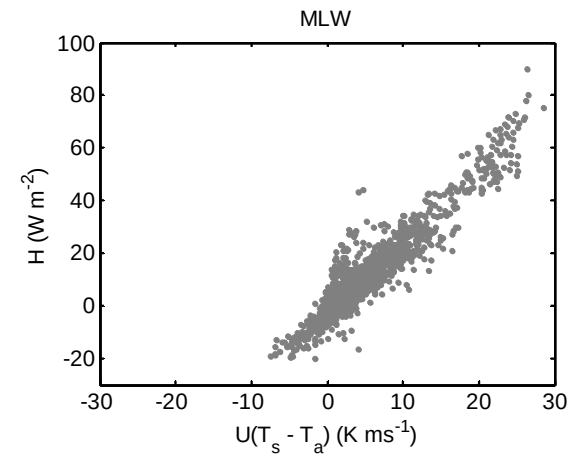
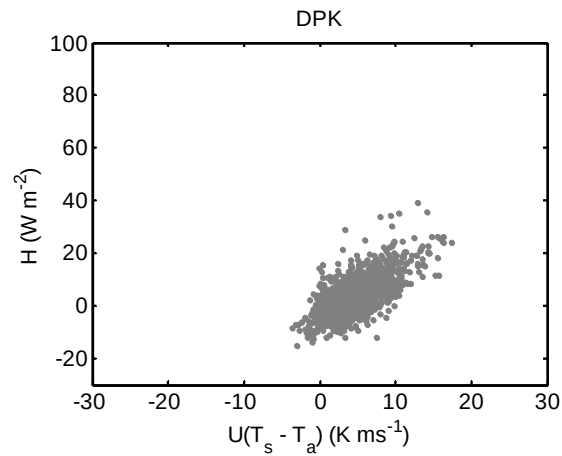
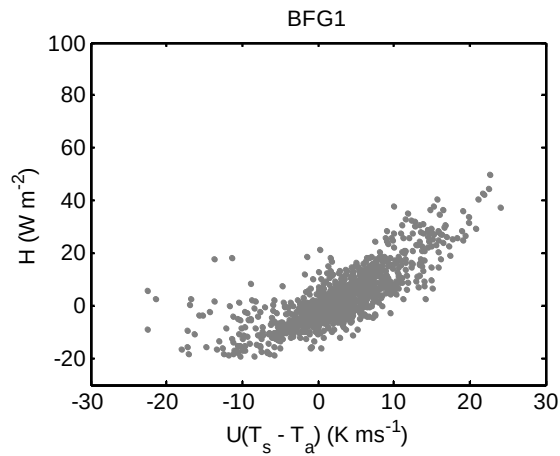
Drag coefficients



MLW > DPK > BFG1

$10^3 C_{D10N}$	BFG1	DPK	MLW
$u < 1$	12.05	19.62	7.36
$1 < u < 2$	3.34	7.34	3.58
$2 < u < 3$	1.44	3.08	2.42
$3 < u < 4$	0.84	1.82	2.19
$4 < u < 5$	0.69	1.64	1.99
$5 < u < 6$	0.68	1.26	2.11
$6 < u < 7$	0.65	1.18	1.99
$7 < u < 8$	0.56	1.15	1.92
$8 < u < 9$	0.59	1.23	2.05

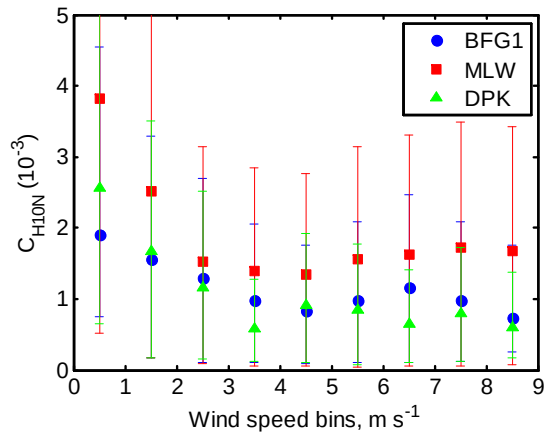
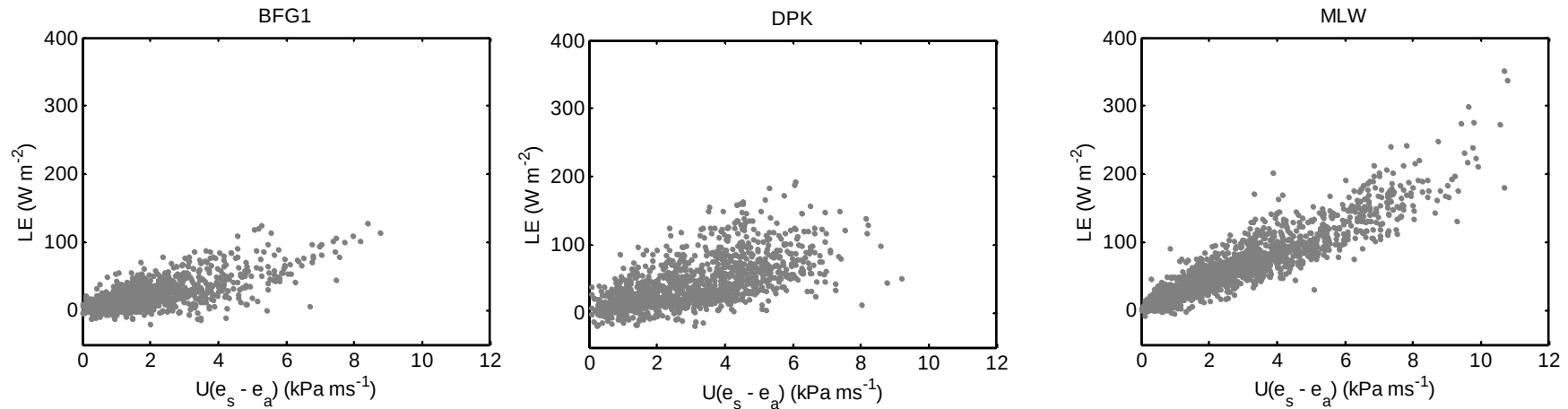
Heat transfer coefficients



MLW > BFG1 > DPK

$10^3 C_{H10N}$	BFG1	DPK	MLW
$u < 1$	1.90	2.56	7.36
$1 < u < 2$	1.56	1.67	3.58
$2 < u < 3$	1.29	1.17	2.42
$3 < u < 4$	0.98	0.58	2.19
$4 < u < 5$	0.83	0.92	1.99
$5 < u < 6$	0.99	0.85	2.11
$6 < u < 7$	1.17	0.66	1.99
$7 < u < 8$	0.98	0.80	1.92
$8 < u < 9$	0.74	0.60	2.05

Water vapor transfer coefficients



MLW > BFG1 > DPK

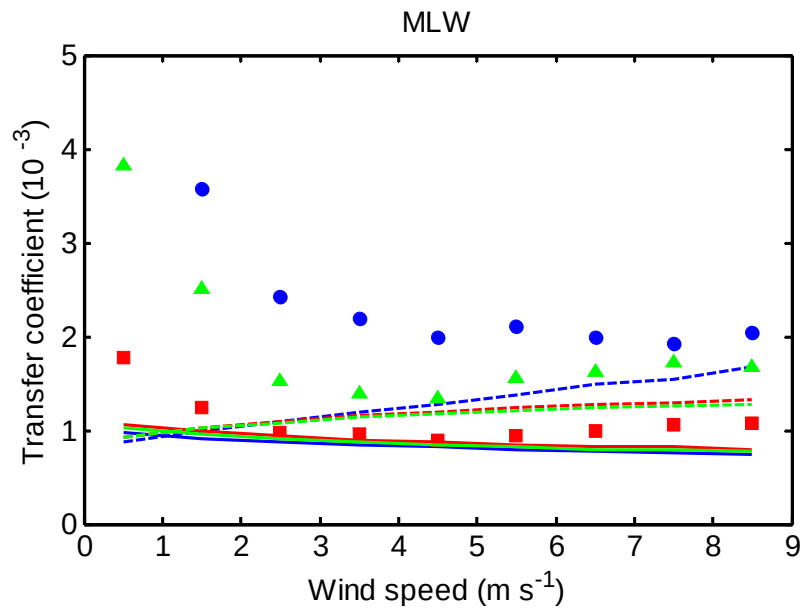
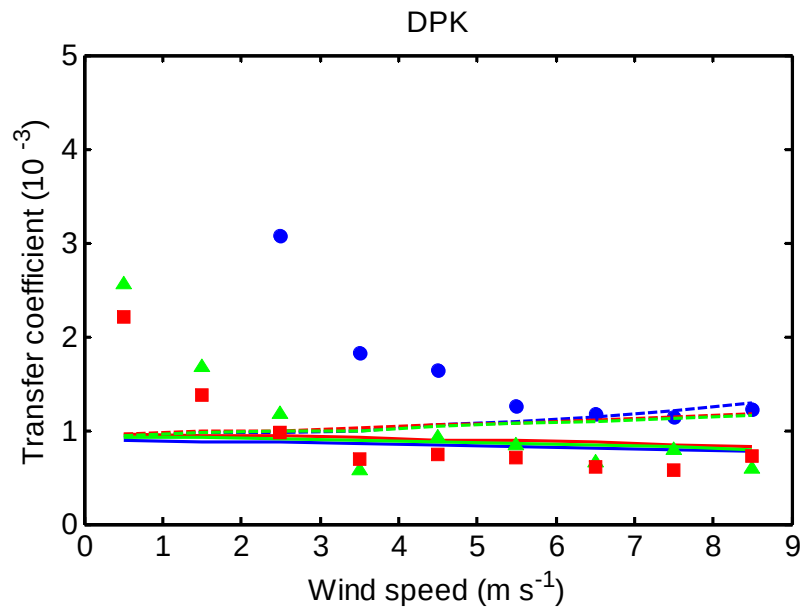
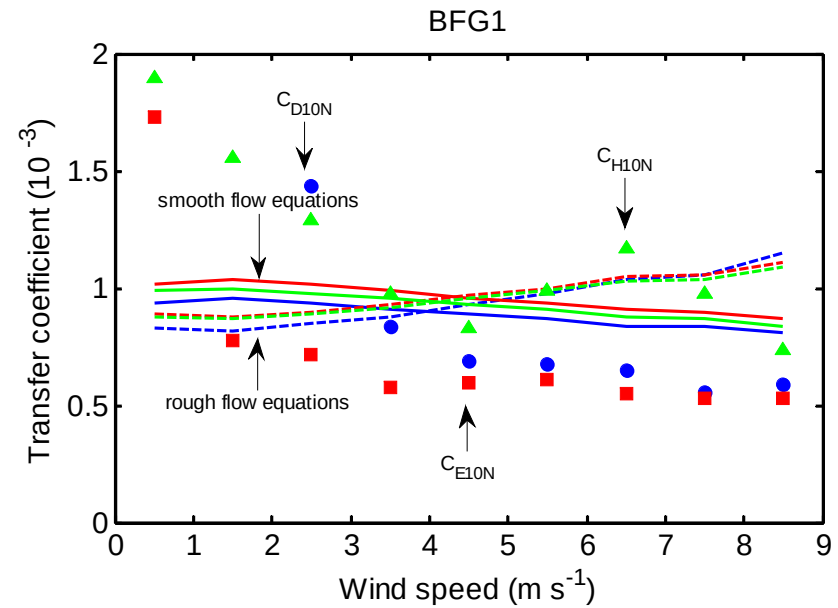
$10^3 C_{E10N}$	BFG1	DPK	MLW
$u < 1$	1.73	2.21	1.78
$1 < u < 2$	0.78	1.38	1.25
$2 < u < 3$	0.72	0.98	0.97
$3 < u < 4$	0.58	0.69	0.96
$4 < u < 5$	0.60	0.75	0.90
$5 < u < 6$	0.61	0.71	0.95
$6 < u < 7$	0.55	0.61	0.99
$7 < u < 8$	0.53	0.58	1.06
$8 < u < 9$	0.53	0.73	1.07

Garratt's model versus optimization results

BFG1: $C_{H10N} > C_{D10N} > C_{E10N}$

DPK: $C_{D10N} > C_{H10N} \sim C_{E10N}$

MLW: $C_{D10N} > C_{H10N} > C_{E10N}$



5. Discussion

- Low wind speed: how to identify the transfer coefficients?
- What drives the spatial variation of transfer coefficient on Lake Taihu? Fetch, depth, or submerged vegetation?

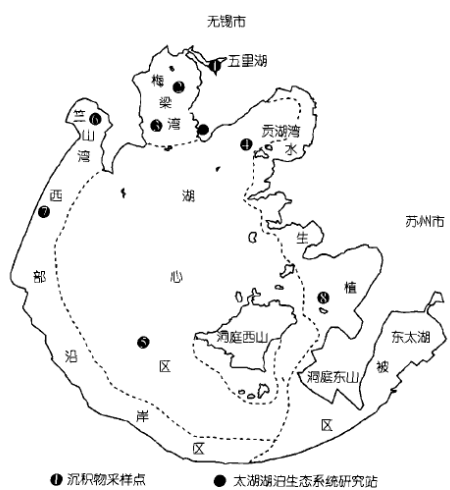
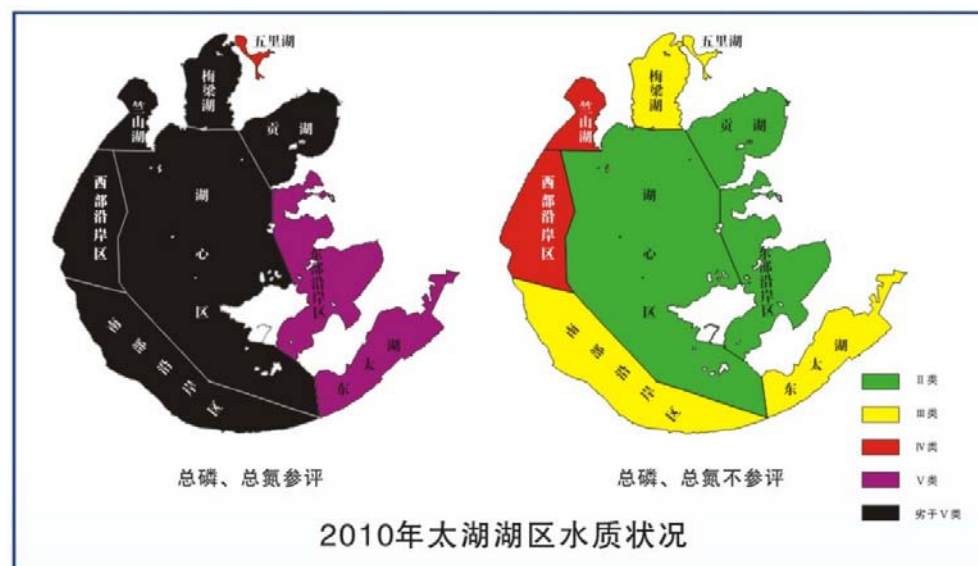


图1 太湖湖区类型与采样点分布图



BFG1(1号避风港)通量站

Location: $120^{\circ} 24'01''\text{E}$, $31^{\circ} 10'28''\text{N}$

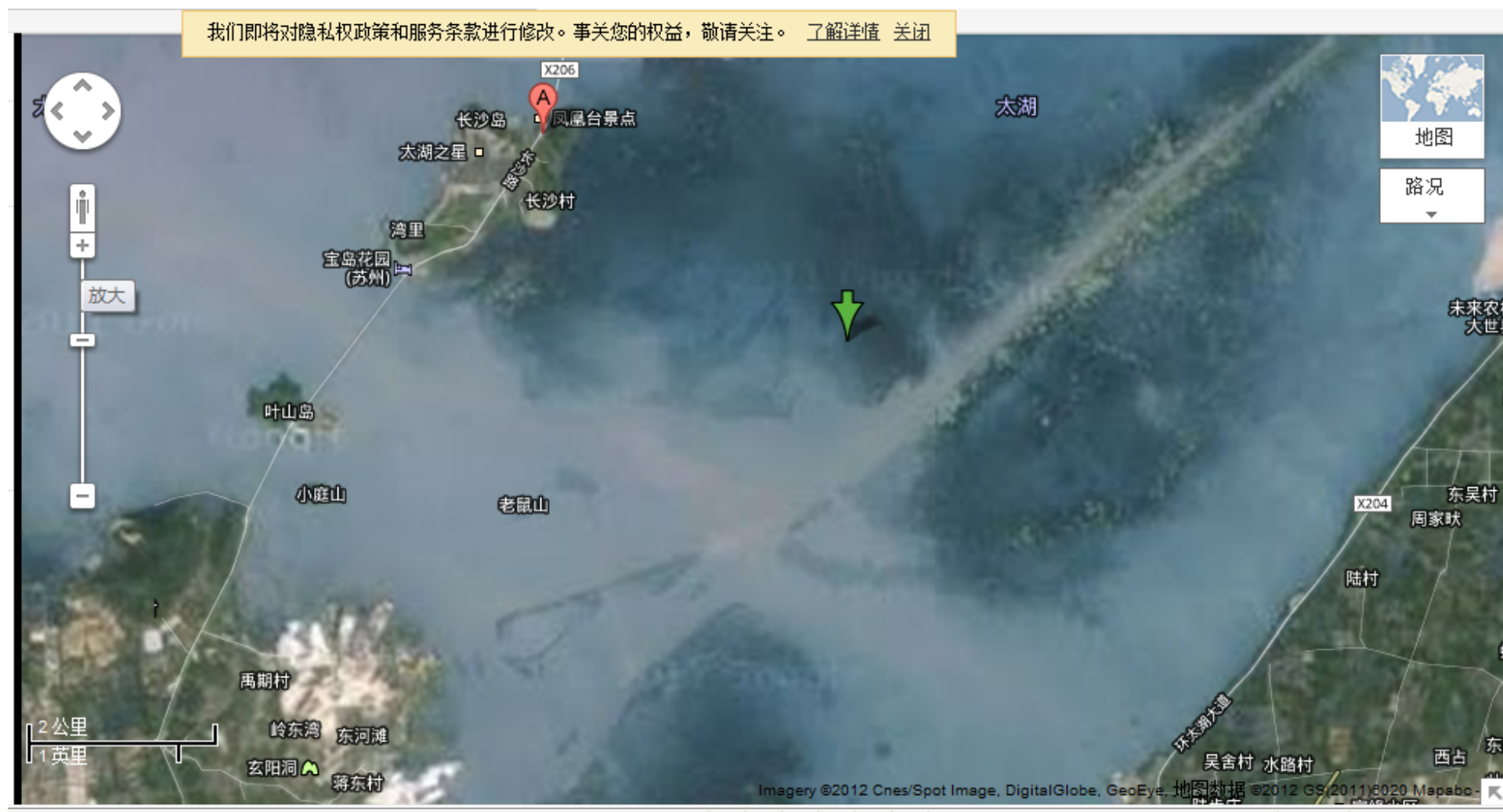
Started from: December 15, 2011

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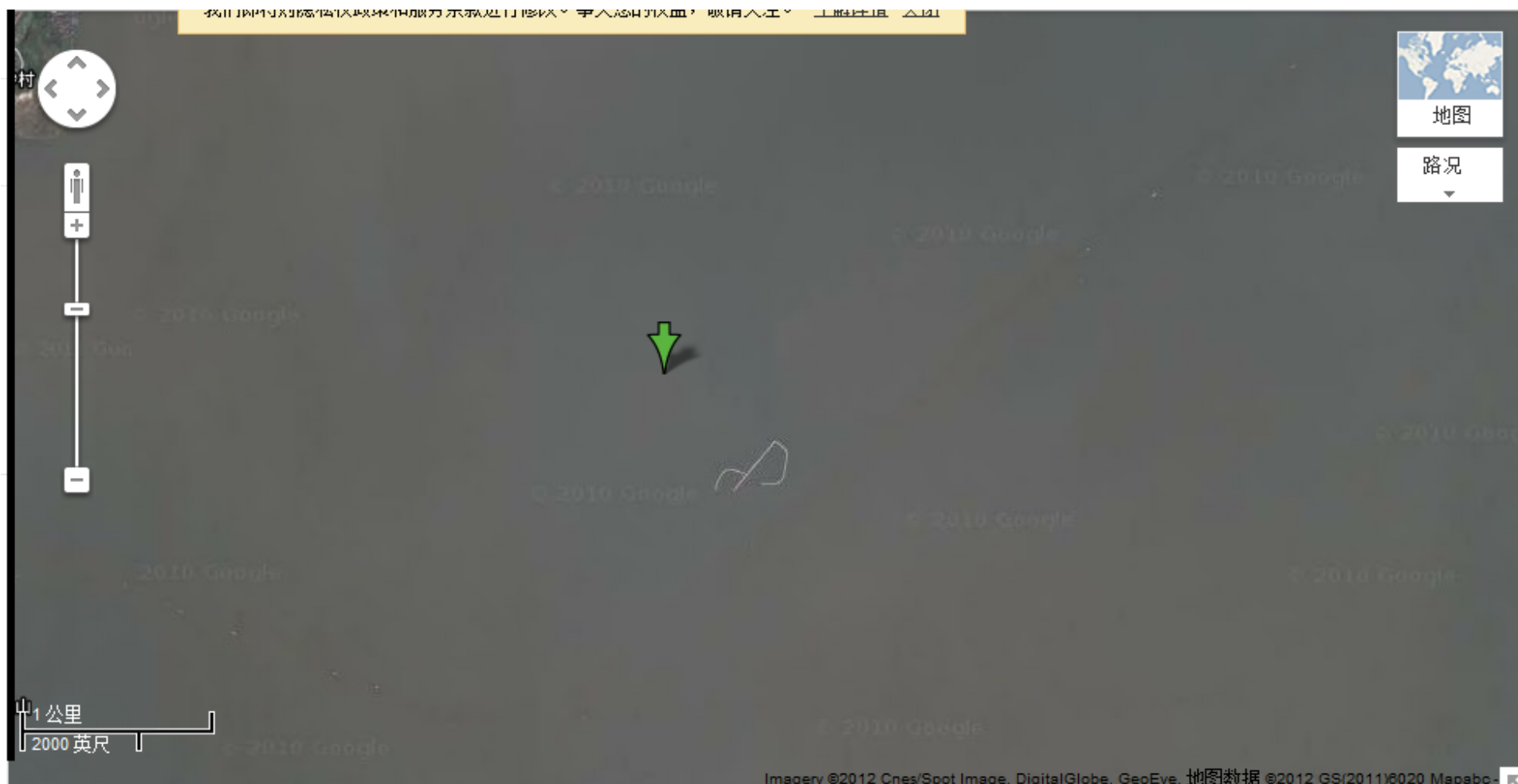


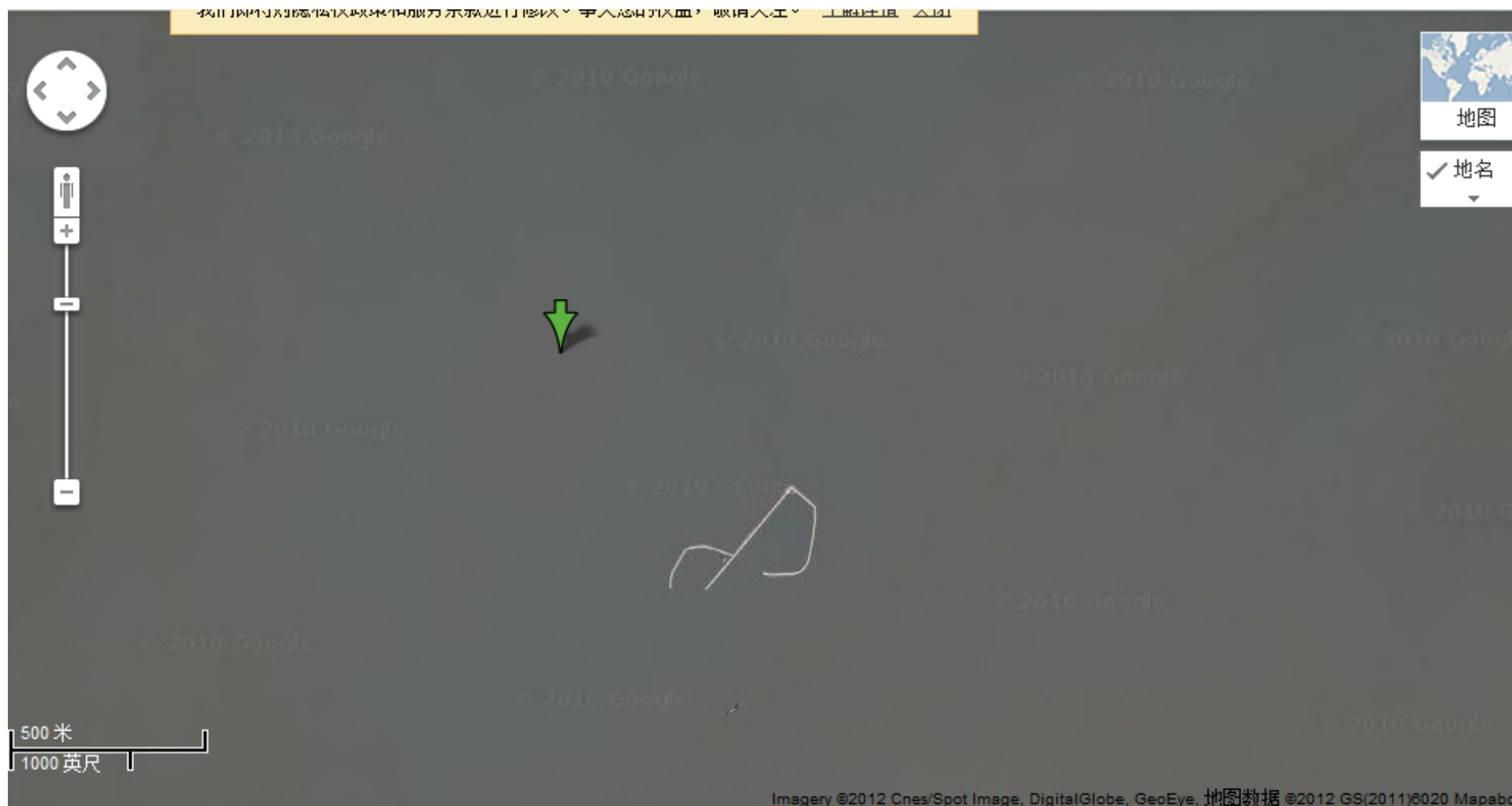


3km

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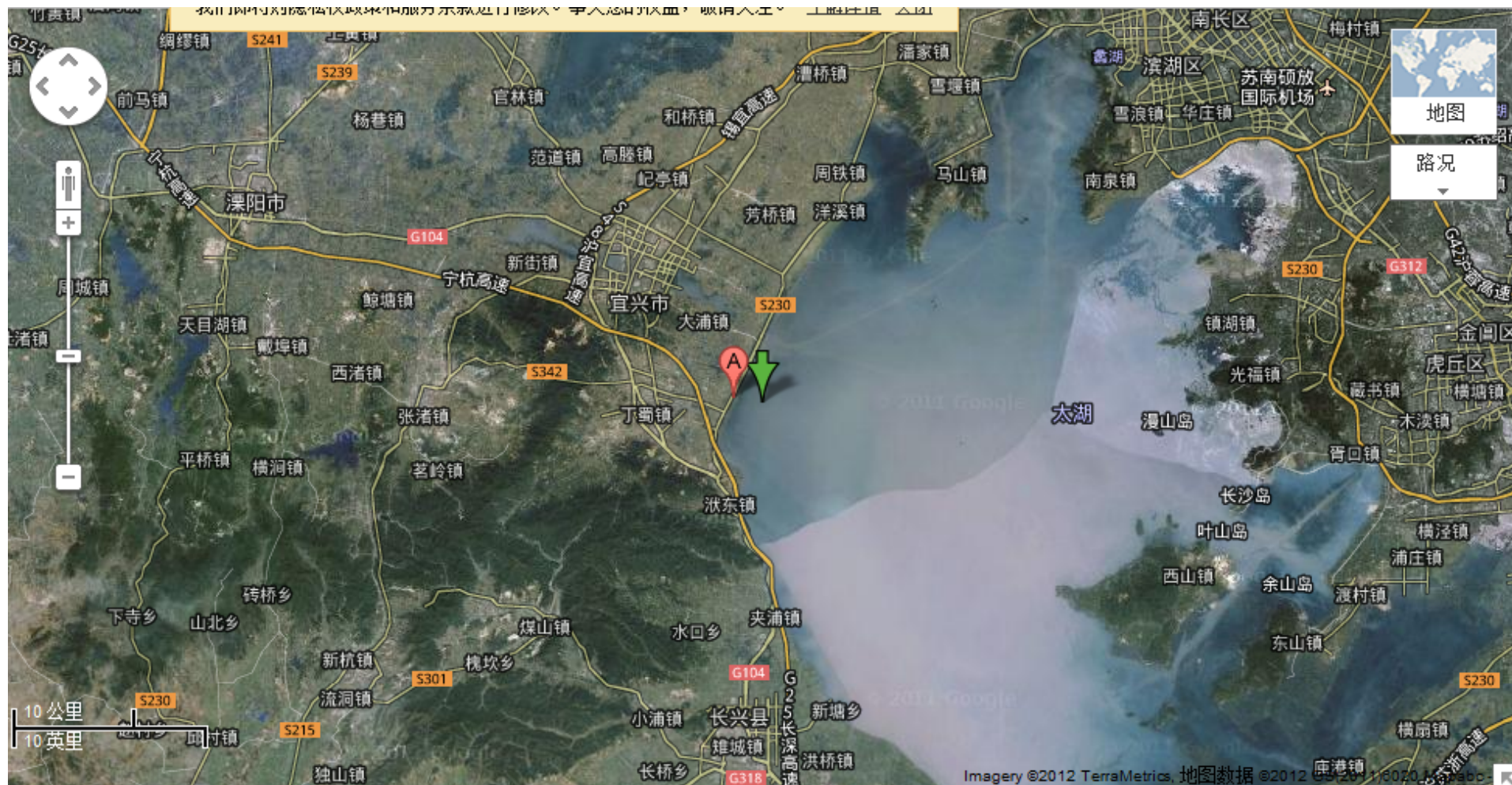


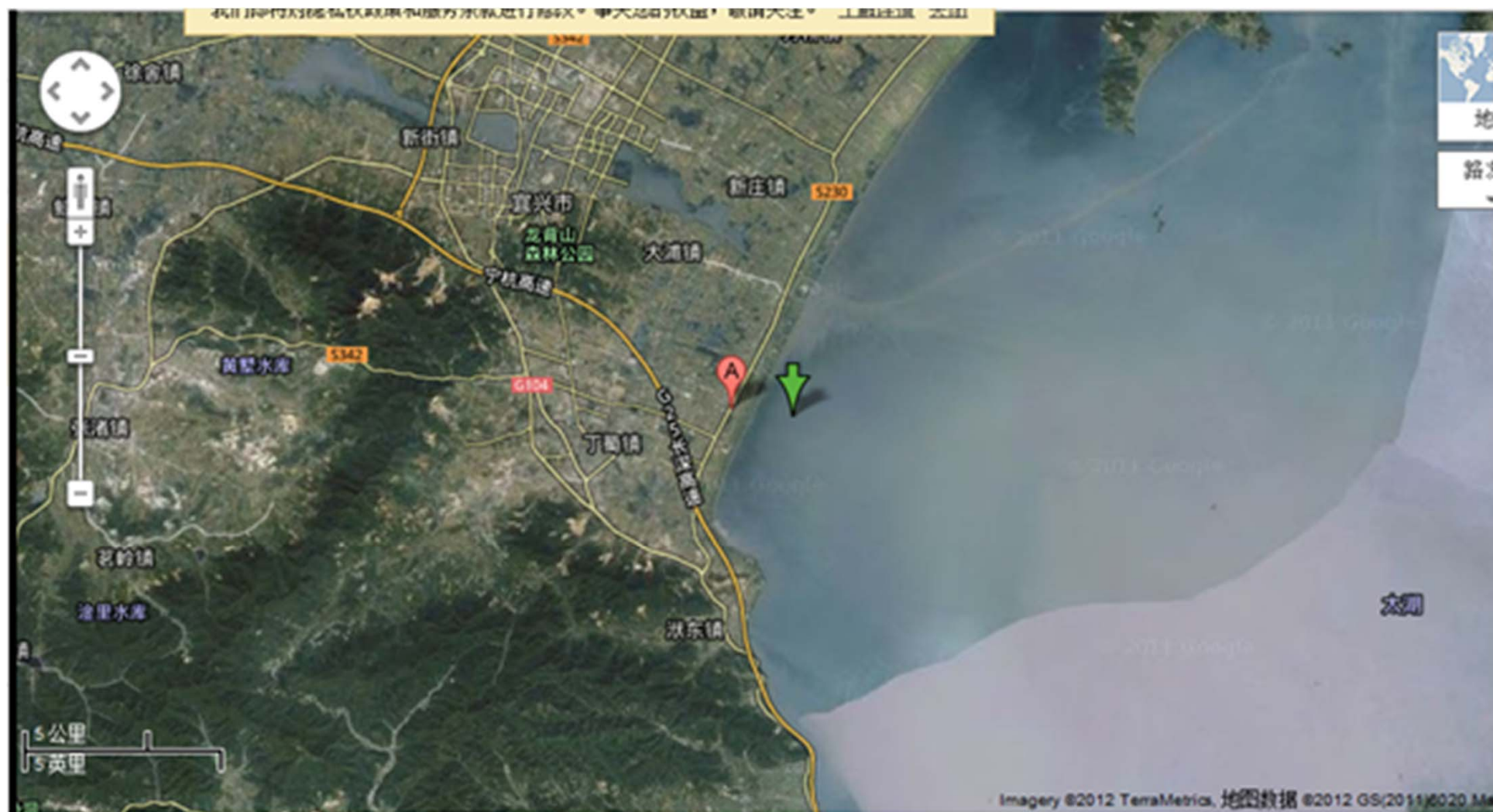


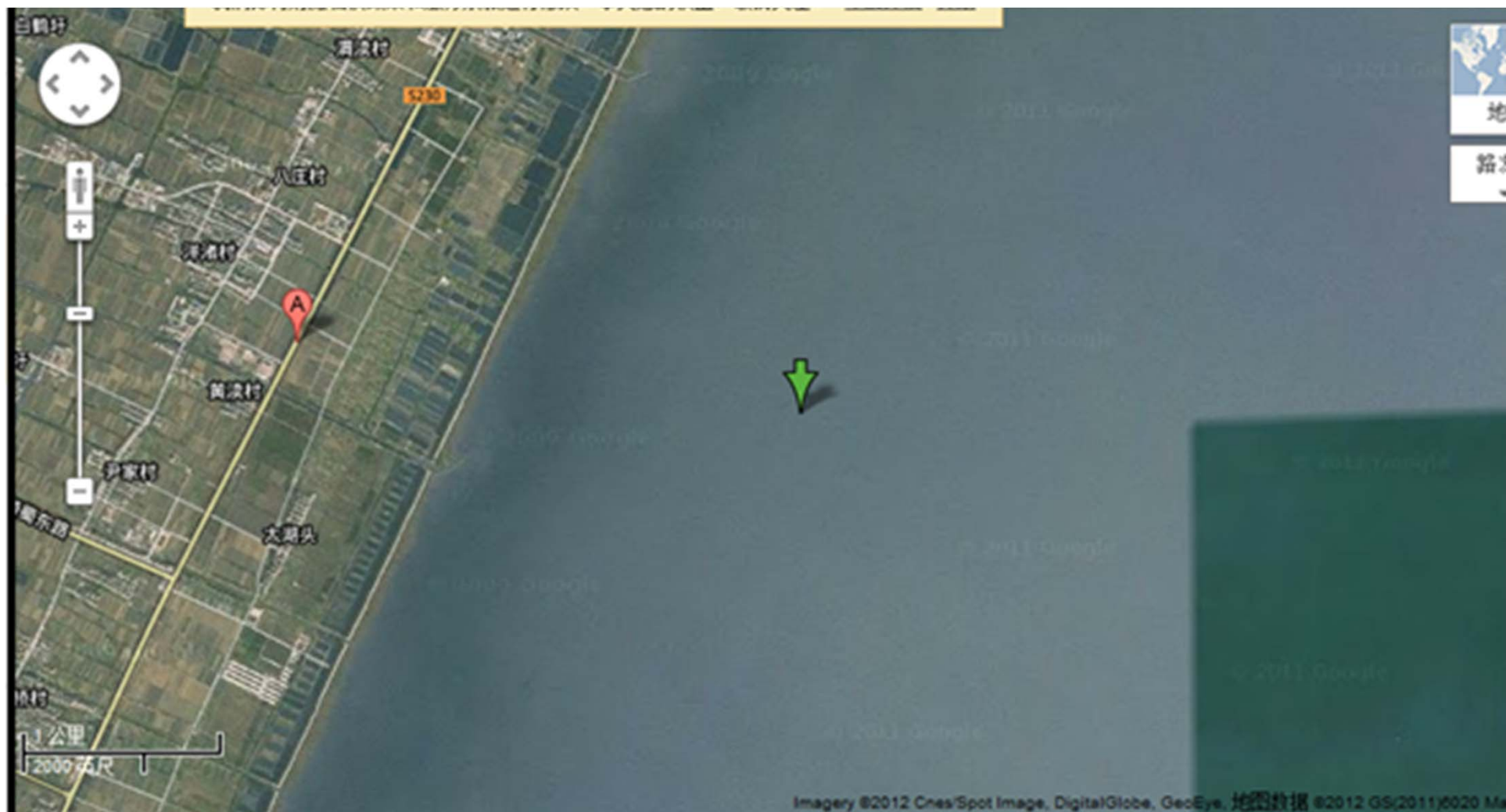
600m

Description of Dapukou(DPK, 大浦口) flux site

- Radiation measurement was started on December 3, 2010
- Eddy covariance, meteorology & water temperature was started from August 17, 2011
- Location: $119^{\circ} 55'52.3''\text{E}$, $31^{\circ} 15'57.8''\text{N}$







1.5km

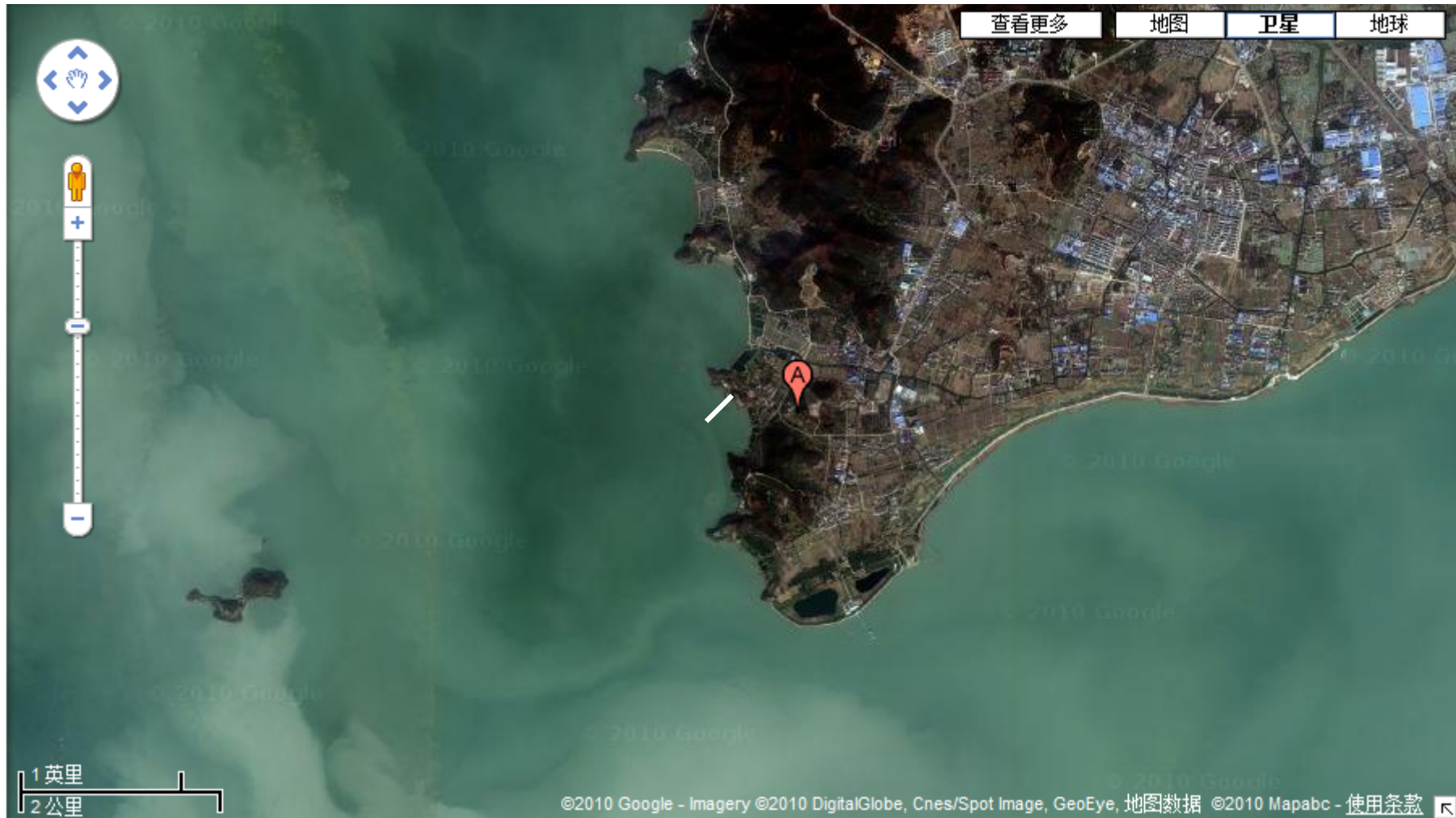
MLW (梅梁湾) site description

Location: $31^{\circ} 24'N$, $120^{\circ} 13'E$



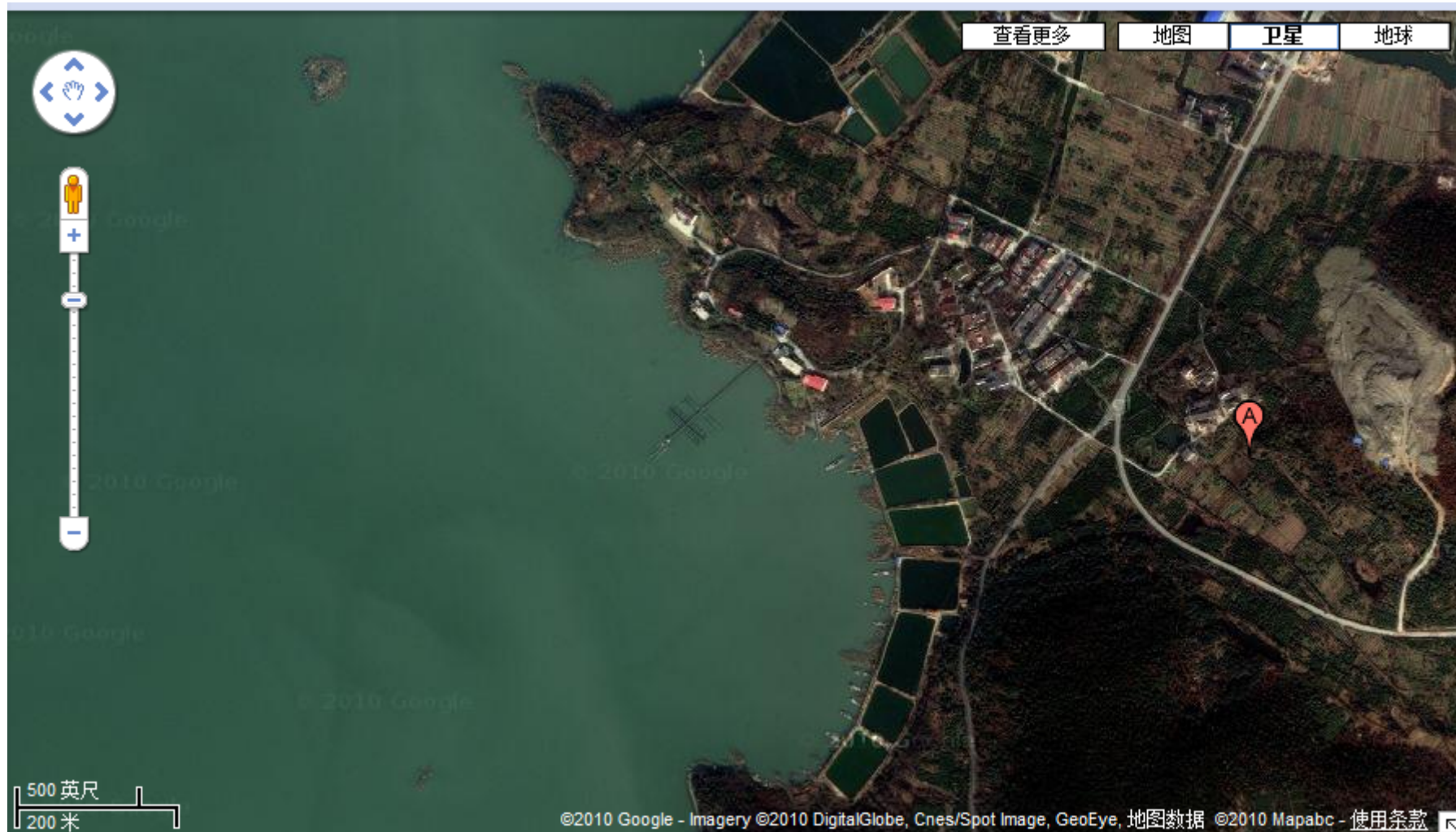


Our Experimental Site

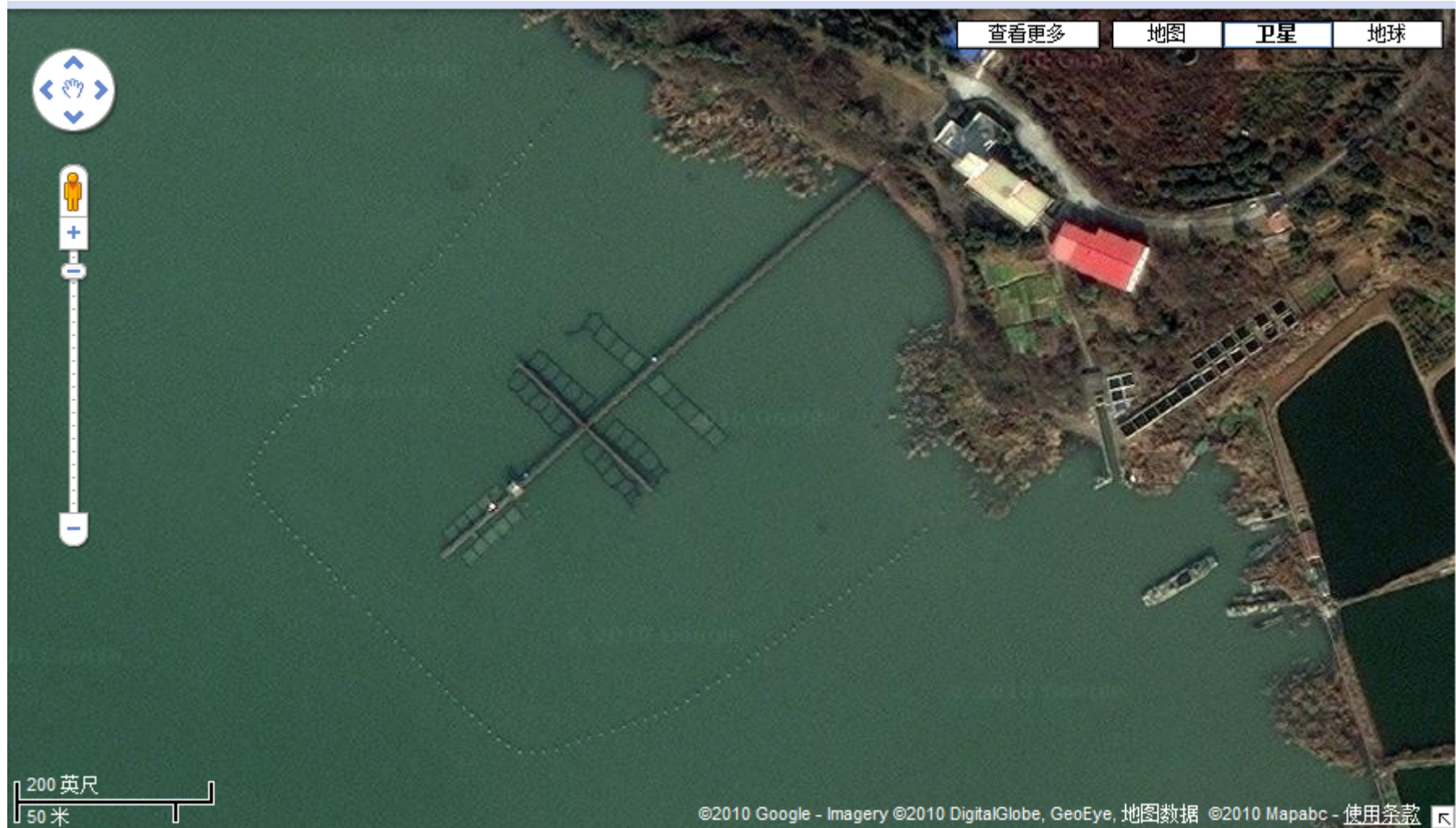


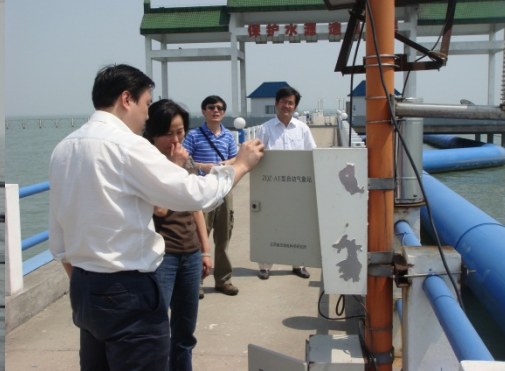
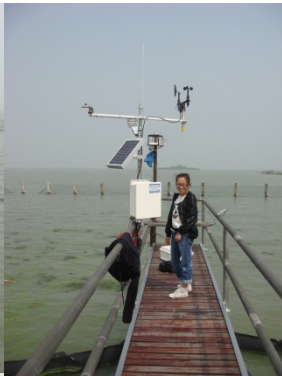
Good fetch: 180~315° south and west wind

Our Experimental Site



Our Experimental Site





Thank you!